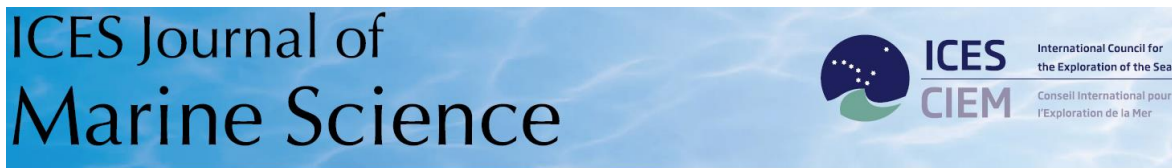


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## Contribution to the Symposium: *'Ecosystem studies of subarctic and Arctic seas'*

### Using natural analogues to investigate the effects of climate change and ocean acidification on Northern ecosystems

Samuel S. P. Rastrick<sup>1\*</sup>, Helen Graham<sup>2</sup>, Kumiko Azetsu-Scott<sup>3</sup>, Piero Calosi<sup>4</sup>, Melissa Chierici<sup>5,6</sup>, Agneta Fransson<sup>7</sup>, Haakon Hop<sup>7,8</sup>, Jason Hall-Spencer<sup>9</sup>, Marco Milazzo<sup>10</sup>, Peter Thor<sup>7</sup>, and Tina Kutti<sup>1</sup>

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3 <sup>1</sup>Institute of Marine Research, PO Box 1870 Nordnes, 5870 Bergen, Norway. <sup>2</sup>Ocean

4 Bergen, Espelandvegen 232, Blomsterdalen, Norway. <sup>3</sup>Department of Fisheries and

5 Oceans, Bedford Institute of Oceanography, Dartmouth, Nova Scotia, Canada

6 <sup>4</sup>Département de Biologie, Chimie et Géographie, Université du Québec à Rimouski, 300

7 Allée des Ursulines, Rimouski, QC G5L 3A1, Canada. <sup>5</sup>Institute of Marine Research, Fram

8 Centre, 9007 Tromsø, Norway <sup>6</sup>University Centre in Svalbard, Longyearbyen, Norway.

9 <sup>6</sup>Norwegian Polar Institute, Fram Centre, 9296 Tromsø, Norway. <sup>7</sup> Department of Arctic

10 and Marine Biology, Faculty of Biosciences, Fisheries and Economics, UiT The Arctic

11 University of Norway, N-9037 Tromsø, Norway. <sup>8</sup>Marine Biology and Ecology Research

12 Centre, School of Marine Science and Engineering, Plymouth University, Drake Circus,

13 Plymouth, Devon PL4 8AA, UK. <sup>9</sup>Department of Earth and Marine Science, Università

14 degli studi di Palermo, CoNISMa, Via Archirafi 20, I-90123 Palermo, Italy.

15

16 \*Corresponding author: [samuel.rastrick@imr.no](mailto:samuel.rastrick@imr.no),

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18 **Running Head:** Analogues of climate change in northern ecosystems

19

20 **Abstract**

21 Northern oceans are in a state of rapid transition. Still, our knowledge on the likely effects

22 of climate change and ocean acidification on key species in the food web, functionally

23 important habitats and the structure of ~~polar-Arctic and sub-Arctic~~ ecosystems is limited

24 and based mainly on short-term ~~single-species~~ laboratory studies on single species. This

25 review ~~demonstrates~~ discusses how tropical and temperate natural analogues to ~~ocean~~

26 ~~acidification~~ carbonate chemistry drivers, such as CO<sub>2</sub> vents, have been used to ~~move~~

27 further our knowledge on the sensitivity of biological systems to predicted climate change,

28 and thus assess the capacity of different ~~tropical and temperate~~ species to show long-term

29 ~~acclimate~~ acclimation and adaptation to elevated levels of pCO<sub>2</sub>. Natural analogues ~~and has~~

30 have also provided the means to scale-up from single-species responses to community and

31 ecosystem level responses. However, to date the application of such approaches is limited

32 in high latitude systems. A range of Arctic and sub-Arctic sites, including CO<sub>2</sub> vents,

33 methane cold seeps, estuaries, up-welling areas and polar fronts, that encompass gradients

34 of pH, carbonate saturation state and alkalinity, are suggested for future high latitude, *in-*

35 *situ* ocean acidification research. It is recommended that combinations of monitoring of the

36 chemical oceanography, laboratory studies, monitoring of the chemical oceanography,

37 observational and experimental (in situ and laboratory) studies of organisms around these

38 natural analogues be used to attain ~~good~~ better predictions of ~~future-the impact~~ effects of

39 ocean acidification and climate change on high latitude species and ecosystems.

40

| 41

## 42 Introduction

43 Rising levels of CO<sub>2</sub> in the atmosphere are causing worldwide modification of seawater  
44 carbonate chemistry, with gradual reductions in pH and carbonate ion (CO<sub>3</sub><sup>2-</sup>)  
45 availability, in a process known as ocean acidification (OA) ([Caldeira and Wickett, 2003](#);  
46 [IPCC, 2014](#)). The Arctic Ocean is particularly sensitive to OA, already having a naturally  
47 low pH and CaCO<sub>3</sub> saturation caused by the large freshwater content, which may increase  
48 due to warming and associated ocean and tundra ice melt and increased river runoff  
49 (Chierici and Fransson, 2009; Chierici et al., 2016). The relatively cold water in these  
50 regions also causes a high solubility of CO<sub>2</sub>, which could exacerbate OA in the future.  
51 Increased ice melt and river runoff not only effects the solubility of CO<sub>2</sub> via alterations in  
52 salinity, but depending on the geology of the runoffs terrestrial catchment can also  
53 inference total alkalinity (e.g. McGrath et al., 2016). The northern Norwegian Sea is a one  
54 region where a large part of the anthropogenic CO<sub>2</sub> has been absorbed during cooling of the  
55 warmer Atlantic water transported north along the Norwegian coast ([e.g. Olsen et al.,](#)  
56 [2006](#)). This has resulted in decreased pH and calcium carbonate (CaCO<sub>3</sub>) saturation, as well  
57 as shoaling of the saturation horizon. In fact, recent observations show that the pH decrease  
58 in the Norwegian Sea is occurring at one of the highest rates globally (Chierici et al., 2017;  
59 [Jones et al., 2018](#)). ~~The Arctic Ocean is particularly sensitive for OA, since it already has a~~  
60 ~~naturally low pH and CaCO<sub>3</sub> saturation caused by the large freshwater content ocean~~  
61 ~~(Chierici and Fransson, 2009; Chierici et al., 2016). The relatively cold water in these~~  
62 ~~regions causes a high solubility of CO<sub>2</sub>, which could progress OA in the future.~~ The Arctic  
63 It is one of the first ~~areas~~ regions already being affected by a rapid expansion in carbonate  
64 undersaturation (Olafsson et al., 2009; AMAP 2013; Qi et al., 2017). ~~The polar ocean and~~  
65 is experiencing one of the most rapid shifts in biogeographic boundaries on the planet due

66 to ~~polar and tundra rapid warming-ice melt~~ coupled with ~~ice melt rapid warming~~ and  
67 ~~acidifications decrease in alkalinity—these all affecting the ecology of marine organisms.~~  
68 Whilst rapid adaptation and borealization of the benthic flora are expected (Brodie et al.,  
69 2014), some benthic animals in the region may often have limited physiological plasticity  
70 which may can reduce their chances of survival (e.g. Rastrick and Whiteley, 2011; 2013;  
71 Calosi et al., 2017).

