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# Evolution of sea-surface conditions on the northwestern Greenland margin during the Holocene

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ABSTRACT

Reconstructions of sea-surface conditions during the Holocene were achieved on two sediment cores from the Northwest Greenland margin (AMD14-204) and Kane Basin (AMD14-Kane2B) based on dinoflagellate cyst assemblages. On the northwest Greenland margin, sea-surface conditions were cold with an extended sea-ice cover prior to 7750 cal a BP associated to the end of the deglaciation. A major change occurred around ca. 7750 cal a BP with enhanced influence of warmer water from the West Greenland Current (WGC), and optimal sea surface conditions were observed around 6000 cal a BP. After 3350 cal a BP, results reflect establishment of the modern assemblages. In the Kane Basin, sea-surface conditions were not favorable for dinocysts productivity prior 7880 cal a BP, as the basin was still highly covered by ice. Presence of warmer water is recorded between 7880 and 7200 cal a BP and highest primary productivity between 5200 and 2100 cal a BP, but sea-surface conditions remained cold with an extended sea-ice cover throughout the Holocene. Overall, the results from this study revealed the strong influence of meltwater discharges and oceanic current variability on the sea-surface conditions.

KEYWORDS. Dinoflagellate cysts; Baffin Bay; Melville Bay; Kane Basin; Sea surface conditions

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# Introduction

Over the last decades, the Arctic has strongly responded to modern climate warming with the reduction of sea-ice extent and cover duration (Serreze and Stroeve, 2015) and to place these rapid changes in a long-term perspective, there is a need for a better understanding of the interactions between sea ice and climate throughout the geological past (Jakobsson *et al.*, 2010; Müller and Stein, 2014). Baffin Bay is an excellent location to study these interactions as sea-ice is present inside the bay during most of the year (Tang *et al.*, 2004) and because it plays the role of oceanic corridor between the Arctic and the North Atlantic oceans (through Labrador Sea, where deep convection contributes to the thermohaline circulation). Moreover, Baffin Bay has been affected by the dynamics of three large ice sheets (Greenland [GIS], Innuitian [IIS] and Laurentide [LIS] ice sheets) since the last glacial maximum (England *et al.*, 2006). Therefore, it is highly relevant to investigate how oceanographic and climatic changes affected the sea surface conditions in Baffin Bay since the last deglaciation.

Dinoflagellates cysts (hereafter, dinocysts) are excellent tracers for the estimation of past sea-surface conditions, as their highly resistant organic membrane (e.g., Fensome *et al.*, 1993) allows them to be preserved in the sediment, particularly in subpolar and polar seas where they are found in abundance and high diversity (e.g., Mudie, 1992; Matthiessen, 1995; Mudie *et al*., 2001). Several oceanographic properties can be reconstructed from dinocyst assemblages including sea-surface temperature (SST), salinity (SSS), seasonal sea-ice cover (SIC), and primary productivity (de Vernal *et al.*, 2001, 2013a; Radi and de Vernal, 2008). Previous studies based on dinocysts, but also diatoms, mollusk, foraminifera and sea ice biomarkers ($IP_{25}$) have contributed to improving our knowledge of past sea-surface conditions in Baffin Bay (e.g., Knudsen *et al.*, 2008; Krawczyk *et al.*, 2017). However, information available on the northeastern Baffin Bay (Melville Bay) and

Nares Strait sea-surface variability during the Holocene is still incomplete, as studies have mainly been focused on the regions of the North Water Polynya (e.g., Levac *et al.*, 2001) and Disko Bay region (e.g., Ribeiro *et al.*, 2012; Allan *et al.*, 2018). In this study, well-preserved dinocyst assemblages of two sediment cores recovered from the Upernavik cross-shelf trough (AMD14-204) and in the Kane Basin (AMD14-Kane2B) were analyzed and data was processed with the modern analogue technique (MAT) to reconstruct sea-surface conditions during the Holocene.

## Environmental setting

Baffin Bay is an oceanic basin 1300 km long and 450 km wide, with a depth of up to 2300 m (Aksu and Piper, 1987). It is connected to the Arctic Ocean through CAA straits and to the Atlantic Ocean through Davis Strait. Nares Strait separates Ellesmere Island from Greenland and extends North of Baffin Bay up to the Arctic Ocean (Fig. 1). Baffin Bay is influenced by the cold Baffin Island Current (BIC, temperature >-1°C, salinity <34) flowing southward along Baffin Island from northernmost Baffin Bay, and by the relatively warm and salty West Greenland Current (WGC; temperature >2°C, salinity >34) flowing northward (Fig. 1; Tang *et al.*, 2004 ; Zweng and Münchow, 2006). The WGC consists of a mix of the cold and low salinity East Greenland Current (EGC; arctic water from Fram Strait) and the warmer and salty Irminger Current (IC; North Atlantic water), both of which merge south of Greenland to form the WGC (Fig.1; Tang *et al.*, 2004). The WGC flows northward along the west Greenland shelf and turns west in Melville Bay before joining the BIC in Smith Sound. However, it has been shown that a small branch of the WGC reaches Kane Basin and circulates at depths below 250 m (Sadler, 1976). Baffin Bay is covered by sea-ice during most of the year, with a minimal extent in September (Tang *et al.*, 2004). Nevertheless, the SIC distribution in Baffin Bay is asynchronous with sea-ice extending further southward on the east than the west side due to the differences in temperature and salinity of the

two major surface currents (WGC and BIC) flowing in opposite directions (Tang *et al.*, 2004). Ice floes and icebergs are present continually in Nares Strait, even during summer (Münchow *et al.*, 2006). Landfast ice during winter leads to the formation of an ice arch in Kane Basin, which marks the boundary with the North Water Polynya (Barber *et al.*, 2001; Kwok *et al.*, 2010).

## Material and methods

Two Calypso Square (CASQ) cores were collected on board the CCGS icebreaker Amundsen in northern Baffin Bay during the 2014 ArcticNet expedition: core AMD14-204 (hereinafter referred to as 204) located on the northwest Greenland margin, and core AMD14-Kane2B (hereinafter referred to as Kane2B) located in the Kane Basin (Fig. 1; Table 1). The chronology of core 204 was established using an approach combining radiocarbon dating ($\Delta R$ = 140±30 years) with paleomagnetic chronostratigraphic markers (Caron *et al.*, 2019) and organic matter dating (Giraudeau *et al.*, under review). The chronology of core Kane2B was established by Georgiadis *et al.* (2018) based on twenty radiocarbon ages from mixed benthic foraminifera and mollusk shell fragments ($\Delta R$ = 240±51 years). Based on the established chronologies, core 204 covers the last ca. 9100 cal a BP, and has a mean sedimentation rate of 95 cm $ka^{-1}$, while core Kane2B covers the last ca. 9000 cal a BP, and has sedimentation rates ranging from 20 to 220 cm $ka^{-1}$ (Fig. S1).

### *Palynological preparations and analyses*

Palynological analyses were performed at 8 cm intervals with a volume of ~5 $cm^3$ of wet sediment following the protocol described by Rochon *et al.* (1999). Marker grains (*Lycopodium clavatum* spore tablets, batch n°414831, Lund University) were added in each sample to estimate palynomorph concentrations following the method described by Matthews (1969) and Mertens *et al.* (2009). The concentrations (cysts $cm^{-3}$) were multiplied by the sedimentation rate (cm $yr^{-1}$) in each sample in order to describe the abundances as influxes (cysts $cm^{-2}$ $yr^{-1}$). This makes it easier

to assess changes in dinoflagellate and benthic (organic linings of foraminifera) productivities in environments where sedimentation rates vary through time. After wet sieving the samples at 100 and 10 µm to eliminate sand, fine silt and clay, the 10-100 µm fraction was chemically treated with hydrochloric (HCl, 10%) to eliminate carbonates and hydrofluoric acids (HF, 49%) to eliminate silicates. The residual organic matter was mounted on microscope slides with glycerin gelatin, and slides were analyzed using a transmitted light microscope at magnifications of 400x to 1000x. Dinocysts and organic linings of foraminifera were identified and counted (Fig. 2) until reaching a minimum of 300 dinocysts and 100 marker-grains in each sample. Reworked and freshwater palynomorphs were also counted. Organic linings of foraminifera observed in the samples are mainly benthic, as they are more resistant to dissolution than those of planktic foraminifera (de Vernal *et al.*, 1992). The identification of dinocyst followed the nomenclature provided by Rochon *et al.* (1999), Head *et al.* (2001), Radi *et al.* (2013), Mertens *et al.* (2018) and Londeix *et al.* (2018).

