# Title: Abiotic and biotic interactions in the diffusive boundary layer of kelp blades create a potential refuge from ocean acidification

Authors: Fanny NOISETTE, Catriona HURD

Affiliation: Institute for Marine and Antarctic Studies, University of Tasmania, 20 Castray Esplanade, Battery Point, 7004 TAS Hobart, Australia

**Corresponding author:** 

Fanny Noisette

Email: <u>fanny.noisette@live.fr</u>

Phone : +33 (0)6 03 10 84 80 (FR) / +1 (0)4 99 24 38 63 (AU)

Type of paper: Original Research Article

**Short title:** Living in the DBL of kelp blades

#### 1 Summary

Seaweeds are able to modify the chemical environment at their surface, in a micro-zone called
 the diffusive boundary layer (DBL), via their metabolic processes controlled by light intensity.
 Depending the thickness of the DBL, sessile invertebrates such as calcifying bryozoans or
 tube-forming polychaetes living on the surface of the blades can be affected by the chemical
 variations occurring in this micro-layer. Especially in the context of ocean acidification, these
 microhabitats might be considered as a refuge from lower pH, because during the day
 photosynthesis temporarily raises the pH to values higher than in the mainstream seawater.

9 2. We assessed the thickness and the characteristics of the DBL at two pH levels (today's average surface ocean pH 8.1 and a reduced pH predicted for the end of the century, pH 7.7) and seawater flows (slow, 0.5 and fast, > 8 cm s<sup>-1</sup>) on *Ecklonia radiata* (kelp) blades. Oxygen and pH profiles from the blade surface to the mainstream seawater were measured with O<sub>2</sub> and pH microsensors for both bare blades and blades colonized by the bryozoan *Membranipora membranacea*.

3. The DBL was thicker in slow compared to fast flow and the presence of bryozoans increased
the DBL thickness and shaped the DBL gradient in dark conditions. Net production was
increased in the low pH condition, increasing the amount of oxygen in the DBL in both bare
and epiphytized blades. This increase drove the daily pH fluctuations at the blade surface,
shifting them towards higher values compared to today's pH. The presence of bryozoans led
to lower oxygen concentrations in the DBL and more complex pH fluctuations at the blade
surface, particularly at pH 7.7.

4. Overall this study, based on microprofiles, shows that, in slow flow, DBL micro-environments
at the surface of the kelps may constitute a refuge from ocean acidification with pH values
higher than those of the mainstream seawater. For calcifying organisms, it could also represent
training ground for harsh conditions, with broad daily pH and oxygen fluctuations. These

chemical micro-environments, biologically shaped by the macrophytes, are of great interest
for the resilience of coastal ecosystems in the context of global change.

28

Key-words (5): bryozoan, ecomechanics, epiphytism, hydrodynamic, seaweed, microhabitat,
pH

31

## 32 Introduction

33 Seaweeds are not only the dominant primary producers in coastal waters but they are important bioengineers that are able to modify their surrounding environment (Hurd et al. 34 2014). In particular, brown macroalgae forming broad communities such as the Laminariales 35 and Fucales act as ecosystem engineers by influencing physical factors including seawater 36 velocity (Gaylord et al. 2007; Rosman et al. 2010) and light penetration (Reed & Foster 1984), 37 and the chemical characteristics of the mainstream seawater, including carbonate chemistry 38 (Delille, Borges & Delille 2009; Cornwall et al. 2013a; Hendriks et al. 2014), oxygen and 39 nutrient availability (Frieder et al. 2012; Saderne et al. 2015). Seaweed communities can 40 modulate their surrounding chemical environment on seasonal and diel cycles (Delille et al. 41 2000; Saderne, Fietzek & Herman 2013), however the most rapid and variable fluctuations in 42 chemical parameters occur in the microenvironments formed at the surface of the seaweeds, in 43 a zone called the diffusive boundary layer (DBL, also termed the concentration boundary layer) 44 (Hurd et al. 2014; Wahl, Saderne & Sawall 2016). 45

The DBL is a discrete micro-layer at the surface of many organisms (e.g. Kühl *et al.* 1995; de Beer *et al.* 2000), including all primary producers (e.g. Koch 1994; Brodersen *et al.* 2015), that buffers them from the mainstream seawater (Vogel 1999). This specific layer is formed when a fluid flows over a solid, such as a kelp blade. The no-slip condition creates a

region of viscously-dominated laminar flow at the seaweed surface. In this thin laminar layer, 50 51 movement of ions and molecules is by molecular diffusion, and the metabolic activity of the organism results in a concentration gradient due to the uptake and release of dissolved 52 substances to and from the organism's surface (Vogel 1999; Hurd 2000). Fluctuations observed 53 in these microhabitats at the surface of seaweed are mainly driven by their photosynthesis and 54 respiration processes under the control of light (Sand-Jensen, Revsbech & Jörgensen 1985; 55 56 Hurd et al. 2011; Cornwall et al. 2013b ; Hofmann, Koch & de Beer 2016). Thus, metabolic activity affects the micro-chemical environment of the DBL which differs from that in the 57 mainstream seawater just micrometres away (Hurd 2015), with implications for the alga itself 58 59 and all the other small organisms living on the blades.

The importance of DBLs in controlling the availability and transfer of nutrients and 60 metabolites to and from algal surfaces is well known (Raven 1997; Hurd 2000; De Beer & 61 Larkum 2001). The presence of a thick DBL formed under slow flows ( $< 2 \text{ cm sec}^{-1}$ ) may, for 62 example, decrease the uptake of dissolved inorganic carbon and nutrients (Wheeler 1980; 63 Kregting et al. 2011). However, these microenvironements are beneficial for other processes 64 such as timing of gamete release (Pearson, Serrão & Brawley 1998) or keeping antifouling 65 agents at their surface (Dworjanyn, De Nys & Steinberg 2006). Interestingly, different 66 organisms such as bacteria, diatoms, larvae and spores live in this thin layer (Schaffelke 1999; 67 Wahl, Saderne & Sawall 2016). Some bigger and calcifying species such as bryozoans or tube 68 forming worms also settle on brown macroalgal blades (Saderne & Wahl 2013) and would be 69 submitted to the fluctuations occurring in these micro-habitats. In particular, bryozoans are 70 known to be epiphytic organisms which influence their algal substratum (Muñoz, Cancino & 71 72 Molina 1991; Hurd et al. 1994a), changing for example nutrient uptake at the seaweed surface (Hurd et al. 1994a; Hepburn & Hurd 2005; Hepburn, Frew & Hurd 2012). According to their 73 size, their structure and their direct contact with the algal substratum, they are typically sessile 74

organisms which would be affected by oxygen and pH fluctuations that occur in the DBL (Irwin
& Davenport 2002; Wahl, Saderne & Sawall 2016). Nevertheless, there is a knowledge gap on
the interactive effects that bryozoans can have with other abiotic factors of the surrounding
water and their role in the formation and gradient creation in the DBL has never been clearly
observed.

80 The different physical and chemical parameters of the environment surrounding macroalgae can affect the thickness of the DBL and the flux of dissolved substances, and thus 81 the concentration gradient which can vary from µm to mm (0.1 to 10 mm; Raven & Hurd 2012). 82 Flow rates directly impact the thickness of the DBL (Vogel 1999; Denny 2015) and variation 83 in metabolic processes can change the concentration gradients (Sand-Jensen, Revsbech & 84 Jörgensen 1985). In seaweeds, photosynthesis and respiration processes not only affect the 85 oxygen gradient of the DBL (e.g. Spilling et al. 2010) but they also change the pH due to the 86 uptake and release of DIC in light and dark, respectively (e.g. De Beer & Larkum 2001; Hurd 87 et al. 2011, Cornwall et al. 2013b ; Short, Pedersen & Kendrick 2015). This ability of 88 macrophytes to metabolically modify their local pH and the influencing effect of 89 90 hydrodynamics on DBL thickness have recently been put forward in the context of ocean acidification (OA) to understand how seaweeds, seagrasses and their associated organisms, 91 92 particularly calcifiers, will respond to the predicted decrease in seawater pH, (Hendriks et al. 2015, 2017; Hurd 2015; Wahl, Saderne & Sawall 2016; Cox et al. 2017). The term OA 93 describes the average decrease in the surface oceanic pH and the associated changes in seawater 94 carbonate system caused by increasing CO<sub>2</sub> release into the atmosphere since the beginning of 95 the industrial revolution (Gattuso & Hansson 2011). There is no doubt that OA is an ongoing 96 97 process and that the pH will be reduced by 0.1 to 0.3 pH units by the end of the century (Stocker et al. 2013), and it is therefore essential to find ways to increase the resilience of sensitive 98 species and discover temporal or spatial refuge from OA (Hurd 2015). 99

In this context, this study aims to better characterize the thickness and concentration 100 101 gradient of the DBL at the blade surface of the kelp Ecklonia radiata, (C.AGARDH) J.AGARDH 102 1848, one of the most common and widespread brown macroalgae in southern hemisphere coastal oceans, which forms large and dense seaweed communities (Steneck et al. 2002). As 103 104 the interactive effect of flow rate, mainstream seawater pH and epiphytism on DBL have never been characterized, we measured oxygen concentrations and pH values from the blade to the 105 106 mainstream seawater, in saturating light and dark conditions, in fast and slow flow, today's pH and that predicted for 2100 and in the presence or absence of the common bryozoan 107 Membranipora mambranacea, LINNAEUS 1767. We hypothesised that, 1. the DBL would be 108 109 thicker in slow flow because low velocity favours the increase of the DBL (Hurd 2000); 2. the DBL would also be thinner for blades with bryozoa because the ciliary motion of the zoides 110 may create turbulence disturbing the laminar properties of the DBL; 3. the DBL oxygen 111 gradient would be thicker in future OA compared to current pH conditions because the 112 productivity of fleshy algae is expected to increase under OA (Kroeker et al. 2013); 4. oxygen 113 concentration will be lower on blades colonized by bryozoans than on bare blade as bryozoan 114 colonies can shade the tissue decreasing the intensity received and so the photosynthetic 115 116 efficiency of the blades (Muñoz, Cancino & Molina 1991). Finally, knowledge of oxygen and 117 pH fluctuations in these different conditions may give some clues on if DBL microenvironment on kelp blades might be considered as refuge from OA because of the buffering effect of 118 seaweed metabolism to their surrounding pH, permitting temporal respite from very low 119 120 mainstream pH.