72 Ocean acidification in Arctic and sub-Arctic seas may have negative effects on  
73 pelagic species such as the copepod *Calanus glacialis* (Thor et al., 2017) and the shell-  
74 bearing pteropod *Limacina helicina* (Bednarsek et al., 2017) that constitute the main prey  
75 items for a ~~large~~ variety of larger zooplankton, juvenile fish, baleen whales and seabirds.  
76 Acidified seawater has also been suggested to impact coralline algae (Brodie et al., 2014)  
77 and cold-water corals (Jackson et al., 2014), which that form extensive biogenic habitats  
78 around northern Norway. To date our best predictions of habitat change in the North  
79 Atlantic are that warming will depleted kelp forests in the south and that ocean acidification  
80 will compromise maerl reefs in the north (Brodie et al., 2014). However, these predictions  
81 are based on a range of published laboratory experiments on signal species in isolation and  
82 so lack an understanding of the complex effects of interactions between species that can  
83 only be studied in naturally assembled systems. This isSuch habitat change is expected to  
84 impact nursery and brood stock areas for commercially important molluscs and fish  
85 (Branch et al., 2013; Sunday et al., 2017). It is predicted that cephalopods and crustaceans  
86 will be mostly unaffected by elevated pCO<sub>2</sub> levels expected by the end of the century, in  
87 contrast to shelled molluscs will be negatively affected (Branch et al., 2013). However,  
88 again these predictions are based on studies that cannot comprehend how the responses of

89 these key species to elevated pCO<sub>2</sub> may be modulated by and in turn modulate wider  
90 community level change.

91 Although a number of studies have shown adult fish to be relatively resistant to  
92 elevated pCO<sub>2</sub>, they many suffer neurological impairment (Milazzo et al., 2016). ~~and a~~  
93 recent study showed increased mortality in Atlantic cod larvae exposed to elevated pCO<sub>2</sub>,  
94 potentially resulting in reduced recruitment to the stock (Stiasny et al., 2016). Our current  
95 understanding of key processes driving the ~~response~~ responses of northern commercially  
96 important species and ecosystems to climate change is, ~~however,~~ limited. In addition, since  
97 the majority of studies conducted so far have been ~~in vitro~~ laboratory, short-term, rapid  
98 perturbation experiments on isolated elements of ~~the~~ ecosystems (e.g. Agnalt et al., 2013;  
99 Andersen et al., 2013; Bednarsek et al., 2017). Most studies have used set stable (but  
100 unrealistic) pCO<sub>2</sub> levels and focus on the responses of organisms that are separated from  
101 their natural suite of competitors, predators, parasites and facilitators, and experiment  
102 conducted are ofteoften too short-term to reveal how organisms may adapt/acclimatise.  
103 Thus, it is difficult to extrapolate from ~~such~~ individual level responses studies to larger  
104 ecological scales, ~~as these are generally such as~~ Population-level effects, but this has also  
105 been have to date been studied done either by applying laboratory rates (e.g. Stiasny et al.,  
106 2016) or with integrated models for effects of ocean warming and acidification  
107 (Koenigstein et al., 2018). ~~the w~~ However it extremely difficult to validate such modals in a  
108 natural context. ~~too short term to reveal how organisms may adapt/acclimatise,~~ have use in  
109 most cases often set stable (but unrealistic) eady pCO<sub>2</sub> levels (which are unrealistic) and use  
110 focus on the responses of organisms that are separated from their natural suite of  
111 competitors, predators, parasites and facilitators.

112 One approach to study ~~the ecosystems' responses to effects of predicted~~ future  
113 ~~predicted~~ chronic ~~elevations-increases~~ in mean  $p\text{CO}_2$ , as well as associated increases in  
114 acute fluctuations in carbonate chemistry due to freshwater run off, is to use natural  
115 analogues. Such analogues for future predicted OA have included ~~so far~~ volcanic vent sites  
116 (e.g. Hall-Spencer et al., 2008; ~~Kroeker et al., 2011, 2013~~), up-welling of deep  $\text{CO}_2$  rich  
117 water (Manzello et al., 2014) and temperate systems which present a mosaic of alkalinity  
118 and pH conditions (Thomsen et al., 2010).

119 These natural analogues provide an opportunity to simultaneously investigate  
120 changes in community structure (Hall-Spencer et al., 2008; Kroeker et al., 2011, 2013) and  
121 the capacity for physiological adaptation/acclimatisation of species in responses to elevated  
122  $p\text{CO}_2$  and low  ~~$\text{CO}_3^{2-}$ -carbonate ion~~ concentration (Calosi et al., 2013a; Harvey et al., 2016).  
123 In addition, broader evolutionary responses to past and future  $p\text{CO}_2$  changes can be  
124 addressed (Garilli et al., 2015). They have also been used to investigate the importance of  
125 natural variability in carbonate chemistry on the mechanisms that set ~~or limit~~ the  
126 distribution of species (Small et al., 2015) and to investigate the effect of multiple stressors  
127 ~~on the ecological performance and distribution of species~~ in naturally fluctuating  
128 environments (Thomsen et al., 2010; Kroeker et al., 2016). ~~However, to date studies~~  
129 ~~utilising natural analogues are limited to temperate and tropical systems.~~

130 Recent studies have identified how shifting boundaries of water and ice in ~~a high-~~  
131 latitude ~~systems-glacial fjord~~ create mosaics of ~~seawater~~ total alkalinity and pH conditions  
132 (Fransson et al., 2015; 2016), which could be used as natural analogues to investigate the  
133 effects of climate change and OA on the physiology, ecology and distribution of flora and  
134 fauna in northern ecosystems. However, to date, despite the potential for natural OA  
135 analogues at higher latitudes, such studies are limited to the Baltic (Thomsen et al., 2010).

136 Consequently, the aim of this paper review is to explore the possible use of natural  
137 analogues for investigating the effects of future changes in carbonate chemistry on northern  
138 species and ecosystems. Herein, we discuss: 1) the advantages and challenges of using  
139 natural analogues to investigate physiological, ecological and evolutionary effects of  
140 climate change and OA, drawing on temperate and tropical studies highlighting the arctic as  
141 an under-studied region; 2) Challenges of using natural analogues in more studied other  
142 ecosystems temperate and tropical to explore both regions, moving from studies of  
143 individual and to multiple carbonate chemistry drivers stressors; 3) how such analogue  
144 approaches could be modified for chemical oceanography and possible natural analogues  
145 that could be used in Arctic and sub-Arctic ecosystems given the present understanding of  
146 chemical oceanography in this region; and 4) what key target habitats and species in  
147 northern ecosystems could be studied using such analogues.

148

149 **The advantages and challenges of using natural analogues to investigate physiological,**  
150 **evolutionary and ecological and evolutionary effects of climate change**

151 Our knowledge of the biological effects occurring under anticipated changes of  
152 ocean chemistry is primarily informed by laboratory experiments. Such studies are  
153 informative, as they enable us to identify the effect of one or a few variables on many  
154 processes such as reproduction, development, physiology and behaviour of different  
155 organisms. However, most laboratory studies are carried out on a single generation of  
156 species in isolation and during short-term exposure to stress, neglecting many processes  
157 involving species in the wild, such as intra- and interspecific interactions and trans-  
158 generational adaptation, or nutrition supply and fluctuations in environmental parameters  
159 fluctuations (Barry et al., 2010). Research using natural analogues allows the investigation



160 of ~~the~~ chronically exposure to ~~ed populations to~~ elevated levels of dissolved  $p\text{CO}_2$ , in  
161 natural populations, thus providing a means by which to scale-up from physiological; to  
162 ecological processes with further extrapolation; -to evolutionary processes that by which  
163 ~~what is known about the effects of~~ OA could structure and modify the community and  
164 ecosystem levels.