### *Statistical approach and data analyses*

The modern analogue technique (MAT) allowed to reconstruct sea surface conditions using a dinocyst database of the Northern Hemisphere that includes 1776 reference sites and 73 taxa (Allan *et al.*, 2018). The calculations were performed using the Bioindic scripts developed for the R platform by Joel Guiot (CEREGE, France) and the procedure described by de Vernal *et al.* (2013b). The MAT relies on the hypothesis that similar biogenic assemblages occur under similar environmental conditions without any assumptions in terms of quantitative relationships (Guiot and de Vernal, 2007). Although there have been some concerns expressed that root mean square error of prediction (RMSEP) might be underestimated (Telford, 2006; Telford and Birks, 2005, 2011), this method (MAT) has been demonstrated as reliable as any reconstruction method using micropaleontological proxies (Guiot and de Vernal, 2011).

Five analogues were used, and a logarithmic transformation was applied on the relative abundance of dinocysts for the reconstructions. The errors of prediction on winter and summer SSTs are ±1.4/1.7°C, and ±1.2/2.1 for SSSs. SIC extent is expressed in number of months with more than 50% ice coverage (based on the 1953-2003 mean) (±1.4 months/year) and productivity in gC $m^{-2}$ $yr^{-1}$ (±61.5 gC $m^{-2}$ $yr^{-1}$). The Heterotrophic/Autotrophic (H/A) dinocysts ratio was calculated to assess temporal changes in the dominant trophic mode. Finally, principal component analyses (PCA) were performed with the software PAST (Hammer *et al.*, 2001) to differentiate the distribution of assemblages (taxa <1% were excluded).

## Results

The two cores used in this study (204 and Kane2B) were described in detail by Caron *et al.* (2019) and Georgiadis *et al.* (2018) and major sedimentological units (Fig. S1) were distinguished based on lithology and physical property of the sediment. Overall, the lithology of core 204 is relatively homogeneous and fine grained and is largely dominated by hemipelagic sedimentation. At the opposite, core Kane2B is showing a more variable sedimentation associated to higher inputs of ice-rafted debris (IRD) and reworked sediments throughout the core (Caron *et al.*, 2019; Georgiadis *et al.*, 2018). Palynological results provide additional details on environmental changes for both cores. A total of 25 dinocyst taxa were identified, 14 of which account for more than 95 % of the assemblages (Table 2 and Fig. 2). One interval at the base of core Kane2B (428-300 cm, ca. 9000-8300 cal a BP) was not considered for palynological reconstructions due to very low dinocyst concentrations. Relative abundance of the main taxa accounting for more than 95 % of the assemblages are presented in Figs. 3 and 5. Main dinocyst assemblage zones were distinguished based on the scores of the first two components of the PCA (Fig. S2), reflecting the occurrence of accompanying taxa and changes in abundance of the dominant species. Reconstructions of sea-

surface properties including SST, SSS, SIC and primary productivity are shown in Figs. 4 and 6, together with organic linings of foraminifera, freshwater (*Halodinium*) and reworked palynomorph influxes. The reconstructed primary productivity represents the productivity of all autotrophic groups (including diatoms, ciliates, etc.; e.g., Cormier *et al.*, 2016).

### *AMD14-204 – Northwestern Greenland margin*

Dinocyst assemblages were recovered from all samples in core 204, with an average estimated flux of 1260 ± 100 cysts $cm^{-2}$ $yr^{-1}$, and a maximum value of 4830 cysts $cm^{-2}$ $yr^{-1}$ between ca. 5700 and 5600 cal a BP (Fig. 3). Overall, reconstructed SST ranged from -1.7 to 7°C, summer SSS ranged from 27 to 32.8, SIC between 3 and 9 months per year and primary productivity between 90 and 190 gC $m^{-2}$ $yr^{-1}$ (Fig. 4). The characteristics of each dinocyst assemblage zone are described below in terms of relative abundance and sea-surface conditions reconstructions:

**Zone 1** (prior to 7750 cal a BP; 738- 604 cm) is characterized by the dominance of heterotrophic taxa, such as *Islandinium minutum* (up to 39 %) and *Brigantedinium* spp. (up to 46 %), as revealed by the H/A ratio (Fig. 4). It is also characterized by the highest relative abundances of *Impagidinium pallidum* (up to 4 %), and *Polykrikos* var. Arctic (up to 3.5 %), which nearly disappears after 7750 cal a BP (Fig. 3). Sea-surface reconstructions reveal the lowest SST values (summer SST averaging ~2°C) and the maximum SIC duration (8-9 months/year) of the entire core (Fig. 4).

**Zone 2** (7750-3350 cal a BP; 604-164 cm) is dominated by "warmer water" autotrophic species (Rochon et al., 1999; Zonneveld *et al.*, 2013), including *Operculodinium centrocarpum* and *Spiniferites elongatus,* and by the gradual increase of the cyst of *Pentapharsodinium dalei* (from ~7 to 37 %). The dominance of these autotrophic species after 7750 cal a BP is reflected by the low H/A ratio (Fig. 4). Assemblage zone 2 was sub-divided into 5 sub-zones, which coincide with

the maximum (sub-zones 2a, 2c and 2e) and minimum (sub-zones 2b and 2d) relative abundance of *O. centrocarpum* (Figs. 3 and S2). Sub-zone 2a (7750-6880 cal a BP; 604-521 cm) is characterized by the sharp increase of *O. centrocarpum*, *S. elongatus* and *Spiniferites ramosus*, the decrease of *I. minutum* and *Brigantedinium* spp. and near disappearance of *Polykrikos* var. Arctic. The transition between sub-zones 2a and 2b (6880-6450 cal a BP; 521-475 cm) is marked by the arrival of *N. labyrinthus*, the low abundance of *O. centrocarpum* and the increase of heterotrophic taxa (*Echinidinium karaense*, *I. minutum* and *Brigantedinium* spp.; Fig. 3). Sub-zone 2c (6450-5100 cal a BP; 475-300 cm) corresponds to the highest abundance of *O. centrocarpum* (up to 67 %), *S. elongatus* (up to 18 %) and *S. ramosus*, and low abundance of all heterotrophic taxa (Fig. 3). It is also characterized by the highest production of organic linings of foraminifera (up to 1500 linings $cm^{-2}$ $yr^{-1}$; Fig. 4). Sub-zone 2d (5100-4750 cal a BP; 300-282 cm) is characterized by the maximum abundance of *E. karaense* and sharp decrease of the gonyaulacacean taxa (*O. centrocarpum*, *S. ramosus*, *S. elongatus* and *N. labyrinthus*; Fig. 3). Finally, sub-zone 2e (4750-3350 cal a BP; 282-164 cm) is characterized by another increase in the abundance of Gonyaulacacean taxa. Overall, reconstructed SST, SSS and primary productivity conditions display a series of 3 pulses of high SST values (sub-zones 2a, 2c and 2e), alternating with 2 short periods of low values (sub-zones 2b and 2d) (Fig. 4). The 3 "warm" pulses are marked by relatively low SIC duration, while the cooler periods are marked by increased SIC, up to 9 months/yr (Fig. 4).

**Zone 3** (3350-700 cal a BP; 164-0 cm) is characterized by the maximum abundance of both the cyst of *P. dalei* (up to 71 %; Fig. 3) and *N. labyrinthus.* It is subdivided into two sub-zones. Sub-zone 3a (3350-2170 cal a BP; 164-106 cm) is dominated by the cyst of *P. dalei* (avg. ~55%), a decrease of *O. centrocarpum* abundance, relatively stable values of the other "North Atlantic" taxa

(*S. elongatus, N. labyrinthus* and *S. ramosus*), and slightly increasing abundances of *I. minutum* and *E. karaense*. Sub-zone 3b (2170-700 cal a BP; 106-0 cm) is characterized by the maximal abundance of cyst of *P. dalei*, an increase of *O. centrocarpum* and a peak in the abundance of *N. labyrinthus*. All heterotrophic taxa are at their lowest abundance in this sub-zone. Sea-surface reconstructions for Zone 3 indicate the highest SST and primary productivity values throughout the core, as well as the lowest values of SIC (Fig. 4).