121

#### 122 Material and methods

#### 123 Sample collection and laboratory acclimation

Around 80 lateral blades from separate individuals of the kelp Ecklonia radiata were 124 125 collected using scissors, by snorkelling, in the Tinderbox reserve (-43.058177 S, 147.330749 E), close to Hobart (Tasmania, Australia) on 17<sup>th</sup> and 24<sup>th</sup> September 2015. Bare blades and 126 blades colonized by the bryozoan Membranipora membranacea were sampled each time. 127 Blades were similar in term of color and length, collected at the same depth ( $\approx$  3-4 m) and the 128 bryozoan patches upon blades had a diameter of  $3.1 \pm 0.1$  cm. Field collections were allowed 129 130 by permit provided by the Department of Primary Industries, Parks, Water and Environment of Tasmania (n°15115). Blades freshly cut were transferred to the laboratory in a cool box to keep 131 them at the field temperature (13.5°C). They were then placed in a temperature controlled room 132 133 (13°C), in 0.1 µm filtered/UV sterilized seawater bubbled with air, under low light conditions  $(50 \,\mu\text{mol photons m}^{-2} \,\text{s}^{-1})$ , until the beginning of the experiment, to acclimate them to laboratory 134 conditions. 135

136

#### 137 Experimental design

All the experiments were carried out after a 2-day acclimation period, between 2 and 5 138 days after collection. Experiments were conducted in a 46 L unidirectional recirculating flume 139 140 (see description in Hurd et al. 1994b) filled with 0.1 µm filtered/UV sterilized seawater to a depth of 15 cm. This type of unidirectional flume is considered as a standard tool to test the 141 interactions between organisms and their surrounding flow even if it cannot perfectly mimic 142 143 the hydrodynamics in the field (Vogel 1999). The flume was initially cleaned with a 1% sodium hydroxide (bleach) solution and neutralized with a 1% thiosulfate solution (Hurd et al. 1994b). 144 It was subsequently rinsed with tap water every day, at the end of the daily experiments, to 145 avoid any biofilm formation inside. The flume was filled with seawater at the beginning of each 146 day and 8 different *Ecklonia* blade samples were measured over the same day. 147

Four blade replicates were measured in each different combination of light, mainstream 148 pH, flow and epiphytic condition i.e. bare blade or presence of bryozoans (64 different blades 149 in total). Measurements were carried out in both the light, at a saturating irradiance of 120 µmol 150 photons m<sup>-2</sup> s<sup>-1</sup> provided by overhead lights (T8 840, 36W, Thorn Lighting, UK) and dark 151 conditions (<  $2 \mu$ mol photons m<sup>-2</sup> s<sup>-1</sup>). Irradiance was measured with a flat underwater quantum 152 sensor LI-250A (LI-COR, Lincoln, USA). Two pH conditions were tested, the current sea-153 154 surface average  $pH_T = 8.09$  (hereafter 8.1) and a lower pH level predicted for the end of the century in the worst-case scenario for ocean acidification,  $pH_T = 7.71$  (hereafter 7.7) (Stocker 155 et al. 2013). The pH was adjusted by bubbling with 100% CO<sub>2</sub> into the seawater until reaching 156 157 the expected level (Hurd et al. 2011). Light and pH treatments were randomized in order to have a single combination of light and pH per experimental day. In total, 8 days were necessary 158 to run the measurements for all the possible combined conditions. 159

160 Every day, two mainstream seawater velocities were randomly tested, a fast one, > 8 cm  $s^{-1}$  and a slow one, < 0.5 cm  $s^{-1}$ . The flow characteristics in the different regimes were 161 determined in the middle of the flume using a field 10 MHz Acoustic Doppler Velocimeter 162 (Sontek, San Diego, CA, USA) before the start of the experiment. These velocities were chosen 163 because they represent speeds at which the DBL can reach maximal and minimal values 164 165 (Hansen, Hondzo & Hurd 2011) and they are similar to those observed within kelp beds in the field (Jackson & Winant 1983; Gaylord et al. 2007; Kregting et al. 2011). Individual kelp blades 166 colonized or not by bryozoans were randomly assigned to the different flow treatments. Each 167 168 of the replicate blades were attached to an aluminium plate covered with plastic film so they could not move during the trials. Measurements were always made at the top of a crenulation, 169 170 in the middle of the blade, in the middle of bryozoan patch (for the bryozoans present condition), in the same position in the flume (Hurd & Pilditch 2011). 171

The characterization of the DBL, i.e. its thickness and the oxygen concentration and pH 174 gradient, was determined using a 50 µm O<sub>2</sub> micro-sensor and a 50 µm micro-pH electrode 175 coupled with a reference electrode (Unisense, Aarhus, Denmark), for 4 replicates in each 176 combined conditions of light, flow, bulk seawater pH and epiphytism (presence or absence of 177 178 bryozoans). The electrodes were attached to a Unisense MM33-2 manual micromanipulator (Unisense, Aarhus, Denmark). A minimum period of 30 min acclimation to each treatment 179 occurred prior to starting vertical profiles, allowing the DBL to form and stabilize. Profiles of 180 181 O<sub>2</sub> concentration and pH were taken by placing the microelectrodes at the surface of the blade and by sequentially increasing the height of the probes towards the mainstream seawater. The 182 measurements were made at 0.05, 0.1, 0.15, 0.2, 0.3, 0.4, 0.5, 0.6, 08, 1, 3, 5, 10 and 50 mm 183 (bulk seawater) above the blade surface, for 2 min each, logging O<sub>2</sub> and pH every second. 184 Values were then averaged over 30 seconds after the initial peak induced by the movement of 185 the electrode. 186

187 *DBL thickness:* DBL thickness was calculated for each  $O_2$  profile, on the raw values, 188 and was defined as the greatest height above the surface of the blade at which the concentration 189 of  $O_2$  was <1 % per 0.1 mm for four subsequent measurements (Hurd *et al.* 2011; Cornwall *et* 190 *al.* 2015).

191  $O_2$  profiles: To describe the profiles,  $O_2$  was first standardized by dividing the 192 concentration at any given profile location by the bulk seawater concentration measured at the 193 end of the profile. Profiles were then fitted and smoothed with an exponential equation: y = y0194 +  $\alpha \exp^{(-\beta x)}$  with y: the standardized  $O_2$  concentration, y0: a constant,  $\alpha$ : the oxygen 195 standardized concentration when x = 0,  $\beta$ : the rate of change x: the distance from the blade. *Interfacial oxygen fluxes*: net production in light and respiration in dark conditions were
 defined as interfacial oxygen fluxes (Hofmann, Koch & de Beer 2016), calculated from the raw
 concentration values of the profiles using Fick's first law (Revsbech & Jørgensen 1986):

$$J = -D \left(\frac{dc}{dx}\right)$$

with J: interfacial oxygen fluxes in  $\mu$ mol O<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>, D: diffusion coefficient of oxygen in seawater in m<sup>-2</sup> s<sup>-1</sup> (D = 1.66 10<sup>-9</sup> at 13.2°C and salinity 37, value calculated with the R package marelac), dc: the change in concentration in the DBL in  $\mu$ mol m<sup>-3</sup> and dx: the thickness of the DBL in m.