165 A shift in community structure and composition favouring algal assemblages over reef  
166 forming species, for example, is consistently documented for chronically exposed benthic  
167 species along natural  $p\text{CO}_2$  gradients (Sunday et al., 2017). However,  $\text{CO}_2$  seep sites used  
168 in such studies are not perfect analogues of future change. A challenge being that within  
169 volcanic seep gradients variability in carbonate chemistry is often more rapid and of a  
170 higher amplitude than is expected due to the effects of gradually rising atmospheric levels  
171 of  $\text{CO}_2$  (Kerrison et al., 2011; Kroeker et al., 2011). This high rate of change in seawater  
172 chemistry is known to be especially detrimental to coralline algae, for example, which are  
173 less sensitive to gradual change (Kamenos et al., 2013). Such patterns in community  
174 structure across natural  $\text{CO}_2$  gradients are likely driven by a combination of direct  
175 physiological effects on habitat forming species (such as, elevated costs of maintaining  
176 homeostats and calcification) and indirect effects such as involving changes in the energy  
177 available in feed, competition, predation and habitat structure. The natural distribution of  
178 ecosystem engineers, such as grazing sea urchins across volcanic  $p\text{CO}_2$  gradients in Italy, is  
179 suggested to be controlled by the physiological capacity to maintain acid-base homeostasis  
180 (Calosi et al., 2013b; Small et al 2015). These species-specific responses will depend on  
181 both the physiological limits and energetic costs of the specific mechanisms employed  
182 (Small et al., 2015). Changes in the energetic costs of maintaining homeostasis across  $p\text{CO}_2$   
183 gradients can lead to energetic trade-offs that effect-impact growth and reproduction

184 affecting populations (Harvey et al., 2016). This demonstrates how natural analogue studies  
185 can be used to scale from direct physiological impacts at the individual level to population  
186 level responses that may have implications at the community level ~~implications~~ within  
187 naturally assembled systems. Fleshy non-reef forming macroalgae, for example, show  
188 marked increases under naturally elevated  $p\text{CO}_2$  levels and appear capable of exploiting  
189  $p\text{CO}_2$  *via* photosynthesis (Cornwall et al., 2017a). This allows seaweed to out-compete  
190 calcifying reef species (e.g. corals and vermetids), whose performance is lowered by  
191 dissolution and increased energetic costs associated ~~to~~ with calcification (Milazzo et al.,  
192 2014). In addition to this, key groups (e.g., crustose coralline algae) that trigger the  
193 recruitment of reef forming species, can be out-competed or cannot survive (Diaz-Pulido et  
194 al., 2011; Milazzo et al., 2014). Overall such responses lead to ecosystem shifts from  
195 calcareous reefs to algal-dominated habitats (Diaz-Pulido et al., 2011).

196 Carbon dioxide seep research carried out in temperate, sub-tropical and tropical  
197 regions has revealed that responses of benthic habitats vary regionally, ~~as no total coral~~  
198 ~~cover reduction~~ Reduction but in species ~~composition~~ diversity of corals, but not in total  
199 cover of corals, has been ~~changes~~ observed in Papua New Guinea and Palau (Fabricius et  
200 al., 2011; Barkley et al., 2015; ~~Fabricius et al., 2014~~), while a shift from hard to soft corals  
201 has been documented in Japan (Inoue et al., 2013) and from corals to macroalgae in  
202 Northern Mariana Islands (Enochs et al., 2015). Hall-Spencer et al. (2008) and Linares et  
203 al. (2015) describe a transition from communities with abundant calcareous organisms to  
204 communities lacking scleractinian corals and a significant reduction in coralline algae off  
205  $\text{CO}_2$  vents in the Mediterranean Sea.

206 A common feature shared by biological systems-chronically exposed to elevated  
207  $\text{CO}_2$ , and demonstrated uniquely by the use of natural analogues, is the general loss of

208 habitat complexity and ~~usually of~~ the associated diversity. ~~Acidification~~ Decreasing pH  
209 conditions also leads to ecological shifts, such as changes in competitive dominance and  
210 habitat provisioning (Sunday et al., 2017). For instance, when kelp/macroalgae and  
211 seagrass habitats shift to low-relief turf-algal habitats off CO<sub>2</sub> seeps in New Zealand and  
212 Italy, this causes loss of fish predators and increase in prey fish species ~~release~~, even  
213 though their antipredator responses were compromised (Nagelkerken et al., 2015).  
214 However, such patterns of response are species-specific (Cattano et al., 2017).

215 In addition to, facilitating the investigation of the interaction between individual,  
216 population and community responses to chronic changes in carbonate chemistry within  
217 naturally assembled systems, natural analogues provide a suitable test bed for studies of  
218 adaptation to many different environmental drivers. While lab studies of adaptation require  
219 multi-generation incubation periods, which is impractical in longer lived species and  
220 species with complex life history patterns, the real strength of natural analogous approaches  
221 is the possibility of tests on populations pre-adapted through many generations. Thus,  
222 natural analogous They may be specifically used to test the potential for a species'  
223 adaptation to future environmental changes, and how such adaptation may potentially  
224 rescue species from local or global extinction. Evolutionary rescue from environmental  
225 changes may be facilitated ~~by either by~~ mutation, evolutionary selection ~~adaptation~~, or  
226 migration (Bell and Collins, 2008; Bell and Gonzalez, 2009). However, in long-lived  
227 multicellular organisms, mutations progress at rates much lower than what would be  
228 needed to facilitate adaptation to ~~Mutation rates are far slower than~~ present-day  
229 environmental changes. ~~in metazoans. However~~ This said, genetically ~~-~~ based phenotypic  
230 variation throughout a species' distribution range provides a constant supply of  
231 physiological alternatives or new possibilities upon which selection can operate to facilitate

232 adaptation to the new set of conditions that occur (Foo et al., 2012; ~~Kelly and Hofmann,~~  
233 ~~2013; Munday et al., 2013~~ Ecology Letters; Reusch, 2014; Sunday et al., 2014; Calosi et al.,  
234 2016). Therefore, adaptation from selection in extant genetic diversity, as well as migration  
235 among locally adapted populations may effectively decrease the risk of extinction in the  
236 face of climate change and OA. Both of these processes can be tested using natural  
237 analogues, given certain conditions. The main prerequisite for differential adaptation  
238 among natural analogues is that populations should be genetically isolated. In addition, The  
239 rates of selection should not be matched by the rate of gene flow among populations  
240 (Kawecki and Ebert, 2004; Savolainen et al., 2013). Therefore, a potential challenge in  
241 using natural gradients is to establish possible connectivity patterns between experimental  
242 populations/sub-populations. Many benthic organisms proliferate in the larval stages and  
243 establishing models of larval dispersal may be important (Cowen et al. 2007). Genetic  
244 isolation may be obtained at a distance of 2-5 times the larval dispersal range (Palumbi  
245 2003). However genetic differences may be compromised with the migration of only a few  
246 individuals per generation (Cowen and Sponaugle 2009, Slatkin 1993). Consequently,  
247 differences in physiological responses of individuals across natural analogues should be  
248 accompanied by an understanding of the phylogenetic relationship between  
249 populations/sub-populations (Hill and Bucklin 2001; Calosi et al., 2013) or if possible the  
250 allelic heterogeneity in genes related to the physiological processes. In addressing these  
251 challenges analogue selection is critical with many seep systems presently used showing  
252 localised steep gradients in carbonate chemistry over distances of 10s to 100s of meters,  
253 allowing organisms to move in or recruit from outside. This may hinder genetic adaptation  
254 (Calosi et al., 2013; Harvey et al., 2016; Turner et al., 2016) and cause short-term  
255 physiological shocks to organisms that are suddenly exposed to hypercapnia (Small et al.,