### *AMD14-Kane2B – Kane Basin*

Dinocysts are less abundant in core Kane2B compared to core 204 and influxes range from extremely low (barren in dinocyst at the base) to 397 cysts $cm^{-2}$ $yr^{-1}$, with maximum values recorded between 4500 and 3000 cal a BP (Fig. 5). Maximum influxes of foraminiferal linings (up to 500 linings $cm^{-2}$ $yr^{-1}$) are recorded between 8100 and 7400 cal a BP (Fig. 6), suggesting higher benthic productivity. Reworked palynomorphs are abundant throughout the core (up to 1200 cells $cm^{-2}$ $yr^{-1}$ at the base of the core, Fig. 6), which suggests high erosion and remobilization of pre-Quaternary sediments. The core is largely dominated by heterotrophic taxa (77-100 %) as reflected by the A/H ratio <1 (Fig. 6). Overall, reconstructed SST ranged from -1.8 to 2.2°C, summer SSS ranged from 27 to 33, SIC between 8 and 10 months per year, and primary productivity between 55 and 120 gC $m^{-2}$ $yr^{-1}$.

**Zone 1**: Prior to 8300 cal a BP (427-298 cm), dinocyst abundance was too low to achieve reconstructions. Nonetheless, the few specimens found were mostly of *Brigantedinium* spp., *E. karaense* and *I. minutum*, suggesting harsh sea-surface conditions, likely associated with a perennial or quasi-perennial SIC. A peak of dinocyst influx at 8100 cal a BP is associated to an increase of *E. karaense* (up to 45 %) and a slight decrease of *Brigantedinium* spp. (Fig. 5).

**Zone 2** (7880-5200 cal a BP; 270-172 cm) is marked by maximum abundances of autotrophic species (*O. centrocarpum* and *S. elongatus*), although the assemblage remains dominated by heterotrophic taxa, such as *Brigantedinium* spp., *E. karaense* and *I. minutum* (Fig. 5 and A/H ratio, Fig. 6). It is subdivided into two sub-zones based on the relative abundances of taxa. Sub-zone 2a (7880-7225 cal a BP; 270-242 cm) is marked by low dinocyst influxes (~60 cysts $cm^{-2}$ $yr^{-1}$) and the appearance of the cyst of *P. dalei*. It is dominated by *Brigantedinium* spp. and accompanied by I. *minutum*, but it also corresponds to the maximal abundance of *O. centrocarpum* (up to 11 %), *S. elongatus* (up to 4 %) and *I. pallidum* (up to 2 %). Sea surface reconstructions indicate maximum values of summer SST and SSS during this period. Sub-zone 2b (7225-5200 cal a BP; 242-172 cm) is dominated by *E. karaense* (up to 49 %) and *I. minutum* (up to 41 %). It is also characterized by the highest relative abundance of *Polykrikos* var. Arctic (up to 5 %), an Arctic species restricted to regions characterized by cold surface waters (Zonneveld *et al.*, 2013). Sea surface reconstruction revealed relatively low summer SST (~1°C), SSS (~27psu) and primary productivity (~60 gC $m^{-2}$ $yr^{-1}$), and an extended SIC during this period (Fig. 6).

**Zone 3** (5200 cal a BP to present; 172-0 cm) is characterized by *Brigantedinium* spp. and a progressive increase of the relative abundance of *I. minutum* (20 to 47 %). A change occurred around ca. 2100 cal a BP, which leads to the subdivision into 2 sub-zones. Sub-zone 3a (5200-2100 cal a BP; 172-45 cm) is characterized by the maximum relative abundance of *Brigantedinium* spp. (up to 57 %), low abundance of *E. karaense* (~15 %), and maximum dinocyst influxes (4500 to 3000 cal a BP). Sub-zone 3b (2100 cal a BP to present; 45-0 cm) is characterized by the maximum abundance of *I. minutum* (up to 48 %) and by increasing abundances of *E. karaense*, *Polykrikos* var. Arctic and cyst of *P. dalei* (Fig. 5). Sea-surface reconstructions reveal a slight increase in SIC and a slight decrease in SST and primary productivity during this period (Sub-zone 3b, Fig. 6).

# Discussion

Dinocyst assemblages from two sites (204 and Kane2B) reflect important oceanographic changes throughout the Holocene. The results indicate a shift from cold and extended SIC to milder postglacial conditions with seasonal SIC that occurred around 7880 cal a BP in the Kane Basin and around 7750 cal a BP in the northeast Baffin Bay. The synthesis on the timing of climatic and environmental changes documented from both cores are presented on Figs. 7 and 8. Results from core 204 (Fig. 7) are compared with regional record from Disko Bugt (MSM343300: Ouellet-Bernier *et al.*, 2014), while results from core Kane2B (Fig. 8) are compared with regional records from Smith Sound (Sites 008 and 012: Levac *et al.*, 2001) and Lancaster Sound (2004-804-009: Ledu *et al.*, 2008).

## *Early Holocene - End of deglaciation*

In both cores, assemblages from this period (prior to 7750 and 7880 cal a BP, respectively) are dominated by *Brigantedinium* spp., *I. minutum*, *E. karaense* and accompanied by *Polykrikos* var. Arctic. The co-dominance of those taxa is often associated with harsh, densely ice-covered conditions and relatively high productivity (de Vernal *et al.*, 2001, 2013a; Head *et al.*, 2001; Rochon *et al.*, 1999; Zonneveld *et al.*, 2013). Low dinocyst influxes and the dominance of these heterotrophic species suggest primary productivity probably dominated by diatoms, which constitute the main food source of heterotrophic dinoflagellates (e.g., Jacobson and Anderson, 1986). Diatoms are usually advantaged (opportunists) compared to dinoflagellates in areas where nutrients are relatively abundant (e.g., Dale, 1996; Lovejoy *et al.*, 2002).

At the site of core 204, the sea surface reconstructions revealed the lowest SST values (~2°C in summer, ~-1.4°C in winter) and the highest SIC duration (up to 9 months/year) during this period, although associated to minimal values of productivity (Fig. 4). Harsh sea-surface conditions likely

prevented the development of phototrophic dinocysts taxa in this area, while primary producers such as diatoms may account for the presence of heterotrophic dinocysts taxa. Previous studies (e.g., Ouellet-Bernier *et al.*, 2014; Gibb *et al.*, 2015) in the region of Disko Bugt have also observed dinocyst assemblages dominated by heterotrophic taxa prior to 7300 and 7400 cal a BP respectively. Besides, Jennings *et al.* (2014) determined a continuous GIS mass loss until 7800 cal a BP based on foraminifera and other proxies. Thus, the relatively low reconstructed salinity values in core 204 support evidence of high meltwater discharge from the GIS prior to 7750 cal a BP, also suggested in Caron *et al.* (2019).

In Kane Basin, immediately after the retreat of grounded ice from the core site, the high concentration of reworked palynomorphs (Fig. 6) associated to a high IRD content support ice-proximal conditions with high meltwater discharge from both IIS and GIS (Caron *et al.*, 2019; Georgiadis *et al.*, 2018). This strong meltwater input, linked to the retreating ice sheets, probably resulted in very harsh sea surface conditions with a quasi-permanent SIC during this period (prior to 8300 cal a BP; Ledu *et al.*, 2010), which would explain the very low concentration of dinocysts observed during this period. Moreover, we consider that this period before 8300 cal a BP is linked to a major oceanographic reorganization associated with the opening of Nares Strait and establishment of the modern oceanic circulation (e.g., Jennings *et al.*, 2011, 2019; Georgiadis *et al.*, 2018).

### *Onset of full interglacial conditions*

On the Northwestern Greenland margin (core 204), a major shift is recorded around 7750 cal a BP, representing a change in dominant species from heterotrophic to phototrophic taxa. It was also associated with a rapid increase of dinocyst influxes and species richness, hence suggesting milder conditions with warmer SST and a reduced SIC during this period (Figs. 3 and 7). Several regional

records (based on dinocysts or foraminifera) have also observed a major shift due to surface and subsurface warming occurring between 7400-7800 cal a BP, which is considered as the transition towards modern postglacial conditions and warmer sea-surface conditions (e.g., Perner *et al.*, 2013; Ouellet-Bernier *et al.*, 2014; Jennings *et al.*, 2014; Gibb *et al.*, 2015). Indeed, the establishment of warmer conditions in Baffin Bay is likely a response to the reinforcement of the warm and saline component (IC) of the WGC, linked with the final retreat of both LIS and GIS (Lloyd *et al.*, 2005; Ouellet-Bernier *et al.*, 2014; Gibb *et al.*, 2015). In addition, the dominance of the phototrophic species *O. centrocarpum* and *S. elongatus* is often associated with the North Atlantic Drift, suggesting surface water temperatures > 0°C throughout the year (Rochon *et al.*, 1999; Mudie *et al.*, 2006; Zonneveld *et al.*, 2013).