Data analyses: DBL thicknesses, interfacial fluxes and O<sub>2</sub> standardized concentrations 204 at the surface of the blade (0 mm) were analyzed using 3-way ANOVAs with pH, flow and 205 presence/absence of bryozoan (factor called "blade") as crossed orthogonal factors followed by 206 post-hoc SNK tests. A 3-way MANOVA with the same factors was used to compare the 207 estimated y0,  $\alpha$  and  $\beta$  coefficients of the curves fitted to the profiles. The analyses were run 208 independently for light and dark conditions, except for comparing DBL thicknesses in which 209 210 measurements in the light and dark conditions were combined because the DBL thickness is affected by physical factors as flow speed and roughness but not light. All statistical analyses 211 were performed using the R software, version 2.15.0 (R Core Team 2013) after the normality 212 and the homoscedasticity of the data have been checked. 213

pH profiles: pH values acquired along the profiles were first transformed in H<sup>+</sup> concentration with the equation  $[H^+] = 10^{-pH}$  (Riebesell *et al.* 2010). H<sup>+</sup> concentrations were then standardized as for oxygen, by dividing the concentration at any given profile location by the mainstream seawater concentration measured at the end of the profile. Standardized values where then converted back in pH values using the equation pH = -log  $[H^+]$  to obtain profiles of pH deviation from the mainstream pH in the DBL. Finally, linear regression wasused to explore the relationship between pH deviation from the mainstream value and the standardized oxygenconcentrations along the DBL.

222

### 223 Seawater parameters

During the eight experimental days, temperature, salinity, pH and nutrients were 224 measured three times a day, i.e. after flume filling, at midday and before flume draining. 225 Temperature was logged using a Precision Multi Digital thermometer (Testo, Lenzkirsh, 226 Germany), salinity was measured with a refractometer (HI96882, Hanna Instruments, 227 228 Woonsocket, RI, USA) and pH was determined immediately after sampling using a Orion star A111 pH meter coupled with a OrionRoss ultra pH/ATC triode (Thermo Scientific, Waltham, 229 MA, USA) and calibrated with 4, 7 and 10 pH<sub>NBS</sub> buffers. pH was standardized on the total 230 scale using TRIS buffers prepared as directed by (Dickson, Sabine & Christian 2007). Samples 231 for nutrients were frozen at -20°C until defrosted and analysed for nitrate (NO<sub>3</sub><sup>-</sup>) and phosphate 232 (PO4<sup>-</sup>) using a QuickChem 8500 series 2 Automated Ion Analyzer (Lachat Instrument, 233 Loveland, CO, USA). Water for total alkalinity (A<sub>T</sub>) was sampled once a day, pre-test showing 234 235 that alkalinity did not vary along the day. AT samples were poisoned with mercuric chloride 236 (0.02% vol/vol; Dickson, Sabine & Christian 2007) pending analyses, made later at the Australian National University, using an automatic built in-house titrator (consisting in a 5 mL 237 Tecan syringe pump (Cavro XCalibur Pump), a Pico USB controlled pH sensor, and a TPS pH 238 electrode). A<sub>T</sub> values were then calculated using the Gran technique (Dickson, Sabine & 239 Christian 2007). Seawater carbonate chemistry, i.e. CO<sub>2</sub> partial pressure (pCO<sub>2</sub>), carbon 240 dioxide, bicarbonate and carbonate ions concentrations ( $[CO_2], [HCO_3^-]$  and  $[CO_3^{2-}]$ ), dissolved 241 inorganic carbon (DIC), and the saturation state of aragonite and calcite ( $\Omega_{Ar}$ ,  $\Omega_{Ca}$ ) were 242 calculated from temperature, salinity, pH and A<sub>T</sub> values measured for each pH level with the 243

Seacarb package (Gattuso, Epitalon & Lavigne 2016) using constants from Mehrbach et al.
(1973) refitted by Dickson & Millero (1987).

246

247 **Results** 

248 Seawater parameters

Average chemical parameters of the mainstream seawater at pH 8.1 and pH 7.7 conditions are given in Table 1. Standard errors of the values showed that the algae did not strongly affect the bulk seawater chemistry, meaning that the physico-chemical parameters did not vary during each daily experiment: the seawater was renewed each day. Seawater nutrients of 3.4  $\mu$ M for nitrate and ammonium and 1.3  $\mu$ M for phosphate were within the normal range for spring (Hepburn, Hurd & Frew 2006). According to the nutrient concentrations and DIC values carbon or nitrogen limitation for photosynthesis are unlikely during our measurements.

256

#### 257 DBL thickness

The average thickness of the DBL varied from  $0.09 \pm 0.03$  mm (mean  $\pm$  SE) in fast flow, pH 7.7 conditions on bare blades to  $0.94 \pm 0.30$  mm in slow flow and pH 8.1 condition, in the presence of bryozoans on the blade (Figure 1). While pH did not affect the thickness of the DBL, the flow (3-way ANOVA, df = 1, F = 25.908, p <0.001) and the presence of bryozoans (3-way ANOVA, df = 1, F = 5.976, p = 0.018) had significant effects on its thickness. In slow flow conditions, the DBL was always thicker than in fast flow conditions. Compared to bare blades, the thickness of DBL was greater in blades colonized by bryozoans.

265

266  $O_2$  profiles and interfacial fluxes

In both light and dark conditions, O<sub>2</sub> profiles in fast flow treatments were much steeper 267 268 than in slow flow (Table 2, Gradient shape, p = 0.001 and 0.003 in light and dark conditions, respectively), with initial oxygen concentration on the surface varying by 0.2 units (Figure 2 A, 269 B). Conversely, the oxygen gradients in slow flow treatments were more gradual, with a broader 270 range of O<sub>2</sub> values from varying by 0.6 units (Figure 2 C, D). Under dark conditions, the 271 272 presence of bryozoans decreased the oxygen content within the DBL to a greater extent than on 273 bare blades regardless of the flow speed and the mainstream pH (Table 2, Gradient shape, p = 0.003). Thus, the shape of the profiles differed between fast and slow flow and in the presence 274 or absence of bryozoans in dark conditions, but no interactive effects were detected either in 275 the light or dark. 276

277 The mean standardized O<sub>2</sub> concentration measured at the surface of the kelp blades in the different conditions were differently affected by pH, flow and presence/absence of 278 279 bryozoans depending on the light conditions (Table 3). Under saturating light, O<sub>2</sub> concentrations were > 1 because photosynthetic processes released oxygen. So the oxygen 280 concentration just above the blade was higher than in the mainstream seawater. In the light, O<sub>2</sub> 281 concentrations were significantly affected by pH (Table 2, surface  $[O_2]$ , p = 0.044) with higher 282 283 concentrations in pH 7.7 condition (1.14  $\pm$  0.03) compared to pH 8.1 (1.25  $\pm$  0.06). O<sub>2</sub> 284 concentrations were also significantly lower in fast than in slow flow (1.10  $\pm$  0.03 and 1.28  $\pm$ 0.05, respectively, Table 2, surface  $[O_2]$ , p = 0.002). In the dark, standardized  $O_2$  concentrations 285 were < 1 because of the use of oxygen by respiration. In contrast to the light conditions, pH did 286 287 not affect the O<sub>2</sub> levels at the surface in the dark. However, similar to the light conditions, they varied according to the flow, being higher in fast flow treatments than in slow flow  $(0.91 \pm 0.02)$ 288 and  $0.59 \pm 0.06$ , Table 2, surface [O<sub>2</sub>], p > 0.001). The presence of bryozoans also significantly 289 decreased the oxygen concentrations at the surface compared to bare blades ( $0.68 \pm 0.06$  and 290

291  $0.85 \pm 0.05$ , respectively, Table 2, surface [O<sub>2</sub>], p = 0.008). No interactive effects were detected 292 either in the light or dark.

Interfacial fluxes in light condition were assimilated to net production (Table 3). Net production increased significantly by 80% in pH 7.7 compared to pH 8.1 ( $0.30 \pm 0.05$  and  $0.17 \pm 0.05 \mu$ mol m<sup>-2</sup> s<sup>-1</sup>, respectively, Table 2, Interfacial flux, p = 0.049). The presence of bryozoans also influenced the interfacial flux in the light (Table 2, Interfacial flux, p = 0.042). Net production was higher on bare than epiphytized blades ( $0.29 \pm 0.06$  and  $0.17 \pm 0.04 \mu$ mol m<sup>-2</sup> s<sup>-1</sup>, respectively), likely due to the oxygen consumption by bryozoans. No significant effect of pH, flow or blade state on respiration rate, i.e. interfacial fluxes in dark, were detected.

300

#### 301 *pH profiles and pH-O*<sub>2</sub> *relationship (slow flow conditions only)*

Among the 64 profiles measured, only 10 pH profiles in slow flow conditions were 302 303 useable due to technical issues with the microprobes. These profiles are presented in Figure 3, which shows the variation in pH units from the mainstream seawater (increase in light 304 conditions and decrease in dark condition) for slow flow conditions, for bare blades and blades 305 with bryozoans, in the different mainstream pH conditions tested. On bare blades, pH variations 306 307 between light and dark were similar in both pH treatments, around 0.5 unit (Figure 3A). 308 However, pH at the blade surface in the pH 7.7 condition shifted towards higher values than those measured in pH 8.1 treatment. In the light, the increase in photosynthesis at pH 7.7 309 elevated the mainstream pH by almost 0.3 units in the first µm of the DBL, compared to only 310 311 0.2 unit at pH 8.1. This means that the pH at the blade surface ranged from 7.8 to 8.3 in current pH conditions and 7.5 to 8 in OA conditions, along daily cycles. For epiphytized blades at pH 312 313 8.1, the range of pH variation between light and dark was narrower than on bare blades ( $\Delta pH$ = 0.26 pH, Figure 3B). At pH 7.7, pH fluctuations on blades with bryozoans varied depending 314 on the profiles, with narrower or larger fluctuations than for the pH 8.1 condition. 315

The relationship between pH deviation from the mainstream value and oxygen standardized values in the DBL was linear for bare blades, in both light and dark conditions ( $R^2$ = 0.95 to 0.99, p < 0.01, Figure 3C). The shape of the relationship for blades with bryozoans were less obviously linear (Figure 3D) even if some of the linear regression were significant ( $R^2 = 0.82$  to 0.96).