256 ~~2015). While our knowledge on the adaptation potential to climate change and OA has been~~  
257 ~~very limited (Dam, 2013; Munday et al., 2013; Reusch, 2014; Sunday et al., 2014; Kelly and~~  
258 ~~Hofmann, 2013), recent studies have brought more insight (Calosi et al., 2016). In~~ Despite  
259 this in benthic animals, adaptation to e.g. OA has been recently demonstrated. For instance,  
260 the polychaete *Platynereis dumerilii* has been shown to adapt to chronic and elevated levels  
261 of  $p\text{CO}_2$ . Populations occupying  $\text{CO}_2$  vent sites on the Italian coast are physiologically and  
262 genetically different from nearby populations that experience low  $p\text{CO}_2$  (Calosi et al.,  
263 2013). However, no adaptation to high  $\text{CO}_2$  conditions was found in the calcifying  
264 spirorbid worm *Simplaria* sp. following a putative multi-year exposure to high OA  
265 conditions no evidence for adaptation to high  $\text{CO}_2$  conditions where found (Turner et al.,  
266 2016). In the sea urchin *Centrostephanus rodgersii*, the presence of tolerant genotypes  
267 indicates a potential to adapt to concurrent warming and acidification (Foo et al., 2012).  
268 Thus ability to adapt to OA conditions does not appear ubiquitous in marine metazoans (see  
269 also e.g., Calosi et al. 2016). Using natural analogues to testing adaptation in planktonic  
270 species may be a specific challenge as ~~But even planktonic species may experience~~  
271 ~~differential adaptation throughout their distribution range. While~~ low genetic  
272 differentiation and efficient dispersal of ~~propagules all life stages may were previously~~  
273 ~~thought to~~ hinder local adaptation, However, recent studies have shown pelagic copepods to  
274 be dispersed in distinct populations locally with little genetic interchange ~~throughout~~  
275 around the Northern hemisphere (Nelson et al., 2009; Unal and Bucklin, 2010; Yebra et al.,  
276 2011; but see Weydmann et al., 2016), although some species have large oceanic  
277 distributions (Wassmann et al., 2015). ~~Coupled with~~ Moreover large population sizes of  
278 planktonic organisms, as opposed to less abundant longer-lived benthic organisms, this can

279 promote effective selection with an increased potential for local adaptation (Charlesworth,  
280 2009; Peijnenburg and Goetze, 2013).

281 Reciprocal transplant tests between locations characterized by different  
282 environmental regimes should be employed to ascertain that differences in phenotypes  
283 among locations are not caused by phenotypic plasticity but occur as a result of adaptation  
284 (Niewiarowski and Roosenburg, 1993). Transplant individuals, once transferred to the new  
285 environment, should show the same phenotype as individuals found in the environment.  
286 ~~such transplant tests, phenotypic plasticity results in transplant individuals attaining native~~  
287 ~~individuals' phenotype.~~ Any deviation from this outcome signals that differences stems  
288 from either ~~pure genetic~~ adaptation or ~~or~~ transgenerational effects ~~caused by~~; for instance  
289 reversible epigenetic or post-transcriptional changes (Bonduriansky et al., 2012). However,  
290 whilst evidence for phenotypic differences are essential, so is the aAnalyses of allelic  
291 differences between populations. This will ~~is needed to~~ in fact further ascertain that  
292 observed differences are indeed genetically based (Calosi et al., 2013; Pespeni et al., 2013;  
293 De Wit et al., 2015). Obviously most of genetic variation involved in adaptation to  
294 environmental changes lies within expressed sequences (Jones et al., 2012). Reverse  
295 transcription sequencing and transcriptomic treatment ~~is~~ are a powerful tools to assess the  
296 connection between physiological differences and allelic changes in expressed genes.  
297 Aligning allelic differences to observed differences in targeted physiological processes can  
298 be used for a hypothesis-testing strategy to detect cellular targets of adaptation to ocean  
299 acidification (De Wit et al., 2015; Thor and Dupont, 2015). This novel approach seems  
300 promising for future studies of effects of environmental changes in ecologically important  
301 non-model organisms.

302

### 303 Use of natural analogues to explore both individual and multiple stressors

304 Hall-Spencer et al. (2008) initiated the use of volcanic CO<sub>2</sub> seeps as analogues for  
305 future OA off the Island of Ischia in the Mediterranean. More recently, similar volcanic  
306 sites ~~have been~~were~~have been~~ investigated, for example, in Papua New Guinea (Lamare et  
307 al., 2016), in the subtropical North East Atlantic reefs (La Palma Island) (Hernández et al.,  
308 2016), in the temperate Pacific Ocean in Japan (Shikine Island) (Agostini et al., 2015), and  
309 Bay of Plenty, New Zealand (Burrell et al., 2015). A natural CO<sub>2</sub> seep was also found in  
310 Salt Dome Juist in the North Sea (McGinnis et al., 2011), although OA studies ~~are~~ have not  
311 yet been conducted there. Most ~~ocean acidification~~OA research using such sites has  
312 focused on sessile benthos; to retain greater control over exposure elevated pCO<sub>2</sub>  
313 conditions to acidified conditions. Organisms in the water column can be exposed to abrupt  
314 changes in seawater carbonate chemistry as they move towards or away from the gas vents  
315 (Kerrison et al., 2011; Kroeker et al., 2011). That said, experiments and observations in the  
316 water column around CO<sub>2</sub> seeps have ~~shown~~ been useful in demonstrating shifts in  
317 naturally assembled plankton communities and greater sensitivity of calcifying plankton  
318 even if their exposure may be transitory (Johnson et al., 2013; Ziveri et al., 2014). For  
319 example, ~~that~~ large diatoms tend to grow well in the acidified conditions at sites with  
320 elevated pCO<sub>2</sub> with a significant increase in chlorophyll concentrations and diatom  
321 abundance observed, however cyanobacteria were reportedly uninfected (Johnson et al.,  
322 2013); In contrast ~~whereas~~ coccolithophores show a decrease in cell concentrations and  
323 diversity as calcite saturation decreased from 6.4 to <1, with malformed *Emiliania huxleyi*  
324 observed at the highest pCO<sub>2</sub> levels do not (Ziveri et al., 2014). Such studies are not a  
325 perfect representation of future pelagic systems due to migration in and out of the vent  
326 system. However, they facilitate a greater understanding of how plankton communities that

327 are critical to marine primary production and biogeochemical cycling naturally assemble  
328 under elevated  $p\text{CO}_2$  conditions in a way not possible using laboratory experiments. -In  
329 temperate and tropical conditions invertebrate recruitment to the benthos is severely  
330 disrupted in the high-elevated  $p\text{CO}_2$  conditions found at volcanic seeps, although these  
331 observations may overestimate the impact of ~~ocean acidification~~OA since drifting larvae  
332 from normal seawater conditions are suddenly exposed to acidified waters  
333 characterised by elevated  $p\text{CO}_2$  (Cigliano et al., 2010; Allen et al., 2016). More realistic,  
334 perhaps, are observed impacts of elevated  $p\text{CO}_2$  acidification on pelagic organisms that  
335 spend long periods ~~of time~~ in areas with naturally acidified conditions. The reproductive  
336 behaviour of nesting fish is affected at  $\text{CO}_2$  seeps (Milazzo et al., 2016). Zooplankton and  
337 fish that use coral habitat are also much less abundant in elevated  $p\text{CO}_2$  acidified  
338 conditions, which ~~is thought to be due to the fact that~~ may be because acidification-elevated  
339  $p\text{CO}_2$  reduces coral reef complexity (Smith et al., 2016).