In Kane Basin, this transition towards interglacial conditions is also marked by the increase of accompanying autotrophic species *O. centrocarpum* and *S. elongatus* from 7880 to 7200 cal a BP, although the assemblage is still dominated by heterotrophic taxa (Fig. 5). This suggests warmer sea-surface conditions, which nonetheless remained relatively cold with an extended SIC, as expected considering the presence of the strong cold Arctic current flowing southwards in the basin (Münchow *et al.*, 2006). This is coherent with Levac *et al.* (2001), whom suggest sea-surface temperatures up to 3°C higher than at present as early as 7800 $^{14}$C a BP (~8000 cal a BP) in Smith Sound (Fig. 8). As *O. centrocarpum* and *S. elongatus* are considered as Atlantic water indicators, these results could reflect the entry of warmer Atlantic water into the basin (via the IC) associated with the enhanced strength of the WGC during this period (Perner *et al.*, 2013), the insolation maximum (Lecavalier *et al.*, 2017) and isostatic readjustment after the retreat of grounded ice (England *et al.*, 2006).

### *Middle Holocene - Holocene Thermal Maximum*

Dinocyst assemblages of core 204 indicate that warm postglacial conditions were established in northeastern Baffin Bay after 7750 cal a BP, although two periods (6880-6450 and 5100-4750 cal a BP) are marked by the decrease of *O. centrocarpum* abundance (warm water taxa) and increase of the abundance of *E. karaense* and *I. minutum* (cold taxa; Fig. 3). These two periods are associated with relatively lower reconstructed SST and productivity values and a higher SIC duration (Fig. 4). These results suggest a temporary return to colder sea-surface conditions during these periods, possibly caused by meltwater pulses from the GIS, also evidenced by moderate peak abundance of *Halodinium* sp. at the top of both sub-zones. Overall, we associate the period from 7750 to 3350 cal a BP to the Holocene Thermal Maximum (HTM), with optimal conditions observed between 6400 and 5500 cal a BP (Fig. 7). The high dinocyst influx, as well as the relatively high reconstructed SST and productivity values support this interpretation (Figs. 4 and 7), which is also consistent with previous regional works (e.g., Ouellet-Bernier *et al.*, 2014; Jennings *et al.*, 2014). Furthermore, a higher influence of the IC (stronger, saltier and warmer WGC) was recorded between 6200 and 3500 a BP by Perner *et al.* (2013).

As mentioned above, our data from Kane Basin suggest warmer conditions between 7880 and 7200 cal a BP, with the highest relative abundance of 'North Atlantic' indicators. This period is also characterized by maximum influxes of foraminiferal linings (high benthic productivity) and very low dinocyst influxes (Fig. 8). Regional studies (Bradley, 1990; Levac *et al.*, 2001; Knudsen *et al.*, 2008; Jennings *et al.*, 2011; Lecavalier *et al.*, 2017) have recorded an early HTM in this area compared to lower latitudes sites (e.g., Disko Bugt or Labrador Sea). Likewise, Ledu *et al.* (2010) observed an increase of the phototrophic taxa relative abundance, linked with the HTM between 8500 and 5500 cal a BP. Nevertheless, the low dinocyst influxes and high benthic productivity

suggest that dinoflagellates may not be solely responsible for the primary productivity in the Nares Strait area, and that diatoms may be involved (Cormier *et al.*, 2016). Thus, our results could reflect relatively open water conditions without ice arch formation during this period, favorable for the development of phototrophic dinocysts species and diatoms.

### *Late Holocene changes in sea surface conditions*

Regionally, the Late Holocene is characterized by a cooling trend which corresponds to the Neoglacial period associated with glacier growth and subsurface cooling along the West Greenland coast (e.g., Briner *et al.*, 2016; Perner *et al.*, 2013; Schweinsberg *et al.*, 2017). This neoglacial cooling period is recorded over the last ca. 4000 cal a BP in regional reconstructions (e.g., Knudsen *et al.*, 2008; Seidenkrantz *et al.*, 2008; Krawczyk *et al.*, 2010; Andresen *et al.*, 2010). The fluctuations observed in the reconstructed sea-surface conditions at both our sites over the last ca. 3500 cal a BP reflect this cooling to a certain extent after ~2000 cal a BP.

However, after 3350 cal a BP, core 204 is marked by a major change in the dinocyst assemblages with a dominance shift from *O. centrocarpum* to cyst of *P. dalei* (Fig. 3), suggesting changes in surface water conditions during this period, which were not observed in other regional dinocyst records (e.g., Ouellet-Bernier *et al.*, 2014; Gibb *et al.*, 2015). In the Arctic, cysts of *P. dalei* are associated with relatively warm, stratified and highly productive waters, and generally require SST higher than 4°C (Rochon *et al.*, 1999; Matthiessen *et al.*, 2005; Ribeiro *et al.*, 2012). Moreover, high abundances of cysts of *P. dalei* are recorded in location with large seasonal temperature gradient (Rochon et al., 1999). Thus, the dominance of this species, which is found in abundance near Iceland (Zonneveld *et al.*, 2013) could reflect a higher seasonal gradient during this period, possibly from an increased contribution of the Atlantic component (IC) to the WGC. Indeed,

warmer SST (increase by ~1.5°C) after ~3.1 cal ka BP were also observed by Gibb *et al.* (2015) and was attributed to an enhanced strength of the IC.

Reconstructions from Kane Basin reflect the abundance of the dominant taxa: *Brigantedinium* spp. for the productivity and SST, and *I. minutum*+*E. karaense* for the SIC duration. We interpret the maximum dinocyst influxes and dominance of *Brigantedinium* spp. between 5200 and 2100 cal a BP to reflect an increased primary productivity linked to the availability of pelagic prey, mostly diatoms (Jacobson and Anderson, 1986; Matthiessen *et al.*, 2005; Cormier *et al.*, 2016), but not necessarily warmer conditions. Besides, we associate the increasing trend of *I. minutum* after 5200 cal a BP to a progressive cooling, possibly linked to the Neoglaciation. This cooling trend intensified over the last 2100 cal a BP (increase of *I. minutum*, *E. karaense* and *Polykrikos* var. Arctic*)* with decreasing dinocyst influxes (Figs. 5 and 8). Indeed, the relative abundance of *I. minutum* increases linearly with the duration of seasonal SIC in Arctic regions (Zonneveld *et al.*, 2013). This is coherent with the work of Levac *et al.* (2001) that observed a general deterioration of sea-surface conditions over the last ~3200 cal a BP in Smith Sound (increased SIC; Fig. 8). Finally, Ledu *et al.* (2008) also recorded a cooling trend in Lancaster Sound starting after 5500 cal a BP. Thus, we suggest that the increase of *I. minutum* and *E. karaense* observed at the top of core (last 2100 cal a BP) could reflect this regional cooling trend of the late Holocene.

## Conclusions

Paleoclimatic reconstructions based on dinocyst assemblages from two sediment cores (AMD14-204 and AMD14-Kane2B) using the MAT provided insight on the evolution of sea-surface conditions during the Holocene. In the northeastern Baffin Bay (core 204), the dinocyst assemblages revealed important changes throughout the Holocene, with harsh conditions prior to 7750 cal a BP, associated to an extended SIC and high meltwater inputs from the retreating GIS.

A shift to warmer conditions occurred around 7750 cal a BP, as a response to the final retreat of the GIS and a strengthened North Atlantic component (IC) of the WGC. From 7750 to 3350 cal a BP, the dominance of phototrophic taxa and highest species richness marks the mid-Holocene thermal optimum associated with warmer SST, higher productivity and reduced SIC, culminating between 6450 and 5100 cal a BP. Finally, after 3350 cal a BP, the progressive establishment of modern dinocyst assemblages is observed, associated with higher seasonality.

In Kane Basin (core Kane2B), our results revealed the dominance of heterotrophic taxa throughout the core. Results suggest relatively cold sea-surface conditions, with abundant SIC during most of the Holocene in this area. Nonetheless, variations of the dinocyst assemblages reveal some changes in the sea-surface conditions following the opening of Nares Strait (and the strengthened southward flow of cold Arctic water through Nares Strait). Enhanced contribution of Atlantic water recorded between 7880 and 7200 cal a BP, marked by the occurrence of autotrophic taxa, suggest oceanographic conditions different than present, likely caused by the insolation maximum and breakup of the IIS. Maximum primary productivity recorded between 5200 and 2100 cal a BP reflects the availability of prey for heterotrophic dinoflagellates. Finally, an increasing abundance of the Arctic taxa *I. minutum* is associated with a progressive cooling trend, which intensified after 2100 cal a BP.