321

## 322 Discussion

323 The fluctuations of oxygen concentrations and pH, occurring along DBL gradients in the 324 different conditions of light, flow, mainstream pH and epiphytism tested in our study showed that interactions between abiotic and biotic factors can lead to unexpected chemical variations 325 in the microhabitats on the blades of the kelps and more generally on macrophytes (Koch 1994; 326 Short, Pedersen & Kendrick 2015; Hofmann, Koch & de Beer 2016). The presence of the 327 bryozoans on kelp blades increased the complexity of the habitat at the microscale. Particularly 328 329 in slow flow, DBL thickness increased, merging the algal DBL with that of the bryozoans'. The interaction between the mainstream pH and the physiology of both kelp and epiphytes also led 330 to variations in oxygen concentrations and pH in the DBL which could afford protection from 331 332 future ocean acidification. This kind of interaction may enhance the resilience of organisms and ecosystems to ocean acidification, even for those occurring on a microscale. 333

The thickness of the DBL was regulated by hydrodynamics, which engineer the laminar layer above the blade where the chemical concentration gradient occurs. So, as we expected, the DBLs were thin in faster flows, and thicker under slow flows. Seawater pH did not influence the thickness of the DBL, in fast and slow flows, because physical processes set the maximum thickness. The average thicknesses recorded for bare blades in our study for slow (~ 0.5 mm) and fast flow (~ 0.1 mm) were within the range reported for *Macrocystis sp.*, a genus closely related to *Ecklonia* and the only other kelp for which DBL thickness have been measured (Hurd
2000 and references therein) and thinner (-50%) than those of coralline algae (Raven & Hurd
2012).

343 The thickness of the DBL was not only a factor of flow speed but was also affected by the presence of bryozoans on E. radiata blades. Opposite to our second hypothesis, the DBL was 344 thicker on blades colonized by bryozoans (~ 0.8 mm in slow flow) than on bare blades (~ 0.3 345 346 mm in fast flow). The ciliary motion of zoides was expected to create turbulence, impacting the building of the laminar layer above the blade. However, the presence of the bryozoan layer 347 increased the thickness of the DBL in all the conditions by creating their own DBL in addition 348 349 to the kelp's one, essentially as a small canopy boundary layer (Cornwall et al. 2015). This phenomenon has also been observed on Fucus species where the presence of hyaline hairs 350 increased the DBL thickness compared to hairless thalli (Spilling et al. 2010; Lichtenberg, 351 Nørregaard & Kühl 2017). The explanation is that small scale (µm - mm) surface topographical 352 features such as corrugations, the presence of hyaline hairs or the cilia of bryozoans likely 353 354 reduce the roughness Reynold's number (Hurd et al. 1997; Hurd & Pilditch 2011; Lichtenberg, 355 Nørregaard & Kühl 2017) or create a local depression (Wangpraseurt et al. 2012) thus increasing the DBL thickness. These engineering factors, flow velocity and the presence of 356 bryozoans, are therefore not only able to affect the thickness of the DBL but also can directly 357 and/or indirectly impact the chemical gradients occuring therein. 358

Measuring the oxygen gradient occurring in the DBL is a good way to understand and characterize the surface micro-chemistry and how variable these fine-scale environments can be (Shashar, Cohen & Loya 1993). Flow was the main factor which influenced the O<sub>2</sub> gradients within the DBL at the surface of *E. radiata*. However, net photosynthesis was similar in fast and slow flow at the surface of *E. radiata*, suggesting that the photosynthetic process was not mass-transfer limited by the thick DBL in fast slow flow conditions. This finding is similar to that found for inorganic nitrogen uptake by the kelp *Macrocystis pyrifera* (Hurd et al. 1996) while other studies report mass transfer limitation of photosynthesis (Carpenter & Williams 2007; Mass *et al.* 2010). As the DBL was very thin in fast flow, the profiles were very steep and sharp in light and dark conditions and the standardized oxygen values were lower than in slow flow because of fast diffusion of the oxygen molecules out of the DBL (Irwin & Davenport 2002; Irwin & Davenport 2010). Thus, the main differences between the other parameters (mainstream pH, epiphytism) were better compared in slow flow conditions, as follows.

Net photosynthesis measured on blades increased with the pH decrease predicted for the 372 end of the century, in both bare and epiphytized blades while respiration remained similar in 373 374 both pH conditions. At pH 7.7, the standardised oxygen values measured in the light were greater in the first 0.2 mm layers of the gradient, close to the kelp surface, than at pH 8.1, due 375 to the increase in photosynthesis. Because E. radiata is a  $HCO_3$ -user for photosynthesis 376 (Falkenberg, Russell & Connell 2013; Cornwall, Revill & Hurd 2015; Britton et al. 2016), the 377 greater CO<sub>2</sub> availability related to the pH decrease may not benefit the photosynthesis process 378 379 (Koch et al. 2013, Britton et al. 2016), as shown for Macrocystis pyrifera (Fernández, Roleda & Hurd 2015). However the conversion of  $HCO_3^-$  to  $CO_2$ , the substratum required by the 380 enzyme RuBisCO, implies carbon concentrating mechanism(s) (CCM) which require more 381 energy than the passive diffusion of CO<sub>2</sub> (Giordano, Beardall & Raven 2005; Raven & Beardall 382 2014). Thus, the increase in both CO<sub>2</sub> availability and passive diffusion at pH 7.7 may provide 383 an advantage to *E. radiata*, reducing energetic costs of CCMs and helping to increase primary 384 production (Cornwall et al. 2012; Raven, Beardall & Giordano 2014). 385

The presence of bryozoans on the kelp blades was correlated to a decrease in the standardized  $O_2$  concentrations in the DBL compared to bare blades. Net photosynthesis was lower on epiphytised blades as the bryozoans likely used oxygen produced by the algae for respiration. Bryozoans might also have caused shading on the algal blade, decreasing its

photosynthetic efficiency (Muñoz, Cancino & Molina 1991). However, no change in 390 391 chlorophyll a content was detected between bare and epiphytised blades (see Supporting Informations), supporting the findings of prior studies that reveal a lack of bryozoan effect on 392 pigment acclimation of kelp blades, and thus on photosynthetic efficiency (Hepburn, Hurd & 393 Frew 2006). The bryozoans' contribution to the oxygen concentration gradient was particularly 394 obvious in the dark at pH 8.1, where standardized oxygen values of blades colonized by 395 396 bryozoans were two-times lower than on bare blades. This trend was less pronounced at pH 7.7 in the dark whereby the DBL oxygen concentrations looked similar for epiphytized and bare 397 blades. The pH decrease in the mainstream seawater did not affect the respiration rates of the 398 399 epiphyte/blade complex at the blade surface likely because bryozoans present a great plasticity 400 and different strategies which enable them to cope with pH decrease (Swezey et al. 2017a; Swezey et al. 2017b). Moreover, the pH flucutations occuring in the micro-environment of 401 402 seaweeds can reach very low levels, from 8.1 down to 7.0 in the dark (De Beer & Larkum 2001; 403 Hurd et al. 2011) and bryozoans living upon the blades may be used to these daily drops in the surrounding pH. 404

405 It is more and more recognized that the biological activity of both the seaweeds and their epibionts can generate pH fluctuations with ranges, rates and magnitude of change that mask 406 407 the long-term trend predicted for the open ocean (Krause-Jensen et al. 2015; Hendriks et al. 2017; Wahl et al. 2017). pH measured at the surface of bare blades of E. radiata was linearly 408 correlated to the oxygen values recorded in the DBL in light as in dark conditions, as it has been 409 shown for other seaweeds (De Beer & Larkum 2001; Larkum, Koch & Kühl 2003; Cornwall et 410 al. 2013b). In slow flow, the mean daily pH in the first layers of the DBL remained higher in a 411 412 mainstream pH 8.1 compared to 7.7, similar to the findings for coralline algae (Cornwall et al. 2013b). However, pH fluctuations occurring in the DBL at low mainstream pH could reach pH 413 values > 7.7 in the light, likely providing respite conditions from the corrosive mainstream 414

415 seawater. This rise in pH at the blade surface compared to the bulk seawater pH has also been 416 measured in other macrophytes, with an increase from 0.4 to 1.2 pH unit (Hendricks et al. 417 2017). In our study, the presence of bryozoans narrowed the pH range experienced at the blade surface in pH 8.1 condition compared to bare blades, restricting the exposure of bryozoans to 418 low pH values. In low pH conditions, however, the intensity of DBL pH fluctuation differed 419 depending on the profiles, maybe because of individual-specific responses of bryozoans to pH 420 421 due to phenotypic variance (Eriander, Wrange & Havenhand 2015). Overall, the shift in the range of pH in the DBL towards higher values than the bulk seawater and narrower fluctuations 422 in the presence of bryozoans could lead to local chemical conditions which may be more 423 424 favorable than those of the mainstream seawater for organisms living on top or under the blades. 425 Thus, DBL microhabitats may constitute a refuge from the general pH decline predicted in the context of ocean acidification. Such variations have also been evidenced at the blade surface of 426 427 green algae and seagrasses (Hendricks et al. 2017), extending the concept of OA refuge for small and understory organisms to other macrophytes. 428