340         The steep gradients in pH and carbonate saturation that occur next to volcanic seeps  
341 consistently have marked effects on the abundance and distribution of sessile calcified  
342 organisms. Studies show how most coralline algae are highly susceptible to these naturally  
343 acidified conditions as their high-Mg calcite skeletons are easily corroded or damaged  
344 (Kamenos et al., 2013; Martin and Hall-Spencer, 2017). Although some species are less  
345 sensitive, ~~it has been shown that~~ very few survive where aragonite saturation levels fall  
346 below 1 for even brief periods of time (Martin et al., 2008; Fabricius et al., 2015), and in  
347 these acidified conditions they are easily outcompeted by fleshy algae (Kamenos et al.,  
348 2016~~7~~). The sensitivity of calcified organisms to to acidified low pH conditions depends  
349 on how well they are able to protect their skeletons or shells. Vent studies have shown that  
350 some corals, for example, can calcify and grow well in acidified-low pH conditions, if they



351 have enough food, as their skeletons are covered in protective tissue (Rodolfo-Metalpa et  
352 al., 2011; 2015). The same is true of certain molluscs, whilst others grow in a dwarf form  
353 to more easily meet the metabolic costs of hypercapnia (Garilli et al., 2015). OA is in ~~its~~  
354 self a multiple stressor, with the effects of low carbonate saturation, low pH and  
355 increased DIC working together to shift the outcome of competition within marine  
356 communities, often to the benefit of ~~r~~ weed-like or ~~r~~-selected species such as turf algae  
357 (Connell and Russell, 2010) and uncalcified polychaetes (Gambi et al., 2016).

358 Whilst a great deal has been learnt from CO<sub>2</sub> seeps worldwide about the likely long-  
359 term ecological effects of acidificationOA, these systems are not perfect analogues for the  
360 future. For example, acidification is happening alongside local or regional warming. One  
361 way to address this is to compare CO<sub>2</sub> seep systems in different biogeographic thermal  
362 regimes and regionszones to ~~find out~~reveal whether related organisms show consistent  
363 responses to elevated CO<sub>2</sub> despite differences in temperature (Johnson et al., 2012).

364 Another approach is to take advantage of marine heat waves to assess the combined stress  
365 of high CO<sub>2</sub> and elevated seawater temperature (Rodolfo-Metalpa et al., 2010; 2011). ~~A~~  
366 ~~drawback with volcanic seep systems is that variability in carbonate chemistry is much~~  
367 ~~more rapid and of a higher amplitude than is expected due to the effects of gradually rising~~  
368 ~~atmospheric levels of CO<sub>2</sub> (Kerrison et al., 2011; Kroeker et al., 2011). This high rate of~~  
369 ~~change in seawater chemistry is known to be especially detrimental to coralline algae, for~~  
370 ~~example, which are less sensitive to gradual change (Kamenos et al., 2013).~~ Confounding  
371 factors may also be present at CO<sub>2</sub> seeps, so a great deal of care is needed to tease apart the  
372 effects of multiple stressors such as low oxygen or elevated metal toxicity (Vizzini et al.,  
373 2013). Consequently, seep sites are selected to reduce confounding factors and focus only  
374 on changes in pCO<sub>2</sub>, with a major challenge being how to develop analogue studies from

375 ~~investigating single to multiple carbonate chemistry drivers.~~ Finally, ~~seep systems are~~  
376 ~~usually very localised with steep gradients in carbonate chemistry over distances of 10s to~~  
377 ~~100s of meters, so organisms may move in or recruit from outside, . This may hinder~~  
378 ~~hindering genetic adaptation (Calosi et al., 2013; Harvey et al., 2015; Lucey et al., 2016~~  
379 ~~Evol. APpl.) and causing cause short term physiological shocks to organisms that are~~  
380 ~~suddenly exposed to hypercapnia (Small et al., 2015).~~

381         Despite ~~drawbacks~~ challenges, data collected at CO<sub>2</sub> seeps currently provide us with  
382 the best window we have into the future for assessing the risks of acidification to marine  
383 communities, habitats and ecosystems. ~~Unfortunately, suitable CO<sub>2</sub>-seeps sites are~~  
384 ~~uncommon.~~ However other analogues for future conditions are available that retain the  
385 advantages of seep systems, i.e. ~~(chronic exposure of entire marine communities to low~~  
386 ~~carbonate saturation and high DIC.)~~ but lack the disadvantages of rapid variations in  
387 carbonate saturation, steep gradients in pH and DIC. These analogues may also lack  
388 potentially confounding factors such as hypoxia or H<sub>2</sub>S toxicity and the influx of organisms  
389 that are exposed to a sudden increase in CO<sub>2</sub> levels as they recruit, swim or drift through  
390 these open systems.

391         Sites where the biology affects the CO<sub>2</sub> of the environment could be used similarly  
392 to the seep sites. Due to carbonate production on coral reefs, some atolls, lagoons and  
393 barrier reefs can exhibit consistently higher surface pCO<sub>2</sub> values than those in offshore  
394 waters (Suzuki and Kawahata, 2004). Sea grass beds, on the other hand, can reduce CO<sub>2</sub>  
395 levels causing increases in pH and aragonite saturation (Unsworth et al., 2012). There are  
396 also natural CO<sub>2</sub> gradients formed in areas of large-scale seaweed culture (for example in  
397 China). These systems typically experience large temporal changes in carbonate chemistry  
398 due to water movement and diurnal photosynthesis patterns. These temporal shifts are even

399 more pronounced in the intertidal environments, for example in tide pools, where  $p\text{CO}_2$  can  
400 reach 1800  $\mu\text{atm}$  due to the respiration of the inhabitants (Andersson et al., 2013).

401 Similarly, mangrove environments (in Bermuda) ~~were~~have been shown to ~~ex~~perience  
402 large fluctuations in carbonate chemistry parameters ( $\text{pH}$ ,  $p\text{CO}_2$  and  $\Omega_{\text{Ca}}$ ) over daily cycles  
403 (e.g. due to groundwater input) with  $p\text{CO}_2$  levels varying from 500 to 4200  $\mu\text{atm}$   
404 (Andersson et al., 2013). It is still important to characterise these habitats because of the  
405 effects ~~the~~that varying carbonate parameters can have on organisms. However, these large  
406 daily fluctuations may make it difficult to disentangle the effects of singular factors (e.g.  
407 salinity or temperature) as they may co-vary.

408 Carbonate chemistry gradients can produce a mosaic pattern where species and  
409 communities ~~are naturally~~may be adapted to diverse conditions, ~~putatively~~, over multiple  
410 generations. This could allow for the study of the effects of multiple stressors, and allow  
411 work on natural analogues to move forward and encompass multiple drivers of climate  
412 change ~~drivers~~. Naturally overlapping carbonate chemistry gradients have been described in  
413 marine habitats throughout the world. One of the most common areas for them to occur in  
414 is coastal regions and estuaries, where several interacting biotic and abiotic stressors are  
415 occurring within the environment. Large estuaries may have limited gene flow between  
416 populations allowing for the investigation of the capacity for adaption. There are several  
417 examples of mosaics identified in estuarine environments globally. One such system of  
418 interacting gradients (total alkalinity, DIC and salinity) has been identified around the coast  
419 of Ireland. Due to the underlying limestone bedrock of river catchment areas, runoff to  
420 coastal and estuarine areas can have high total alkalinity (TA) values (2864  $\mu\text{mol Kg}^{-1}$  TA  
421 at salinity 15.86, Shannon plumes) creating these TA gradients, which expand outwards  
422 from the rivers and estuaries (McGrath et al., 2015). Interacting gradients of salinity, with

423 O<sub>2</sub>, DIC and pH, have been highlighted in the Strait of Georgia (British  
424 Columbia, Canada). The water masses connecting the semi-enclosed estuary of the Fraser  
425 river to the outer shelf of the Pacific Ocean are subject to the effects of large scale  
426 upwelling and downwelling on the outer coast leading to different trends in pH and  
427 aragonite saturation ~~differing~~ in the tributaries (Haro and Juan de Fuca) feeding the Strait of  
428 Georgia ~~and causing as well as~~ overlapping gradients ~~to occur~~ (Ianson et al., 2016). The use  
429 of these dynamic coastal environments ~~can may~~ offer insight in-to the long-term effects  
430 and adaptation of organisms to changing oceanic conditions. Unlike the CO<sub>2</sub> seep systems  
431 these estuarine habitats cover a greater spatial range and ~~due to this these, thus,~~ carbonate  
432 chemistry gradients are likely to be less steep possibly limiting gene flow among  
433 populations to a rate lower than selection, facilitating adaptation (Kawecki and Ebert, 2004;  
434 Savolainen et al., 2013):-