Overall, this study indicates that sea-surface conditions in northeastern Baffin Bay are more favorable for the growth of autotrophic dinoflagellates than in Nares Strait. This is explained by the influence of the warm and relatively salty WGC on the eastern side of Baffin Bay compared to the strong southward Arctic flow in Nares Strait. Finally, this study confirmed the important changes in sea-surface conditions that occurred from the end of the deglaciation until present, both for the Northwestern Greenland shelf and Nares Strait. These variations can be associated to

changes in meltwater discharges or varying strength of oceanic currents (e.g., IC or EGC influence) themselves related to the main climatic changes of the Holocene.

*Abbreviations.* BIC, Baffin Island Current; CAA, Canadian Arctic Archipelago; EGC, East Greenland Current; GIS, Greenland Ice Sheet; HTM, Holocene Thermal Maximum; IC, Irminger Current; IIS, Innuitian Ice Sheet; IRD, ice-rafted debris; LGM, Last Glacial Maximum; LIS, Laurentide Ice Sheet; MAT, modern analogue technique; RDL, rapidly-deposited layer; SIC, Sea-ice cover; SST, Sea-surface temperature; SSS, Sea-surface salinity; WGC, West Greenland Current.

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## Table captions

**Table 1.** Coordinates of the two cores used in this study.

| Core ID | Region | Latitude (°N) | Longitude (°W) | Water depth (m) | Core length (cm) |
|---|---|---|---|---|---|
| AMD14-204 | Upernavik Isstrøm | 73°15.663’ | 57°53.987’ | 987 | 734 |
| AMD14-Kane2B | Kane Basin | 79°30.908’ | 70°49.742’ | 220 | 425 |

**Table 2.** List of the dominant dinoflagellate cyst taxa from the two cores location, shown in Fig. 2. G=Gonyaulacales, P=Peridiniales, A=Autotrophic and H=Heterotrophic taxa.

| Dinoflagellate taxa | | Trophic mode | Grouped as |
|---|---|---|---|
| *Operculodinium centrocarpum* sensu Wall and Dale | G | A | *Operculodinium centrocarpum* |
| *O. centrocarpum* short processes | G | A | *Operculodinium centrocarpum* |
| *O. centrocarpum* Arctic | G | A | *Operculodinium centrocarpum* |
| *Nematosphaeropsis labyrinthus* | G | A | |
| *Spiniferites elongatus* | G | A | |
| *Spiniferites ramosus* | G | A | |
| Cyst of *Pentapharsodinium dalei* | P | A | |
| *Impagidinium pallidum* | G | A | |
| *Islandinium minutum* | P | H | |
| *Islandinium brevispinosum* | P | H | *Islandinium minutum* |
| *Islandinium? cezare* | P | H | *Islandinium minutum* |
| *Echinidinium karaense* | P | H | |
| *Brigantedinium* spp. | P | H | |
| *Brigantedinium simplex* | P | H | *Brigantedinium* spp. |
| *Brigantedinium cariacoense* | P | H | *Brigantedinium* spp. |
| *Polykrikos* var. Arctic | P | H | |
| *Polykrikos quadratus* | P | H | *Polykrikos* var. Arctic |

**Figures legends**

**Figure 1:** a) Map of Baffin Bay and Nares Strait indicating sampling location for cores AMD14-Kane2B (Kane Basin) and AMD14-204 (Upernavik cross-shelf trough), as well as cores mentioned in this study (MSM343300, sites 008 and 012, LS009). Regional surface oceanic circulation: BIC= Baffin Island Current (Arctic water in blue); WGC = West Greenland Current (in red; mix of IC = Irminger Current, North Atlantic water and EGC = East Greenland Current). NS: Nares Strait; MB: Melville Bay; DB: Disko Bugt; DS: Davis Strait; LS: Lancaster Sound; SS: Smith Sound; NOW: North Water Polynya. b) Satellite image of the study area indicating the sea ice cover (image from 22 May 2018; credit: NASA Worldview: https://worldview.earthdata.nasa.gov).

**Figure 2:** Photoplate of main dinocyst taxa and palynomorphs (foraminiferal lining) identified in this study: a) *Operculodinium centrocarpum* sensu Wall and Dale; b) Cyst of *Pentapharsodinium dalei*; c) *Nematosphaeropsis labyrinthus*; d) *Spiniferites elongatus*; e) *Spiniferites ramosus*; f) *Impagidinium pallidum*; g) *Islandinium minutum*; h) *Echinidinium karaense*; i) *Polykrikos* var. Arctic*; j) Brigantedinium* simplex; k) foraminiferal lining; l) *O centrocarpum* (left) + *I. minutum* (right).

**Figure 3:** Dinocyst influx (x$10^2$ cysts $cm^{-2}$ $yr^{-1}$), relative abundance (%) of the main dinocyst taxa, and dinocyst assemblage zones in core AMD14-204. Data are represented with the thin dotted dark grey lines, 3-points running mean are represented by green (autotrophs) or brown (heterotrophs) lines.

**Figure 4:** Reconstructions of winter and summer salinity (SSS) and sea-surface temperature (SST), seasonal sea-ice cover in months/year, and productivity (in gC $m^{-2}$ $yr^{-1}$) based on the modern analogue technique (MAT) applied to dinocyst assemblages in core AMD14-204. Reconstructions are shown with the black lines, the 3-points running mean are represented by bold colored lines and the maximum and minimum possible values according to the set of five analogues are shown in light gray. Modern values are indicated by a black arrow at the top. Heterotrophs/Autotrophs (H/A) ratio and influxes of organic linings of foraminifera, *Halodinium* sp. and reworked palynomorphs are displayed on the right. Statistical minimum and maximum distances between fossil and modern analogues are 0.12 and 0.20 respectively.

**Figure 5:** Dinocyst influx (x$10^2$ cysts $cm^{-2}$ $yr^{-1}$), relative abundance (%) of the main dinocyst taxa, and dinocyst assemblage zones in core AMD14-Kane2B. Data are represented with the thin dotted dark grey lines, 3-points running mean are represented by green (autotrophs) or brown (heterotrophs) lines. In samples containing sparse assemblages (counts <50 cysts), occurrences are reported by dots on the diagram.

**Figure 6:** Reconstructions of winter and summer salinity (SSS) and sea-surface temperature (SST), seasonal sea-ice cover in months/year, and productivity (in gC $m^{-2}$ $yr^{-1}$) based on modern analogue technique (MAT) applied to dinocyst assemblages in core AMD14-Kane2B. Reconstructions are shown with the black lines, the 3-points running means are represented by colored lines and the

maximum and minimum possible values according to the set of five analogues are shown in light gray. Modern values are indicated by a black arrow at the top. Autotrophs/Heterotrophs (A/H) ratio and influxes of organic linings of foraminifera, *Halodinium* sp. and reworked palynomorphs are displayed on the right. Statistical minimum and maximum distances between fossil and modern analogues are 0.08 and 0.11 respectively.

**Figure 7:** Comparison of the palynological records from core AMD14-204 (this study) with a) the $\delta^{18}O$ of Camp Century (Vinther *et al.*, 2009), b-c: 'North Atlantic' taxa (*O. centrocarpum*, *S. elongatus*, *S. ramosus)* relative abundance of (a) core MSM343300, Disko Bugt (Ouellet-Bernier *et al.*, 2014) and (b) core 204, d-e: 'Polar' taxa (*I. minutum*, *E. karaense*, *Polykrikos* var. Arctic*)* relative abundance (note reversed scale) of (d) core MSM343300 and (e) core 204, f) dinocysts and foraminiferal linings influxes ($cm^{-2}$ $yr^{-1}$) of core 204 and g) reconstructed SST anomaly (°C) compared to modern value for core 204. The Holocene Thermal Maximum is delimited by the two dashed lines, and blue shaded areas indicate the two cold pulses (zones 2b and 2d).

**Figure 8:** Comparison of the palynological records from core AMD14-Kane2B (this study) with a) the $\delta^{18}O$ temperature reconstructions from the Agassiz Ice Cap (Vinther *et al.*, 2009), b) SIC duration (months/year) from sites 008 and 012 (Levac *et al.*, 2001), c) Gonyaulacal/Peridinoid (G/P) ratio from site LS009 (Ledu *et al.*, 2008), d) 'warm water' taxa (**O. centrocarpum*, *S. elongatus*) relative abundance from sites 008 and 012, e) foraminiferal lining influxes $cm^{-2}$ $yr^{-1}$, f) *O. centrocarpum* relative abundance, g) *I. minutum* relative abundance, h) dinocyst influxes in cyst $cm^{-2}$ $yr^{-1}$. Data from sites 008, 012 and LS009 (core 2004-804-009) were downloaded from the GEOTOP dinocyst database (de Vernal *et al.*, 2013a).