429 The substantial fluctuations in O<sub>2</sub> concentration and pH reported here in the DBL and in other studies on macrophytes (e.g. Hofmann et al. 2016; Hendrick et al. 2017) may represent 430 an alternation between stressful periods and release from pH stress for organisms associated 431 432 with seaweeds (Shashar, Cohen & Loya 1993; Wahl, Saderne & Sawall 2016). Living in microhabitats with high fluctuations of environmental variables would be particularly 433 interesting for calcifying organisms in the context of ocean acidification. In fluctuating 434 environments, calcifiers may be less exposed, at least for a shorter period (Wahl et al. 2015), to 435 corrosive low pH/carbonate chemistry parameters which might decrease the dissolution process 436 437 and help them to cope with ocean acidification (Cornwall et al. 2014; Wahl et al. 2017). Calcifiers such as mussels are able to take advantage of these fluctuations by shifting the 438 majority of their costly physiological processes, including calcification, to times when the 439

surrounding chemical conditions are more favorable (i.e. daytime) (Wahl et al. 2017). The DBL 440 441 chemical microhabitat could also constitute refuge and benefit other small invertebrates for whom ocean acidification can affect the regulation of cellular homeostasis (Melzner et al. 442 2009). They might then represent interesting and attractive settlement sites for early life stage 443 of invertebrates which are supposed to lack efficient acid-base regulation systems (Byrne 2011). 444 Organisms which regularly encounter such strong fluctuations might be better able to 445 446 survive in the lower average pH seawater of the future, because they may present higher phenotypic plasticity (Hurd et al. 2011; Boyd et al. 2016). Transient transgressions of tolerance 447 thresholds followed by relaxation periods in fluctuating regimes may not only favor high 448 449 phenotypic plasticity but may also select for more robust genotypes (Melzner et al. 2009; Frieder et al. 2014). Thus the microenvironment of the DBL, particularly in slow flows, may 450 act as a stop-gap, enabling calcifiers and other organisms to acclimatize or adapt to lower pH 451 452 conditions projected for the coming decades (Hurd 2015).

The concept of refuge highlighted in this study can be generalized for the DBL 453 environments at the surface of kelp blades towards larger scales under the influence of biogenic 454 fluctuations created by macrophytes. Different studies have shown that oxygen and pH 455 456 fluctuations recorded in seagrass meadows (Unsworth et al. 2012; Hendriks et al. 2014; Saderne 457 et al. 2015) and seaweeds beds (Buapet, Gullström & Björk 2013; Pajusalu et al. 2013; Wahl et al. 2017) were masking the global pH decrease predicted in the context of OA. These specific 458 chemical habitats dominated by macrophytes could help in mitigating the negative impact of 459 460 ocean acidification on calcifiers which live on the blades, in the canopy and in understory (Semesi, Beer & Bjork 2009; Cox et al. 2017). They would not only provide temporal and 461 spatial refuge, enabling species to better cope with stressful conditions, but they could also 462 constitute selective areas facilitating the hardening of some populations and then, helping in 463 conserving resistant locally adapted populations. 464

# 466 Acknowledgments

467	The authors thank Clara Peron for her assistance during sampling and data analyses,
468	Pamela Fernandez for nutrient analyses and Michael Ellwood for alkalinity determination.
469	Fanny Noisette was supported by an Endeavour Research Fellowship from the Australian
470	Government. This work included the use of invertebrates (bryozoans) and seaweeds. Their
471	collection was allowed by a permit provided by the Department of Primary Industries, Parks,
472	Water and Environment of Tasmania (n°15115). All applicable institutional and/or national
473	guidelines for the care and use of living organisms were followed.
474	
475	Author contribution:
476	FN and CH conceived and designed the experiments. FN performed the experiments and
477	analyzed the data. FN and CH wrote the manuscript.
478	
479	Data accessibility
480	All data from this article will be available on Pangaea database, under OA-ICC data
481	compilation, with a digital object identifier DOI assigned later.
482	
483	References
484	Boyd, P.W., Cornwall, C.E., Davison, A., Doney, S.C., Fourquez, M., Hurd, C.L., Lima, I.D.
485	& McMinn, A. (2016) Biological responses to environmental heterogeneity under
486	future ocean conditions. Global Change Biology, 22 (8), 2633–2650

487	Britton, D., Cornwall, C.E., Revill, A.T., Hurd, C.L. & Johnson, C.R. (2016) Ocean
488	acidification reverses the positive effects of seawater pH fluctuations on growth and
489	photosynthesis of the habitat-forming kelp, Ecklonia radiata. Scientific Reports, 6,
490	26036.

Brodersen, K.E., Lichtenberg, M., Paz, L.-C. & Kühl, M. (2015) Epiphyte-cover on seagrass
(Zostera marina L.) leaves impedes plant performance and radial O2 loss from the
below-ground tissue. *Frontiers in Marine Science*, 2, 58.

494 Buapet, P., Gullström, M. & Björk, M. (2013) Photosynthetic activity of seagrasses and

495 macroalgae in temperate shallow waters can alter seawater pH and total inorganic

496 carbon content at the scale of a coastal embayment. *Marine and Freshwater Research*,

**64,** 1040-1048.

Byrne, M. (2011) Impact of ocean warming and ocean acidification on marine invertebrate
life history stages: vulnerabilities and potential for persistence in a changing ocean.

500 Oceanography and Marine Biology: An Annual Review (eds R. Gibson, R. Atkinson,

501 J. Gordon, I. Smith & D. Hughes), pp. 1-42. Taylor & Francis.

- Carpenter, R.C. & Williams, S.L. (2007) Mass transfer limitation of photosynthesis of coral
   reef algal turfs. *Marine Biology*, 151, 435-450.
- Cornwall, C.E., Boyd, P.W., McGraw, C.M., Hepburn, C.D., Pilditch, C.A., Morris, J.N.,
  Smith, A.M. & Hurd, C.L. (2014) Diffusion boundary layers ameliorate the negative
  effects of ocean acidification on the temperate coralline macroalga *Arthrocardia corymbosa. Plos One*, **9**, e97235.
- 508 Cornwall, C.E., Hepburn, C.D., McGraw, C.M., Currie, K.I., Pilditch, C.A., Hunter, K.A.,
- 509 Boyd, P.W. & Hurd, C.L. (2013a) Diurnal fluctuations in seawater pH influence the
- response of a calcifying macroalga to ocean acidification. *Proceedings of the Royal Society B: Biological Sciences*, 280, 2013-2201.

Cornwall, C.E., Hepburn, C.D., Pilditch, C.A. & Hurd, C.L. (2013b) Concentration boundary
layers around complex assemblages of macroalgae: Implications for the effects of ocean
acidification on understory coralline algae. *Limnology & Oceanography*, 58, 121-130.

515 Cornwall, C.E., Hepburn, C.D., Pritchard, D., Currie, K.I., McGraw, C.M., Hunter, K.A. &
516 Hurd, C.L. (2012) Carbon-use strategies in macroalgae: differential responses to

lowered pH and implications for ocean acidification *Journal of Phycology*, **48**, 137-144.

- Cornwall, C.E., Pilditch, C.A., Hepburn, C.D. & Hurd, C.L. (2015) Canopy macroalgae
  influence understorey corallines' metabolic control of near-surface pH and oxygen
  concentration. *Marine Ecology Progress Series*, 525, 81-95.
- 521 Cornwall, C.E., Revill, A.T. & Hurd, C.L. (2015) High prevalence of diffusive uptake of CO2
  522 by macroalgae in a temperate subtidal ecosystem. *Photosynthesis research*, **124**, 181523 190.
- Cox, T.E., Nash, M., Gazeau, F., Déniel, M., Legrand, E., Alliouane, S., Mahacek, P., Le Fur,
   A., Gattuso, J.-P. & Martin, S. (2017) Effects of in situ CO2 enrichment on Posidonia
   oceanica epiphytic community composition and mineralogy. *Marine Biology*, 164,
- 527 103.

517

- De Beer, D., Kühl, M., Stambler, N. & Vaki, L. (2000) A microsensor study of light enhanced
  Ca 2+ uptake and photosynthesis in the reef-building hermatypic coral Favia sp. *Marine Ecology Progress Series*, 194, 75-85.
- De Beer, D. & Larkum, A.W.D. (2001) Photosynthesis and calcification in the calcifying algae
   *Halimeda discoidea* studied with microsensors. *Plant Cell and Environment*, 24, 1209 1217.
- De Beer, D. & Larkum, A.W.D. (2001) Photosynthesis and calcification in the calcifying
  algae *Halimeda discoidea* studied with microsensors. *Plant Cell and Environment*, 24,
  1209-1217.