435 To date there has been limited research using mosaics to look at the chronic and/or  
436 acute effects of carbonate chemistry drivers -on organismal performance and fitness. One of  
437 such study examined the effect of salinity and pH gradients on the spatio-temporal variation  
438 in communities of phytoplankton in Sungai Brunei and Brunei bay estuarine system which  
439 identified the highest algal densities occurring at the highest pH (pH 7.8) and highest  
440 salinity (salinity 27) and the number of taxa present decreased with decreasing pH  
441 (Majewska et al., 2017). A similar salinity and pH gradient located in the same estuary  
442 (Sungai Brunei estuary) has also been used to determine correlations between shell  
443 dissolution in the gastropod *Thais gradate* and ~~water chemical properties~~ carbonate  
444 chemistry drivers (pH, salinity, calcium concentration). At decreased pH ( $6.83 \pm 0.39$ ),  
445 lower calcium and low salinity ( $13.50 \pm 5.49$ ), individuals exhibited higher levels of shell  
446 erosion and smaller standardised shell length compared to individuals acclimatised to

447 higher pH ( $8.02 \pm 0.15$ ) and salinity ( $27.17 \pm 3.0$ ). ~~Although not pronounced as smaller-~~  
448 ~~scale habitats (tide pools, mangroves),~~ ~~†~~These estuarine scale gradients can ~~still be~~ ~~can still~~  
449 ~~be~~ subject to temporal changes based on daily cycles (e.g. salinity, temperature) although  
450 these are usually not as pronounced as smaller scale habitats such as tide pools and  
451 mangroves.

452 There are however larger scale mosaics, ~~such as continental upwelling,~~ which offer  
453 the opportunity to study the effects of changing carbonate chemistry at continental scales.  
454 For example, at the continental shelf upwelling on the western coast of North America from  
455 central Canada in northern Mexico, where although seasonal upwelling of low pH water is  
456 a natural phenomenon the extent of the affected area is increasing with OA (Feely et al.,  
457 2008). One such mosaic has been identified in the California current system where, due to  
458 ~~the water source and~~ constant upwelling, a spatial mosaic of carbonate chemistry is formed.  
459 ~~Previously this analogue has been used to analyse the potential interactions between~~ This  
460 large (1280 km of coastline) environmental mosaic of overlapping temperature, pH  
461 carbonate chemistry and chlorophyll-*a* gradients has been used to investigate how multiple  
462 interacting stressors associated with global change and the combined impact they may have  
463 ~~on~~ the growth and predation vulnerability of the California blue mussel *Mytilus*  
464 *californianus* (Kroeker et al., 2016). This study demonstrated that dynamic environments  
465 with frequent exposure to low pH seawater and consistent food showed highest growth rate  
466 and lowest predation vulnerability. Whereas, growth was limited in areas with frequent low  
467 pH and less consistent food availability and with extremes in low tide body temperature  
468 (Kroeker et al., 2016). Other potential sites for large scale mosaics can be found in the  
469 eastern Pacific Ocean -the Arabian sea where overlapping gradients in temperature, oxygen,  
470 and carbonate chemistry have been used to assess the relative influence of these climate

471 change associated drivers on macrofaunal diversity and evenness (Sperling et al., 2016). In  
472 this study oxygen levels expand most of the variation in species diversity, while, carbonate  
473 chemistry was the best explanatory variable in the Arabian sea it explains less of the  
474 variation in the Pacific. , and the north-western coast of America (Feely et al., 2008).

475

## 476 **Chemical oceanography and possible natural analogues in Arctic and sub-Arctic** 477 **ecosystems**

478 The potential is large for investigating natural analogues of multiple carbonate  
479 chemistry drivers to better understand the possible effects of climate change on the  
480 physiological, ecological and adaptive-evolutionary capacity of individuals and  
481 communities is clear. However, to date such sites are-have been exclusively located  
482 identified in temperate and tropical seas, despite the importance and possibly greater  
483 sensitivity of sub-Arctic and Arctic regions to OA. Potential sub-Arctic volcanic CO<sub>2</sub>  
484 seepvent sites may be found in Iceland. Other possible study sites may be found at varying  
485 depths around Jan Mayen in vent fields between the Greenland and Norwegian Seas, and  
486 off the west coast of Spitsbergen. The consequence of OA on ecosystems and  
487 biogeochemical cycling in this area is unknown. The increased acidity due to CO<sub>2</sub>,  
488 associated with corrosive volcanic input and increased Arctic water masses, and potential  
489 vulnerability of key species, calls for the investigation of pH and other carbonate chemistry  
490 variables, as well as vulnerable species and processes (Fauchald et al., 2014). Other  
491 gradients in carbonate chemistry may be associated with large stocks of methane hydrate  
492 that exist in the sub-Arctic and the Arctic. Warming of seawater can destabilize methane  
493 hydrate and releasing methane (CH<sub>4</sub>) to the water column. A recent study demonstrated  
494 that the majority of this CH<sub>4</sub> is oxidized in the water column to CO<sub>2</sub> without escaping to the

495 atmosphere (Myhre et al., 2016). Release of CH<sub>4</sub> from the sediment and subsequent  
496 oxidation to CO<sub>2</sub> in the water column were modelled and attributed to prolonged OA  
497 (Boudreau, et al, 2015). Methane cold seep sites are found extensively in the polar oceans,  
498 for example, East Siberian Shelf (Shakhova et al., 2017), off Svalbard (Myhre et al., 2016),  
499 Baffin Island Shelf (Punshon et al., 2014), and the Barents Sea (Hong et al., 2017; Serov et  
500 al., 2017). Although, cold seeps may provide sites for studying natural gradients in  
501 carbonate chemistry in polar oceans investigations of the chemical oceanography, including  
502 confounding effects of CH<sub>4</sub> methane, and the biology of such sites are in their infancy and  
503 their depth and remoteness makes them logistically difficult. Other mosaics in  
504 carbonate chemistry drivers that may drive plankton communities can be found marginal  
505 ice zone of the Arctic ocean (e.g. Lewis et al., 2013; Barber et.al., 2015; Jule et al., 2018).

506 Coastal regions in high-latitude oceans are influenced by freshwater such as rivers,  
507 glaciers and sea ice melt. Each freshwater source has different carbonate chemistry and  
508 affects local acidification states (Chierici and Fransson, 2009). In the Hudson Bay  
509 (Canada), the rivers flowing into the western bay have higher alkalinity than those in the  
510 eastern bay due to the difference in watershed geology. Watershed of the western rivers is  
511 composed of lime stones, while that of eastern rivers is basaltic. Consequently, carbonate  
512 saturation states and pH in the western bay is higher than those in the eastern bay (Azetsu-  
513 Scott et al., 2014). Similar gradients of pH, carbonate saturation state and alkalinity are  
514 observed in the Spitsbergen/Svalbard fjord system in different years due to drainage of  
515 meltwater containing calcareous minerals from the bedrock (Fransson et al., 2015; 2016).  
516 These heterogeneous chemical environments can be studied in detail together with  
517 individual species and ecosystem responses. In addition to freshwater runoff from glacial  
518 melt water in Arctic fjords, there can also be upwelling near the glacier fronts affecting the

519 biogeochemical gradients and ecosystem (e.g. Straneo et al. 2012; Lydersen et al. 2014;  
520 Meire et al. 2015). Chemical gradients are also observed in polar fronts where warm,  
521 Atlantic water meets cold, polar water, such as in the Barents Sea and Fram Strait (Chierici  
522 et al. 2016). To use naturally occurring gradients as analogues to study effects of climate  
523 change and OA to marine ecosystems, we also need to clarify how the sites represent the  
524 future marine environments beyond the carbonate system. For example, some sites may  
525 provide pH gradients with an extreme dissolved oxygen range, or pH may fluctuate outside  
526 that predicted for OA. Also, pH variation can be attributed to other factors than the  
527 carbonate system, such as H<sub>2</sub>S at volcanic seeps. Variable responses by organisms and  
528 ecosystems in chosen sites may be the results of multiple environmental drivers such as,  
529 nutrients, salinity, temperature, dissolved oxygen and the TA of terrestrial runoff from  
530 glaciers and rivers. To address these questions, it is necessary to understand temporal and  
531 spatial variability of carbonate chemistry, controlling mechanisms of OA and chemical  
532 compositions of seawater such as heavy metals, nutrients, hydrogen sulfide and methane.  
533 Although many confounding factors associated with chemical composition can be  
534 controlled by monitoring, modeling and careful site selection, ~~others factors~~ such as  
535 salinity, TA and temperature may be more variable particular in coastal environments.