## Supplementary material

Additional supporting information:

**Fig. S1.** High-resolution images, CT-Scan, age models and sedimentary units for cores AMD14-204 and AMD14-Kane2B. Radiocarbon dates are from Georgiadis *et al.* (2018) for Kane2B and from Caron *et al.* (2019) for core 204; paleomagnetic stratigraphic markers from Caron *et al.* (2019) and organic matter is from Giraudeau *et al.* (under revision). RDL: Rapidly deposited layer.

**Fig. S2.** Results of the principal component analysis as a function of depth (cm) for cores AMD14-204 and AMD14-Kane2B. PC scores are given at the top of each graph, which represent >95% total variance. Dark grey shade (300-320 cm) in core Kane2B corresponds to the RDL (Rapidly Deposited Layer) which marks the final opening of Nares Strait (Georgiadis *et al.*, 2018).

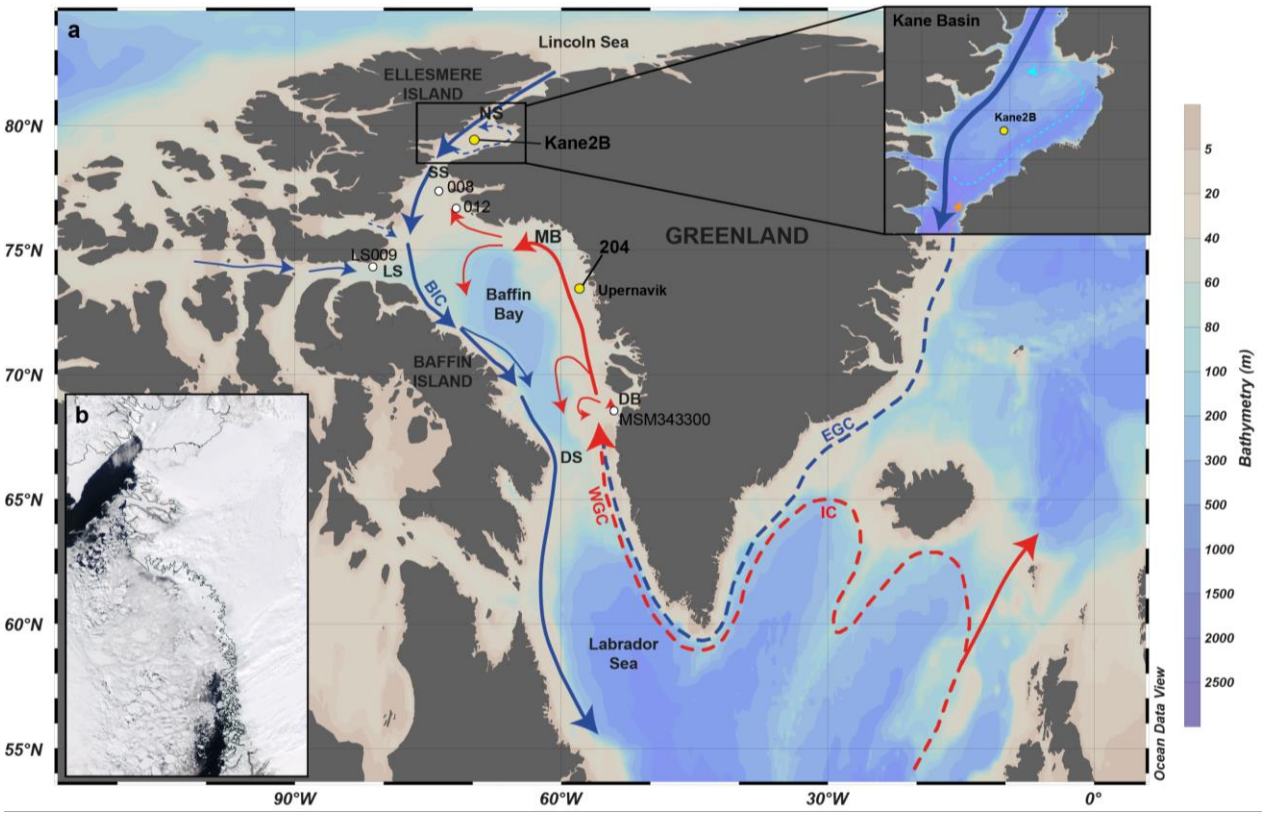
a
b
Lincoln Sea
ELLESMERE ISLAND
NS
Kane2B
SS
008
012
MB
204
GREENLAND
LS009
LS
BIC
Baffin Bay
Upernavik
BAFFIN ISLAND
DB
MSM343300
DS
WGC
EGC
IC
Labrador Sea
Kane Basin
Kane2B
80°N
75°N
70°N
65°N
60°N
55°N
90°W
60°W
30°W
0°
Bathymetry (m)
5
20
40
60
80
100
200
300
500
1000
1500
2000
2500
Ocean Data View

**Figure 1**

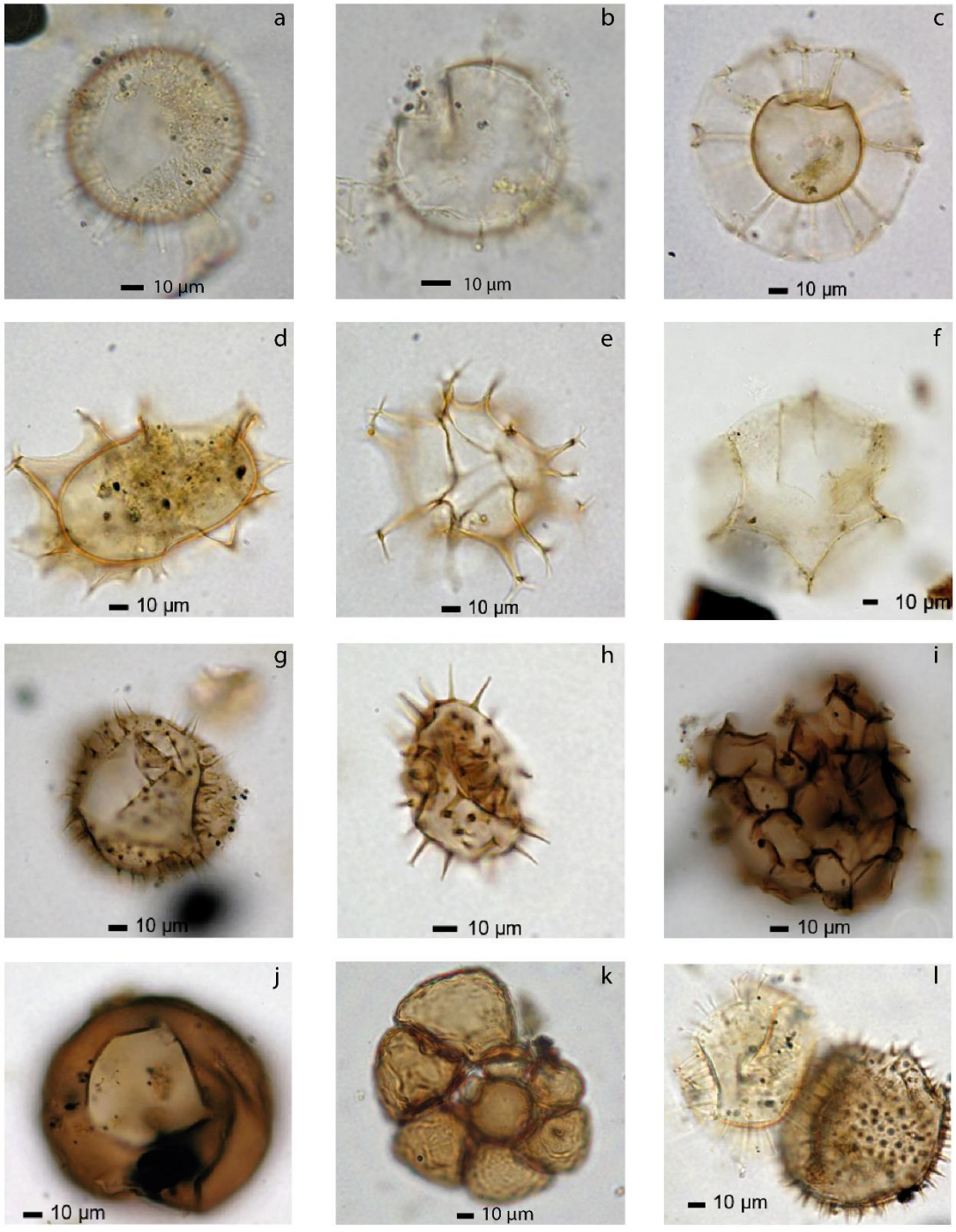
a
10 µm
b
10 µm
c
10 µm
d
10 µm
e
10 µm
f
10 µm
g
10 µm
h
10 µm
i
10 µm
j
10 µm
k
10 µm
l
10 µm