- 537 Delille, B., Borges, A. & Delille, D. (2009) Influence of giant kelp beds (Macrocystis pyrifera)
  538 on diel cycles of pCO 2 and DIC in the Sub-Antarctic coastal area. *Estuarine, Coastal*539 *and Shelf Science*, **81**, 114-122.
- Delille, B., Delille, D., Fiala, M., Prevost, C. & Frankignoulle, M. (2000) Seasonal changes of
  pCO2 over a subantarctic Macrocystis kelp bed. *Polar Biology*, 23, 706-716.
- 542 Denny, M. (2015) *Ecological mechanics: Principles of life's physical interactions*. Princeton
  543 University Press.
- 544 Dickson, A.G. & Millero, F.J. (1987) A comparison of the equilibrium constants for the
  545 dissociation of carbonic acid in seawater media. *Deep Sea Research*, 34, 1733-1743.
- 546 Dickson, A.G., Sabine, C.L. & Christian, J.R. (2007) Guide to best practices for ocean CO<sub>2</sub>
  547 measurements. *PICES special publication*, pp. 176. North Pacific Marine Science
  548 Organization, Sidney, British Columbia.
- 549 Dworjanyn, S., De Nys, R. & Steinberg, P. (2006) Chemically mediated antifouling in the red
  550 alga Delisea pulchra. *Marine Ecology Progress Series*, **318**, 153-163.
- 551 Eriander, L., Wrange, A.-L. & Havenhand, J.N. (2015) Simulated diurnal pH fluctuations
- radically increase variance in—but not the mean of—growth in the barnacle Balanus
  improvisus. *ICES Journal of Marine Science: Journal du Conseil*.
- Falkenberg L, Russell B, Connell S (2013) Contrasting resource limitations of marine primary
  producers: implications for competitive interactions under enriched CO2 and nutrient
  regimes. Oecologia, 172, 575-583.
- Fernández, P.A., Roleda, M.Y. & Hurd, C.L. (2015) Effects of ocean acidification on the
   photosynthetic performance, carbonic anhydrase activity and growth of the giant kelp
   *Macrocystis pyrifera. Photosynthesis research*, 124 (3) 293-304.

- Frieder, C.A., Gonzalez, J.P., Bockmon, E.E., Navarro, M.O. & Levin, L.A. (2014) Can
  variable pH and low oxygen moderate ocean acidification outcomes for mussel larvae? *Global Change Biology*, 20 (3), 754-764
- Frieder, C., Nam, S., Martz, T. & Levin, L. (2012) High temporal and spatial variability of
  dissolved oxygen and pH in a nearshore California kelp forest. *Biogeosciences*, 9, 39173930.
- Gattuso, J.-P., Epitalon, J.-M. & Lavigne, H. (2016) seacarb: Seawater Carbonate Chemistry.
  R package version 3.0.13. http://CRAN.R-project.org/package=seacarb.
- 568 Gattuso, J.-P. & Hansson, L. (2011) Ocean acidification. pp. 326. Oxford University Press.
- Gaylord, B., Rosman, J.H., Reed, D.C., Koseff, J.R., Fram, J., MacIntyre, S., Arkema, K.,
  McDonald, C., Brzezinski, M.A. & Largier, J.L. (2007) Spatial patterns of flow and
  their modification within and around a giant kelp forest. *Limnology and Oceanography*,
  572 52, 1838-1852.
- Giordano, M., Beardall, J. & Raven, J.A. (2005) CO<sub>2</sub> concentrating mechanisms in algae:
  mechanisms, environmental modulation, and evolution. *Annual Review of Plant Biology*, pp. 99-131. Annual Reviews, Palo Alto.
- Hansen, A.T., Hondzo, M. & Hurd, C.L. (2011) Photosynthetic oxygen flux by Macrocystis
  pyrifera: a mass transfer model with experimental validation. *Marine Ecology Progress Series*, 434, 45-55.
- Hendriks, I.E., Duarte, C.M., Olsen, Y.S., Steckbauer, A., Ramajo, L., Moore, T.S., Trotter,
  J.A. & McCulloch, M. (2015) Biological mechanisms supporting adaptation to ocean
  acidification in coastal ecosystems. *Estuarine, Coastal and Shelf Science*, 152, A1-A8.
- Hendriks, I.E., Duarte, C.M., Marbà, N. & Krause-Jensen, D. (2017) pH gradients in the
  diffusive boundary layer of subarctic macrophytes. *Polar Biology*, 1-6.

- Hendriks, I.E., Olsen, Y.S., Ramajo, L., Basso, L., Steckbauer, A., Moore, T.S., Howard, J. &
  Duarte, C.M. (2014) Photosynthetic activity buffers ocean acidification in seagrass
  meadows. *Biogeosciences*, 11, 333-346.
- Hepburn, C.D., Frew, R.D. & Hurd, C.L. (2012) Uptake and transport of nitrogen derived from
  sessile epifauna in the giant kelp Macrocystis pyrifera. *Aquatic Biology*, 14, 121-128.
- Hepburn, C.D. & Hurd, C.L. (2005) Conditional mutualism between the giant kelp Macrocystis
  pyrifera and colonial epifauna. *Marine Ecology Progress Series*, **302**, 37-48.
- Hepburn, C.D., Hurd, C.L. & Frew, R.D. (2006) Colony structure and seasonal differences in
  light and nitrogen modify the impact of sessile epifauna on the giant kelp Macrocystis
  pyrifera (L.) C Agardh. *Hydrobiologia*, 560, 373-384.
- Hofmann, L.C., Koch, M. & de Beer, D. (2016) Biotic Control of Surface pH and Evidence of
  Light-Induced H+ Pumping and Ca2+-H+ Exchange in a Tropical Crustose Coralline
  Alga. *Plos One*, **11**, e0159057.
- Hurd, C.L. (2000) Water motion, marine macroalgal physiology, and production. *Journal of Phycology*, **36**, 453-472.
- Hurd, C.L. (2015) Slow flow habitats as refugia for coastal calcifiers from ocean acidification. *Journal of Phycology*, **51**, 599-605.
- 601 Hurd, C.L., Cornwall, C.E., Currie, K., Hepburn, C.D., McGraw, C.M., Hunter, K.A. & Boyd,
- 602 P.W. (2011) Metabolically induced pH fluctuations by some coastal calcifiers exceed
- projected 22<sup>nd</sup> century ocean acidification: a mechanism for differential susceptibility?
- 604 *Global Change Biology*, **17**, 3254-3262.
- Hurd, C.L., Durante, K.M., Chia, F.S. & Harrison, P.J. (1994a) Effect of bryozoan colonization
  on inorganic nitrogen acquisition by the kelps Agarum fimbriatum and Macrocystis
  integrifolia. *Marine Biology*, **121**, 167-173.

- Hurd, C.L., Harrison, P.J., Bischof, K. & Lobban, C.S. (2014) *Seaweed ecology and physiology*.
  Cambridge University Press.
- Hurd, C.L. & Pilditch, C.A. (2011) Flow induced morphological variations affect diffusion
  boundary layer thickness of *Macrocystis pyrifera* (Heterokontophyta, Laminariales) *Journal of Phycology*, 47, 341-351.
- Hurd, C.L., Quick, M., Stevens, C.L., Laval, B.E., Harrison, P.J. & Druehl, L.D. (1994b) A
  low-volume flow tank for measuring nutrient uptake by large macrophytes. *Journal of Phycology*, **30**, 892-896.
- Hurd, C.L., Stevens, C.L., Laval, B.E., Lawrence, G. & Harrison, P. (1997) Visualization of
- seawater flow around morphologically distinct forms of the giant kelp Macrocystis
  integrifolia from wave-sheltered and exposed sites. *Limnology and Oceanography*, 42,
  156-163.
- Irwin, S. & Davenport, J. (2002) Hyperoxic boundary layers inhabited by the epiphytic
  meiofauna of Fucus serratus. *Marine Ecology Progress Series*, 244, 73-79.
- Irwin, S. & Davenport, J. (2010) Oxygen microenvironment of coralline algal tufts and their
  associated epiphytic animals. *Biology and Environment: Proceedings of the Royal Irish Academy*, pp. 185-193. JSTOR.
- Jackson, G.A. & Winant, C.D. (1983) Effect of a kelp forest on coastal currents. *Continental Shelf Research*, 2, 75-80.
- Koch, E.W. (1994) Hydrodynamics, diffusion-boundary layers and photosynthesis of the
  seagrasses Thalassia testudinum and Cymodocea nodosa. *Marine Biology*, **118**, 767776.
- Koch, M., Bowes, G., Ross, C. & Zhang, X.H. (2013) Climate change and ocean acidification
  effects on seagrasses and marine macroalgae. *Global Change Biology*, **19**, 103-132.