536

### 537 *Target habitats and species*

538 In all ecosystems, there are some species or taxonomic groups that play a  
539 disproportionately important functional role, e.g. as prey, as habitat engineers or in the  
540 recycling of nutrients. If such species suddenly ~~become more or less~~ increase or decrease in  
541 abundance, (due to changes in the abiotic environment), the community structure might  
542 change or even regime shifts ~~to entire~~ ecosystems may occur in the ecosystem (Kortsch et



543 al., 2012; Fossheim et al., 2015). Polar ocean ecosystems in general are characterized by  
544 having pathways of energy flow from lower to higher trophic levels dominated by a small  
545 number of species (Murphy et al., 2016), e.g. the copepod *Calanus glacialis* and the shell-  
546 bearing pteropod *Limacina helicina* constitute the main food source for a large variety of  
547 large zooplankton, juvenile fish, baleen whales and birds in northern ecosystems (Last,  
548 1980; Lowry, 1993; Karnovsky et al., 2003; Hop and Gjørseter, 2013).

549 Shell-bearing ~~Pteropods~~ pteropods are generally thought to be extremely sensitive to  
550 reduced pH because they exert little control over the pH and carbonate chemistry of their  
551 calcifying fluid (Ries, 2012; Manno et al., 2017). Indeed, shell dissolution has been  
552 observed in some ocean areas where aragonite saturation state is around 1 (Bednaršek et al.,  
553 2012). However, recent studies from the California Current Ecosystem, (that naturally  
554 experiences undersaturated waters with respect to aragonite due to seasonal up-welling)  
555 have shown that ~~the pteropod~~ *Limacina helicina* individuals that originate from areas with  
556 naturally lower aragonite saturation state exhibit a higher survival rate as  
557 ~~compared~~ compared to individuals from areas with a higher aragonite saturation state when  
558 exposed to high  $p\text{CO}_2$  waters in laboratory (Bednarsek et al., 2017). Similarly, different  
559 responses to ~~ocean acidification~~ OA were observed in three geographically separated  
560 populations of *C. glacialis* (Thor et al., 2017). ~~Copepodites~~ Copepodids collected from  
561 Kongsfjorden and Billefjorden (Svalbard) showed severe reductions in ingestion and an  
562 increased metabolic cost when exposed to high  $p\text{CO}_2$  waters in laboratory whereas no  
563 effects were observed in ~~copepodites~~ copepodids collected from Disko Bay (west  
564 Greenland). ~~Studies on another copepod species, *Pseudocalanus acuspes*, show fast~~  
565 ~~adaption to OA and that physiological differences among locally isolated populations may~~  
566 ~~occur as a result of local adaptation (Thor and Dupont 2015; De Wit et al., 2015).~~ In

567 ~~addition. Northern populations of the gastropod *Littorina littorea* in Northern Norway~~  
568 ~~were the only ones along the entire geographical range of distribution of this species to~~  
569 ~~showed no ability to modulate their metabolic rates in response to the exposure to OA~~  
570 ~~conditions. In addition, they showed the greatest shift in metabolomic profiles and shell~~  
571 ~~dissolution (Calosi et al., 2017). This may suggest that northern populations of warm-~~  
572 ~~adapted species may are not be able to deal with low pH regimes typical of subpolar and~~  
573 ~~polar environment, being adapted to warmer higher pH seawater conditions.~~

574         Gradients in pH, carbonate saturation state and alkalinity occurs in several  
575 Svalbard/Spitsbergen fjord systems where pteropods and copepods could be studied (e.g.  
576 Fransson et al., 2015; 2016). However as pelagic organisms can be exposed to abrupt  
577 changes in seawater carbonate chemistry as they move across smaller scale natural  
578 gradients (e.g. Kroeker et al., 2011), larger scale open-ocean gradients in carbonate  
579 chemistry, such as in the Barents Sea and Fram Strait polar fronts (Chierici et al., 2016)  
580 may be more suitable as naturally laboratories to investigate the sensitivity of different  
581 zooplankton populations to OA, compared to smaller special gradients.  
582 Polar coastal ecosystems support extensive biogenic habitats in the form of coralline algae  
583 beds and cold-water coral gardens and reefs. Coralline algae and cold-water corals are  
584 considered ecosystem engineers as they play a prominent role in the polar carbonate cycle  
585 and act as habitats for thousands of other benthic species (see e.g. [Freiwald and Henrich,](#)  
586 [1994; Teichert 2014;](#) Henry and Roberts, 2017; ~~Teichert 2014; Freiwald and Henrich,~~  
587 ~~1994~~). It is expected that any changes in the secondary production of these benthic ~~habitat~~  
588 ~~habitat~~-building taxa would have scaling effects on ~~the~~ polar food webs. Studies assessing  
589 the effects of OA on them are therefore of utmost interest for both fisheries and  
590 environmental management. In some ~~of the~~ Arctic fjords (in both northern Norway and

591 Greenland), cold-water corals have been observed where chemical gradients can be  
592 investigated (~~T. Kutti, pers. comm.~~). Extensive cold-water corals reefs and coral gardens  
593 occur in northern Norway, Iceland, western Greenland and eastern Canada (Buhl-  
594 Mortensen et al., 2016). Coralline algae beds have been described from the euphotic zone  
595 in the fjords of northern Norway and Svalbard (Brodie et al., 2014; Teichert and Freiwald,  
596 2014), eastern and western Greenland (Jørgensbye and Halfar, 2017) and eastern Canada  
597 (Halfar et al., 2013). Laboratory studies have yielded contrasting results regarding coralline  
598 algae and cold-water coral performance under elevated temperature ~~elevation~~ and ~~increased~~  
599  $p\text{CO}_2$  (~~see~~ e.g. Ragazzola et al., 2013; Büscher et al., 2017 and references therein). Several  
600 factors, such as seasonality, food availability and species-specific traits, have been ~~used~~  
601 proposed to explain the discrepancy between studies. What is clear is that many coralline  
602 algae and cold-water coral species have the ability to upregulate pH at the site of  
603 calcification and thus continue to grow even in corrosive water (McCulloch et al., 2012;  
604 Cornwall et al., 2017b); ~~;~~ however, often at a reduced rate. Furthermore, the dissolution of  
605 the unprotected skeleton (forming the main part of the coralline algae beds and cold-water  
606 coral reefs) is always ~~enhanced-faster~~ under corrosive conditions. Short-term laboratory  
607 studies, however, cannot test the ability of these species, ~~that-which~~ have very slow growth  
608 and long generation times, to acclimate and/or adapt to ocean warming and acidification.  
609 ~~Coralline algae beds have been described from the euphotic zone in the fjords of northern~~  
610 ~~Norway and Svalbard (Brodie et al., 2014; Teichert and Freiwald, 2014), eastern and~~  
611 ~~western Greenland (Jørgensbye and Halfar, 2017) and eastern Canada (Halfar et al., 2013).~~  
612 ~~Extensive cold-water corals reefs and coral gardens occur in northern Norway, Iceland,~~  
613 ~~western Greenland and eastern Canada (Buhl-Mortensen et al., 2016).~~ Many of these sites  
614 with corals or coralline algae show mosaics in carbonate chemistry, temperature and food

615 availability (driven by up-welling and glacial melt) and could hence be used to assess the  
616 long-term ability of these species to acclimate/adapt to higher CO<sub>2</sub> and temperature. It is  
617 plausible that, similarly to pteropods and copepods, geographically separated populations  
618 might respond differently to exposure to high CO<sub>2</sub> waters. However, such future studies  
619 would require long-term data on carbonate chemistry, temperature, and perhaps even  
620 ecological parameters such as food availability/seston concentration and quality. Although at  
621 specific locations there is a growing understanding of natural spatial and temporal gradients  
622 in carbonate chemistry for example Hudson Bay (Azetsu-Scott et al., 2014) and  
623 Kongsfjorden (Fransson et al., 2015) an increased monitoring and modelling efforts will be  
624 required. In general, the Arctic Ocean and adjacent seas (Barents Sea and Norwegian Sea)  
625 are little investigated with regard to carbonate chemistry. The longest, northernmost time  
626 series study site is located in the Iceland Sea (Olafsson et al., 2009). On other sites such as  
627 Svalbard fjords, Fram Strait and the Barents Sea repeat measurements along hydrography  
628 sections was initiated in 2011. Biological and physical parameters have been sampled  
629 systematically in the Barents Sea since 1950's through the Institute of Marine Research  
630 (Norway) ecosystem surveys, The Fram Strait has annual cruises with carbonate chemistry  
631 since 2011 and in recent years also pteropod collection. There are also several moorings  
632 with proposed plans to extend the sensor systems to include carbonate chemistry for data  
633 collection throughout the year. However, in the Arctic Ocean, data still relies on research  
634 projects and distributed biological observatories (DBO) such as in the Chukchi Sea. DBO's  
635 which may be extended to include more parts of the Arctic Ocean.

636

637

638

## 639 **Conclusions**

640 High latitude oceans are particularly sensitive to climate change, due to naturally low pH  
641 and CaCO<sub>3</sub> saturation caused by changing freshwater content, and high solubility of CO<sub>2</sub>  
642 due to relatively low temperatures ([AMAP AOA 2013](#); Chierici and Fransson, 2009;  
643 Chierici et al. 2016). However, despite the northern oceans being in a state of rapid  
644 transition, little is known about the possible effects of predicted OA on ecologically  
645 important species and communities. What is known comes mainly from relatively  
646 short-term laboratory incubations to [isolated](#) carbonate chemistry drivers (i.e. pCO<sub>2</sub>,  
647 salinity or temperature independently). These studies give little opportunity to investigate  
648 either the bidirectional interactions between responses at the individual, population and  
649 community levels, or the possible capacity for transgenerational adaptation of species to  
650 climate change drivers. However, in a number of tropical and temperate ecosystems these  
651 questions have been successfully addressed by using natural gradients in carbonate chemistry  
652 as analogues for predicted OA, where species have potentially adapted/acclimatised over  
653 multiple generations within naturally assembled communities. Despite the clear potential  
654 for using such analogues also at high latitudes, this remains to be investigated.

655 To date such studies in tropical and temperate ecosystems have mostly focused on  
656 volcanic CO<sub>2</sub> vent sites ~~that~~[which](#) are often carefully selected to avoid confounding factors  
657 such as H<sub>2</sub>S, heavy metals, salinity or temperature. Potential Arctic volcanic CO<sub>2</sub> vent sites  
658 are reported in the Jan Mayen vent fields and off the west coast of Spitsbergen.  
659 Furthermore, methane cold seeps are found extensively in northern oceans. However, the  
660 chemical oceanography, including possible confounding effects of methane, H<sub>2</sub>S or heavy  
661 metals needs further investigation. The depth and remoteness of these sites also makes them

662 logistically ~~default~~difficult~~demanding~~. The research value strength of CO<sub>2</sub> vent sites  
663 studied to date is their relatively simple experimental design with relatively short gradients  
664 and limited confounding drivers. However, this can also be a limitation due to increased  
665 gene flow across shorter gradients making it difficult to investigate adaptive capacity, and a  
666 lack of opportunity to study the interaction of multiple drivers. Consequently, a number of  
667 recent studies have investigated larger scale sites such as areas of coastal upwelling and  
668 estuaries where gradients in carbonate chemistry (e.g. salinity, total alkalinity, pCO<sub>2</sub> and  
669 temperature) intersect to produce an environmental mosaic. This has the potential to allow  
670 work on natural analogues to move forward and encompass multiple climate change  
671 drivers. At high latitudes, such mosaics occur at different scales from coastal regions and  
672 fjords influenced by fluctuations in temperature, salinity and total alkalinity of freshwater  
673 (e.g. rivers, glaciers and sea ice melt) to polar fronts where warm, Atlantic water meets  
674 cold, Arctic water (e.g. Barents Sea and Fram Strait). Although beyond the scope here, it  
675 should be noted that similar mosaics in carbonate chemistry drivers have been described in  
676 the Southern Ocean. With seasonal wintertime minimum in carbonate ion concentration  
677 south of the Antarctic Polar Front, which is predicted to lead to aragonite undersaturation  
678 when atmospheric CO<sub>2</sub> levels reach above 450 ppm (McNeil and Matear, 2008). In the  
679 Ross sea surface pH varies from (7.890–8.033) with the highest values in Terra Nova Bay  
680 and Ross Sea polynyas. Intrusion of the Circumpolar Deep Water can also lead to low pH  
681 values (7.969±0.025) in the Ross Sea shelf area (Rivaró et al., 2014). Such natural  
682 fluctuations in carbonate saturation have been shown to infer the shell dissolution in  
683 the Southern Ocean pteropod *Limacina helicina Antarctica* (Bednaršek et al., 2012).

684 It is important that appropriate analogue sites are selected to investigate particularly  
685 key species, habitats or processes. For example, smaller scale analogues within fjords may

686 be used to investigate the effects of multiple interacting drivers on key benthic biogenic  
687 habitats (e.g. coralline algae and cold-water corals) scaling between individual  
688 physiological effects and community level responses. Larger scale gradients in big fjords  
689 and coastal upwelling areas are more suitable for species with a higher mobility. Larger  
690 analogues allow for target populations to be more genetically isolated and are therefore  
691 more suitable for investigating the capacity for local and regional adaptation. When  
692 investigating ecologically important pelagic organisms (e.g. zooplankton) that can be  
693 exposed to abrupt changes in seawater carbonate chemistry as they move across smaller  
694 scale natural gradients, larger scale open-ocean analogues may have some advantage: (e.g.  
695 gradients in carbonate chemistry across polar fronts in the Barents Sea and Fram Strait).  
696 However, as the scale of gradients increases it becomes more difficult to understand the  
697 role of individual confounding drivers, which are often connected and correlated. Although  
698 many confounding factors can be controlled for by using appropriate monitoring and  
699 modeling to inform site selection, some carbonate chemistry drivers in more complex  
700 systems are likely to co-vary. Consequently, as we move toward a more realistic  
701 understanding of multiple carbonate drivers in the field, it is not suggested that such studies  
702 replace laboratory incubations but rather that both methods complement each other with  
703 natural analogues used to validate responses observed in the laboratory and laboratory  
704 incubations used to disentangle confounding drivers observed in natural systems. It is clear,  
705 that if used and selected appropriately to fit the question, and combined with monitoring  
706 and modeling of the chemical oceanography, natural analogues will be a powerful tool to  
707 achieve a better understanding of the possible effects of climate change on high latitude  
708 species, communities and ecosystems.

709

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