- Krause-Jensen, D., Duarte, C.M., Hendriks, I.E., Meire, L., Blicher, M.E., Marbà, N. & Sejr,
  M.K. (2015) Macroalgae contribute to nested mosaics of pH variability in a subarctic
  fjord. *Biogeosciences*, 12, 4895-4911.
- Kregting, L.T., Stevens, C.L., Cornelisen, C.D., Pilditch, C.A. & Hurd, C.L. (2011) Effects of
  a small-bladed macroalgal canopy on benthic boundary layer dynamics: implications
  for nutrient transport. *Aquatic Biology*, 14, 41-56.
- Kroeker, K.J., Kordas, R.L., Crim, R., Hendriks, I.E., Ramajo, L., Singh, G.S., Duarte, C.M. &
  Gattuso, J.-P. (2013) Impacts of ocean acidification on marine organisms: quantifying
  sensitivities and interaction with warming. *Global Change Biology*, **19**, 1884-1896.
- Kühl, M., Cohen, Y., Dalsgaard, T., Jørgensen, B. & Revsbech, N.P. (1995) Microenvironment
  and photosynthesis of zooxanthellae in scleractinian corals studied with microsensors
  for 02, pH and light. *Mar. Ecol. Prog. Ser*, **117**, 159-172.
- Larkum, A.W.D., Koch, E.M.W. & Kühl, M. (2003) Diffusive boundary layers and
  photosynthesis of the epilithic algal community of coral reefs. *Marine Biology*, 142,
  1073-1082.
- Lichtenberg, M., Nørregaard, R.D. & Kühl, M. (2017) Diffusion or advection? Mass transfer
  and complex boundary layer landscapes of the brown alga Fucus vesiculosus. *Journal of The Royal Society Interface*, 14, 20161015.
- Mass, T., Genin, A., Shavit, U., Grinstein, M. & Tchernov, D. (2010) Flow enhances
  photosynthesis in marine benthic autotrophs by increasing the efflux of oxygen from
  the organism to the water. *Proceedings of the National Academy of Sciences*, 107, 25272531.
- Mehrbach, C., Culberso.Ch, Hawley, J.E. & Pytkowic, R.M. (1973) Measurement of apparent
  dissociation-constants of carbonic-acid in seawater at atmospheric-pressure. *Limnology*& Oceanography, 18, 897-907.

657	Melzner, F., Gutowska, M.A., Langenbuch, M., Dupont, S., Lucassen, M., Thorndyke, M.C.,
658	Bleich, M. & Pörtner, HO. (2009) Physiological basis for high CO <sub>2</sub> tolerance in marine
659	ectothermic animals: pre-adaptation through lifestyle and ontogeny? <i>Biogeosciences</i> , 6,
660	2313-2331.

- Muñoz, J., Cancino, J. & Molina, M. (1991) Effect of encrusting bryozoans on the physiology
  of their algal substratum. *Journal of the Marine Biological Association of the United Kingdom*, **71**, 877-882.
- Pajusalu, L., Martin, G., Põllumäe, A. & Paalme, T. (2013) Results of laboratory and field
  experiments of the direct effect of increasing CO2 on net primary production of
  macroalgal species in brackish-water ecosystems. *Proceedings of the Estonian Academy of Sciences*, **62**, 148-154.
- Pearson, G.A., Serrão, E.A. & Brawley, S.H. (1998) Control of gamete release in fucoid algae:
  sensing hydrodynamic conditions via carbon acquisition. *Ecology*, **79**, 1725-1739.
- 670 R Core Team (2013) R: a language and environment for statistical computing. R Foundation
  671 for Statistical Computing, Vienna, Austria.
- Raven, J. (1997) Inorganic carbon acquisition by marine autotrophs. *Adv. Bot. Res*, 27, 85-209.
- Raven, J.A. & Beardall, J. (2014) CO2 concentrating mechanisms and environmental change. *Aquatic Botany*, **118**, 24-37.
- Raven, J.A., Beardall, J. & Giordano, M. (2014) Energy costs of carbon dioxide concentrating
  mechanisms in aquatic organisms. *Photosynthesis research*, **121**, 111-124.
- Raven, J.A. & Hurd, C.L. (2012) Ecophysiology of photosynthesis in macroalgae. *Photosynthesis research*, **113**, 105-125.
- Reed, D.C. & Foster, M.S. (1984) The effects of canopy shadings on algal recruitment and
  growth in a giant kelp forest. *Ecology*, 65, 937-948.

- Revsbech, N.P. & Jørgensen, B.B. (1986) Microelectrodes: their use in microbial ecology.
   *Advances in microbial ecology*, pp. 293-352. Springer.
- Riebesell, U., Fabry, V.J., Hansson, L. & Gattuso, J.-P. (2010) Guide to best practices for ocean
  acidification research and data reporting. pp. 258. Luxembourg: Publications Office of
  the European Union
- Rosman, J.H., Monismith, S.G., Denny, M.W. & Koseff, J.R. (2010) Currents and turbulence
  within a kelp forest (Macrocystis pyrifera): Insights from a dynamically scaled
  laboratory model. *Limnology and Oceanography*, 55, 1145.
- 689
- Short, J., Pedersen, O. & Kendrick, G. (2015) Turf algal epiphytes metabolically induce local
  pH increase, with implications for underlying coralline algae under ocean acidification. *Estuarine, Coastal and Shelf Science*, 164, 463-470.
- Saderne, V., Fietzek, P., Aßmann, S., Körtzinger, A. & Hiebenthal, C. (2015) Seagrass beds as
  ocean acidification refuges for mussels? High resolution measurements of pCO2 and
  O2 in a Zostera marina and Mytilus edulis mosaic habitat. *Biogeosciences Discuss.*, 12,
  11423-11461.
- 697 Saderne, V., Fietzek, P. & Herman, P.M.J. (2013) Extreme variations of  $pCO_2$  and pH in a 698 macrophyte meadow of the baltic sea in summer: evidence of the effect of 699 photosynthesis and local upwelling. *Plos One*, **8**, e62689.
- Saderne, V. & Wahl, M. (2013) Differential responses of calcifying and non-calcifying
  epibionts of a brown macroalga to present-day and future upwelling *p*CO<sub>2</sub>. *Plos One*, 8,
  e70455.
- Sand-Jensen, K., Revsbech, N.P. & Jörgensen, B.B. (1985) Microprofiles of oxygen in epiphyte
   communities on submerged macrophytes. *Marine Biology*, 89, 55-62.

- Schaffelke, B. (1999) Particulate organic matter as an alternative nutrient source for tropical
  Sargassum species (Fucales, Phaeophyceae). *Journal of Phycology*, **35**, 1150-1157.
- Semesi, I.S., Beer, S. & Bjork, M. (2009) Seagrass photosynthesis controls rates of calcification
   and photosynthesis of calcareous macroalgae in a tropical seagrass meadow. *Marine Ecology-Progress Series*, 382, 41-47.
- Shashar, N., Cohen, Y. & Loya, Y. (1993) Extreme diel fluctuations of oxygen in diffusive
  boundary layers surrounding stony corals. *The Biological Bulletin*, 185, 455-461.
- Spilling, K., Titelman, J., Greve, T.M. & Kuhl, M. (2010) Microsensor Measurements of the
   External and Internal Microenvironment of Fucus Vesiculosus (Phaeophyceae). *Journal*
- 714 *of Phycology*, **46**, 1350-1355.
- Steneck, R.S., Graham, M.H., Bourque, B.J., Corbett, D., Erlandson, J.M., Estes, J.A. & Tegner,
   M.J. (2002) Kelp forest ecosystems: biodiversity, stability, resilience and future.
   *Environmental conservation*, 29, 436-459.
- 718 Stocker, T.F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S.K., Boschung, J., Nauels, A., Xia,
- 719 Y., Bex, V. & Midgley, P.M. (2013) Climate Change 2013. The Physical Science Basis.
- 720 Working Group I Contribution to the Fifth Assessment Report of the Intergovernmental
- Panel on Climate Change. (ed. C.U. Press). Groupe d'experts intergouvernemental sur
  l'evolution du climat/Intergovernmental Panel on Climate Change-IPCC, C/O World
  Meteorological Organization, 7bis Avenue de la Paix, CP 2300 CH-1211 Geneva 2
  (Switzerland), Cambridge, United Kingdom and New York, NY, USA.
- Swezey, D.S., Bean, J.R., Hill, T.M., Gaylord, B., Ninokawa, A.T. & Sanford, E. (2017a)
  Plastic responses of bryozoans to ocean acidification. *The Journal of experimental biology*, 220, 4399-4409.
- Swezey, D.S., Bean, J.R., Ninokawa, A.T., Hill, T.M., Gaylord, B. & Sanford, E. (2017b)
  Interactive effects of temperature, food and skeletal mineralogy mediate biological

- responses to ocean acidification in a widely distributed bryozoan. *Proceedings of the Royal Society B: Biological Sciences*, 284.
- Unsworth, R.K., Collier, C.J., Henderson, G.M. & McKenzie, L.J. (2012) Tropical seagrass
   meadows modify seawater carbon chemistry: implications for coral reefs impacted by
   ocean acidification. *Environmental Research Letters*, 7, 024026.
- Vogel, S. (1999) *Life in Moving Fluids: The Physical Biology of Flow*, 2nd edn. Princeton
  University Press, Princeton, NJ, USA.
- Wahl, M., Buchholz, B., Winde, V., Golomb, D., Guy-Haim, T., Müller, J., Rilov, G., Scotti,
  M. & Böttcher, M.E. (2015) A mesocosm concept for the simulation of near-natural
  shallow underwater climates: The Kiel Outdoor Benthocosms (KOB). *Limnology and Oceanography: Methods*, 13, 651-663.
- Wahl, M., Saderne, V. & Sawall, Y. (2016) How good are we at assessing the impact of ocean
  acidification in coastal systems? Limitations, omissions and strengths of commonly
  used experimental approaches with special emphasis on the neglected role of
  fluctuations. *Marine and Freshwater Research*, 67, 25-36.
- Wahl, M., Schneider Covachã, S., Saderne, V., Hiebenthal, C., Müller, J.D., Pansch, C. &
  Sawall, Y. (2017) Macroalgae may mitigate ocean acidification effects on mussel
  calcification by increasing pH and its fluctuations. *Limnology and Oceanography*, in
  press, 10.1002/lno.10608
- 749 Wangpraseurt, D., Weber, M., Røy, H., Polerecky, L., de Beer, D., Suharsono & Nugues, M.M.
- 750 (2012) *In situ* oxygen dynamics in coral-algal interactions. *Plos One*, **7**, e31192
- Wheeler, W. (1980) Effect of boundary layer transport on the fixation of carbon by the giant
  kelp Macrocystis pyrifera. *Marine Biology*, 56, 103-110.
- 753

## 754 **Supporting information**

Additional supporting information about chlorophyll contents may be found in the onlineversion of this article.

757

758

**Table 1:** Seawater parameters in current and low pH conditions. Temperature, salinity, pH<sub>T</sub>, total alkalinity (A<sub>T</sub>) and nutrients were measured directly in the surrounding seawater while all the other parameters of the carbonate chemistry ( $pCO_2$ , [CO<sub>2</sub>], [HCO<sub>3</sub><sup>-</sup>], [CO<sub>3</sub><sup>2-</sup>], DIC,  $\Omega_{Ar}$  and  $\Omega_{ca}$ ) were calculated.

	pH 8.1 condition			pH 7.7 condition		
	n	Mean	SE	n	Mean	SE
Temperature (°C)	13	13.19 ±	0.12	12	13.17 ±	0.12
Salinity	13	$36.85$ $\pm$	0.10	12	$37.25$ $\pm$	0.25
$pH_{T}$	13	$8.09 \ \pm$	0.02	12	7.71 ±	0.01
$A_T (\mu Eq kg^{-1})$	5	$2099.82 \ \pm$	26.48	5	$2242.72 \ \pm$	16.93
$pCO_2$ (µatm)	13	$330.87 \ \pm$	16.81	12	932.33 $\pm$	22.89
$[CO_2] \ (\mu mol \ kg^{-1})$	13	$12.92 \ \pm$	0.65	12	36.34 ±	0.86
[HCO3 <sup>-</sup> ] (µmol kg <sup>-1</sup> )	13	$1712.79 \ \pm$	12.69	12	$2049.86~\pm$	4.17
$[CO_3^{2-}] (\mu mol kg^{-1})$	13	$151.58~\pm$	4.97	12	$76.56 \pm$	1.62
DIC (µmol kg <sup>-1</sup> )	13	$1877.29\ \pm$	8.36	12	$2162.76~\pm$	3.41
$\Omega_{\mathrm{Ar}}$	13	$2.29$ $\pm$	0.08	12	$1.15$ $\pm$	0.02
$\Omega_{ m ca}$	13	$3.57$ $\pm$	0.12	12	$1.80$ $\pm$	0.04
[NH4] (µM)	12	$3.03$ $\pm$	0.37	12	3.51 ±	0.11
[NO <sub>3</sub> ], [NO <sub>2</sub> ] (µM)	13	$3.64 \pm$	0.14	12	$3.15 \pm$	0.22
[PO <sub>4</sub> ] (µM)	13	$1.26$ $\pm$	0.01	12	1.26 ±	0.02

763

764

**Table 2**: Statistical results from 3-way ANOVA and MANOVA investigating the simple and interactive effects of the mainstream pH (8.1 or 7.7), the flow speed (fast or slow), and the state of the blade (bare or with bryozoans) on interfacial O<sub>2</sub> fluxes, O<sub>2</sub> standardized concentrations at the blade surface and DBL O<sub>2</sub> gradient parameters (i.e. y0,  $\alpha$  and  $\beta$  from the fitted exponential growth shape curve). Number in bold indicate significant effects.

771

772

		Interfacial flux Surface [O <sub>2</sub> ]		ce [O <sub>2</sub> ]	Gradient shape (y0, $\alpha$ and $\beta$ from fitted equation)		
		3-way	ANOVA	3-way ANOVA		3-way MANOVA	
	Df	F-value	p-value	F-value	p-value	F-value	p-value
pН	1	4.435	0.049	4.560	0.044	1.403	0.273
Flow	1	0.018	0.895	12.730	0.002	8.117	0.001
Blade	1	4.747	0.042	0.597	0.448	0.414	0.745
pH:Flow	1	0.251	0.622	1.681	0.208	0.837	0.490
pH:Blade	1	0.326	0.575	0.004	0.952	0.656	0.589
Flow:Blade	1	2.782	0.112	3.320	0.082	0.731	0.546
pH:Flow:Blade	1	0.430	0.520	1.256	0.275	0.950	0.436
	Df	F-value	p-value	F-value	p-value	F-value	p-value
pН	1	1.310	0.268	0.453	0.508	1.150	0.356
Flow	1	2.102	0.165	30.088	>0.001	6.668	0.003
Blade	1	1.668	0.214	8.598	0.008	6.684	0.003
pH:Flow	1	0.773	0.392	0.002	0.962	0.504	0.684
pH:Blade	1	2.009	0.174	3.356	0.081	2.174	0.126
Flow:Blade	1	0.054	0.819	1.319	0.264	2.449	0.097
pH:Flow:Blade	1	0.054	0.819	1.624	0.217	0.857	0.481

773

774

775

**Table 3:** Mean ( $\pm$  SE) standardized O<sub>2</sub> concentrations and interfacial fluxes (J in µmol O<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>) at the surface of the blade in the different experimental conditions of pH, flow, presence/absence of bryozoans and in saturated light and dark conditions. N = 3 or 4.

pН	Flow	Blade state	Light condition		Dark con	ndition
8.1	Fast	Bare	1.07	0.04	0.96	0.02
8.1	Fast	Bryozoans	1.08	0.06	0.81	0.04
8.1	Slow	Bare	1.24	0.06	0.81	0.11
8.1	Slow	Bryozoans	1.16	0.09	0.39	0.09
7.7	Fast	Bare	1.05	0.05	0.97	0.02
7.7	Fast	Bryozoans	1.17	0.07	0.90	0.05
7.7	Slow	Bare	1.49	0.14	0.65	0.09
7.7	Slow	Bryozoans	1.29	0.07	0.60	0.14

Mean standardized O<sub>2</sub> concentration at the blade surface

# Interfacial flux J (µmol m<sup>-2</sup> s<sup>-1</sup>)

рН	Flow	Blade state	Light condition		Dark condition	
8.1	Fast	Bare	0.16	0.12	-0.18	0.05
8.1	Fast	Bryozoans	0.13	0.10	-0.33	0.05
8.1	Slow	Bare	0.27	0.07	-0.23	0.14
8.1	Slow	Bryozoans	0.09	0.08	-0.37	0.07
7.7	Fast	Bare	0.31	0.09	-0.15	0.11
7.7	Fast	Bryozoans	0.31	0.02	-0.17	0.07
7.7	Slow	Bare	0.47	0.13	-0.31	0.06
7.7	Slow	Bryozoans	0.16	0.05	-0.28	0.05

**Figure 1:** DBL thickness in mm in fast (on the left) and slow (on the right) flow conditions, in pH 8.1 (white bars) and pH 7.7 (grey bars), in the absence (plain bars) or presence of bryozoans (striped bars). Values are mean  $\pm$  SE, n = 6 or 8.

790

**Figure 2:**  $O_2$  standardized profiles in fast (2 graphs on the left) and slow (2 graphs on the right) conditions, in pH 8.1 (2 graphs on the top) and pH 7.7 (2 graphs on the bottom). White symbols are for  $O_2$  concentrations measured in saturating light while black ones are for dark measurements. Circle represent measurements on bare blade while triangles are for blades colonized by bryozoans. Values are mean  $\pm$  SE, n = 3 or 4

796

797 Figure 3: pH variations and relationship with oxygen concentration in the DBL. Individual pH profiles in slow flow conditions are shown for bare blades (graph A, 4 profiles) and blades 798 colonized by bryozoans (graph B, 6 profiles). Solid lines represent pH deviation from 799 800 mainstream pH condition for pH 8.1 and dashed ones for pH 7.7, in both light (grey lines) and 801 dark (black lines) conditions. Graphs C and D present the relationship between pH variation and oxygen standardized concentrations in the DBL for bare blades and blades colonized by 802 803 bryozoans, respectively. Data are from the same individual profiles as graphs A and B, with grey and black symbols for light and dark conditions, respectively. Circles are for pH<sub>T</sub> 8.1 as 804 both triangles show data for  $pH_T$  condition 7.7. 805

806

807

808