**Figure 2**

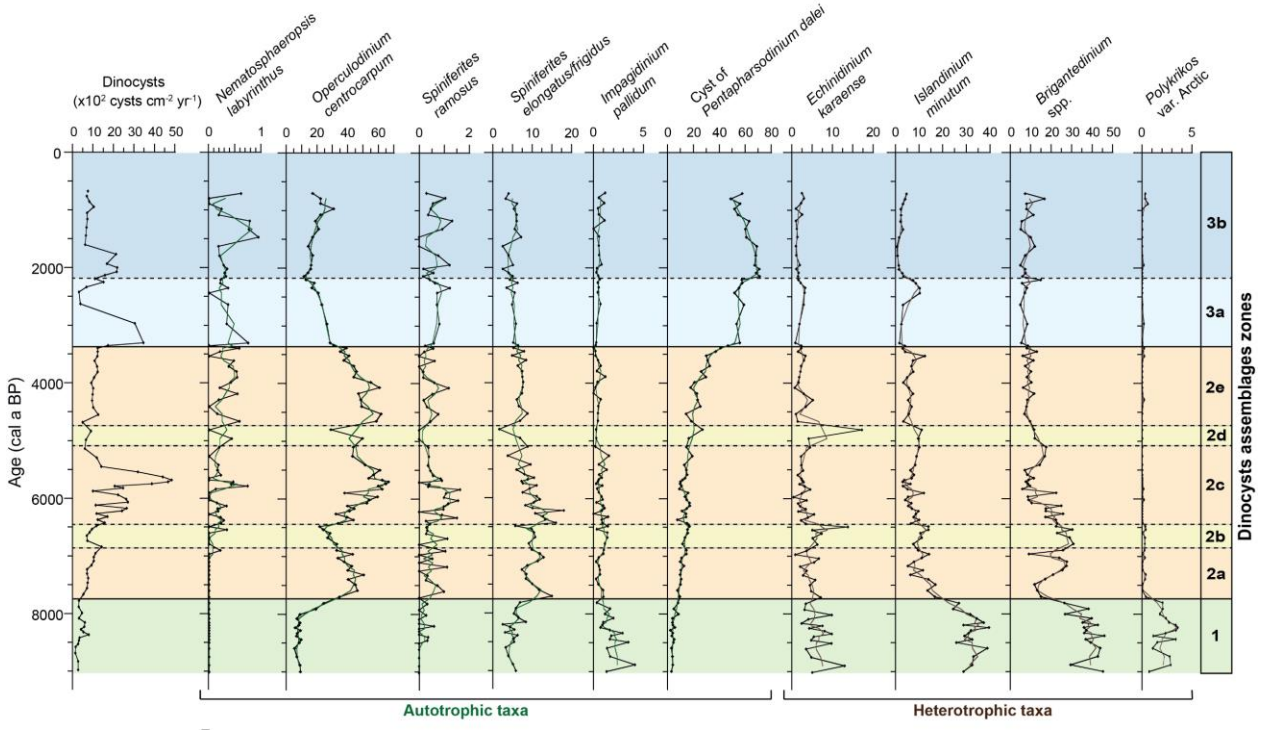


**Figure 3**

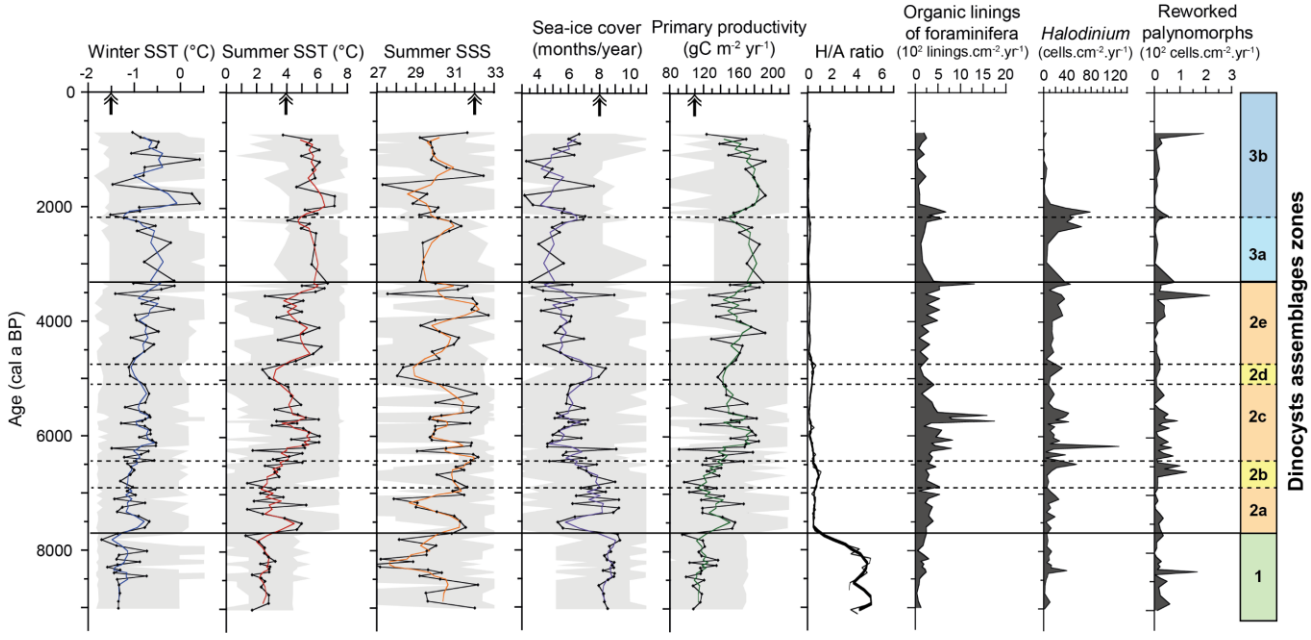


**Figure 4**

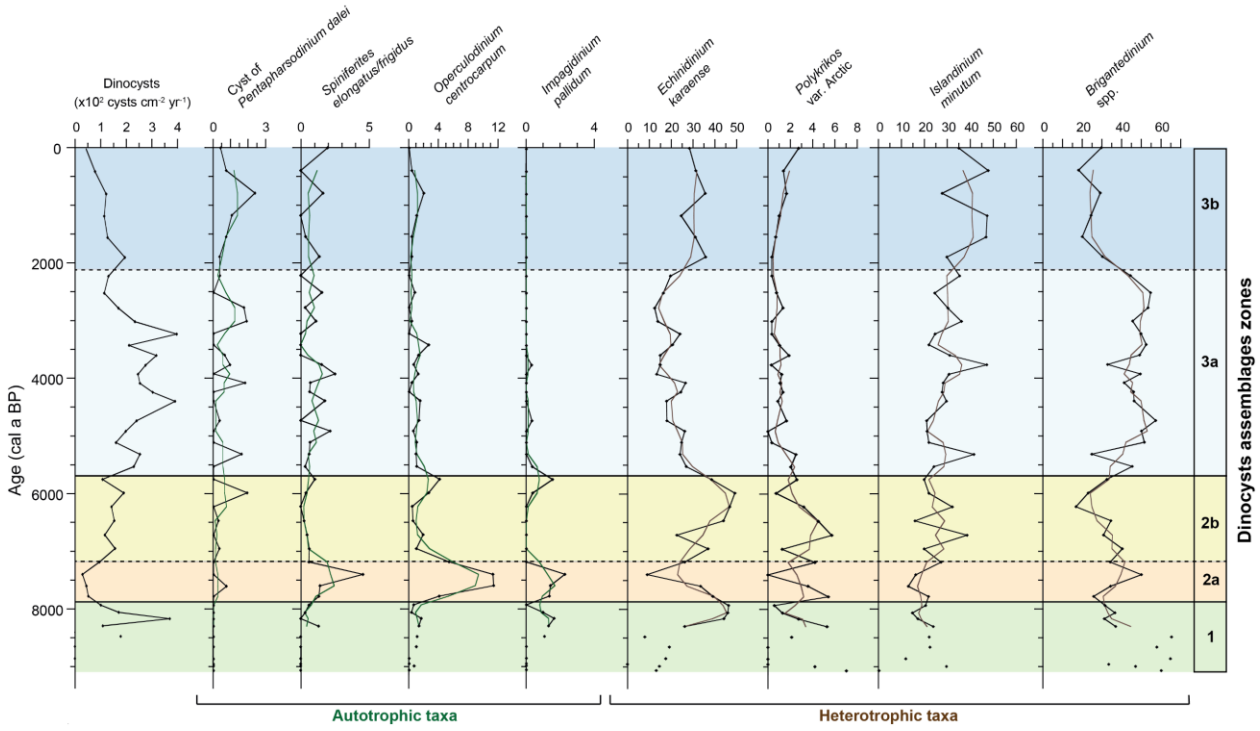


**Figure 5**

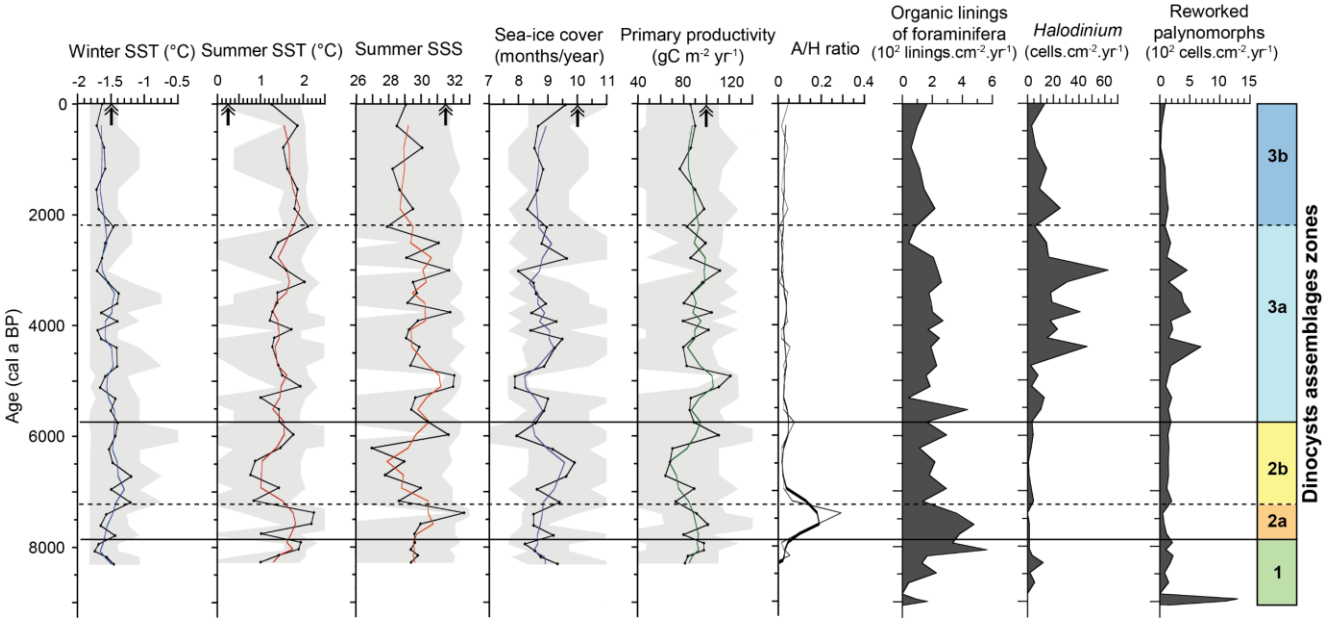


**Figure 6**

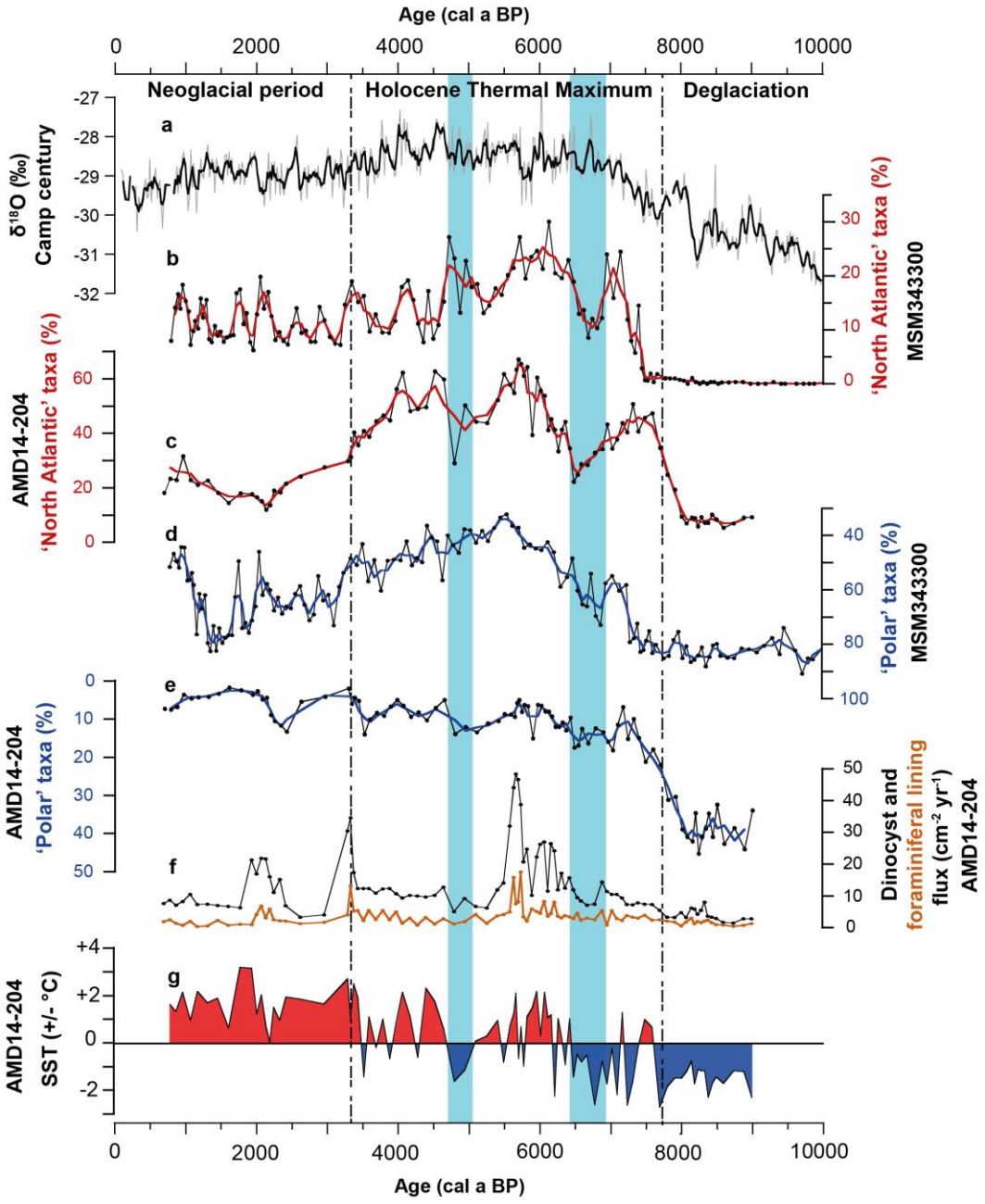
Age (cal a BP)
0
2000
4000
6000
8000
10000
Neoglacial period
Holocene Thermal Maximum
Deglaciation
a
b
c
d
e
f
g
δ18O (‰)
Camp century
-27
-28
-29
-30
-31
-32
'North Atlantic' taxa (%)
MSM343300
30
20
10
0
AMD14-204
'North Atlantic' taxa (%)
60
40
20
0
'Polar' taxa (%)
MSM343300
40
60
80
100
AMD14-204
'Polar' taxa (%)
0
10
20
30
40
50
Dinocyst and foraminiferal lining flux (cm-2 yr-1)
AMD14-204
50
40
30
20
10
0
AMD14-204
SST (+/- °C)
+4
+2
0
-2
Age (cal a BP)

**Figure 7**

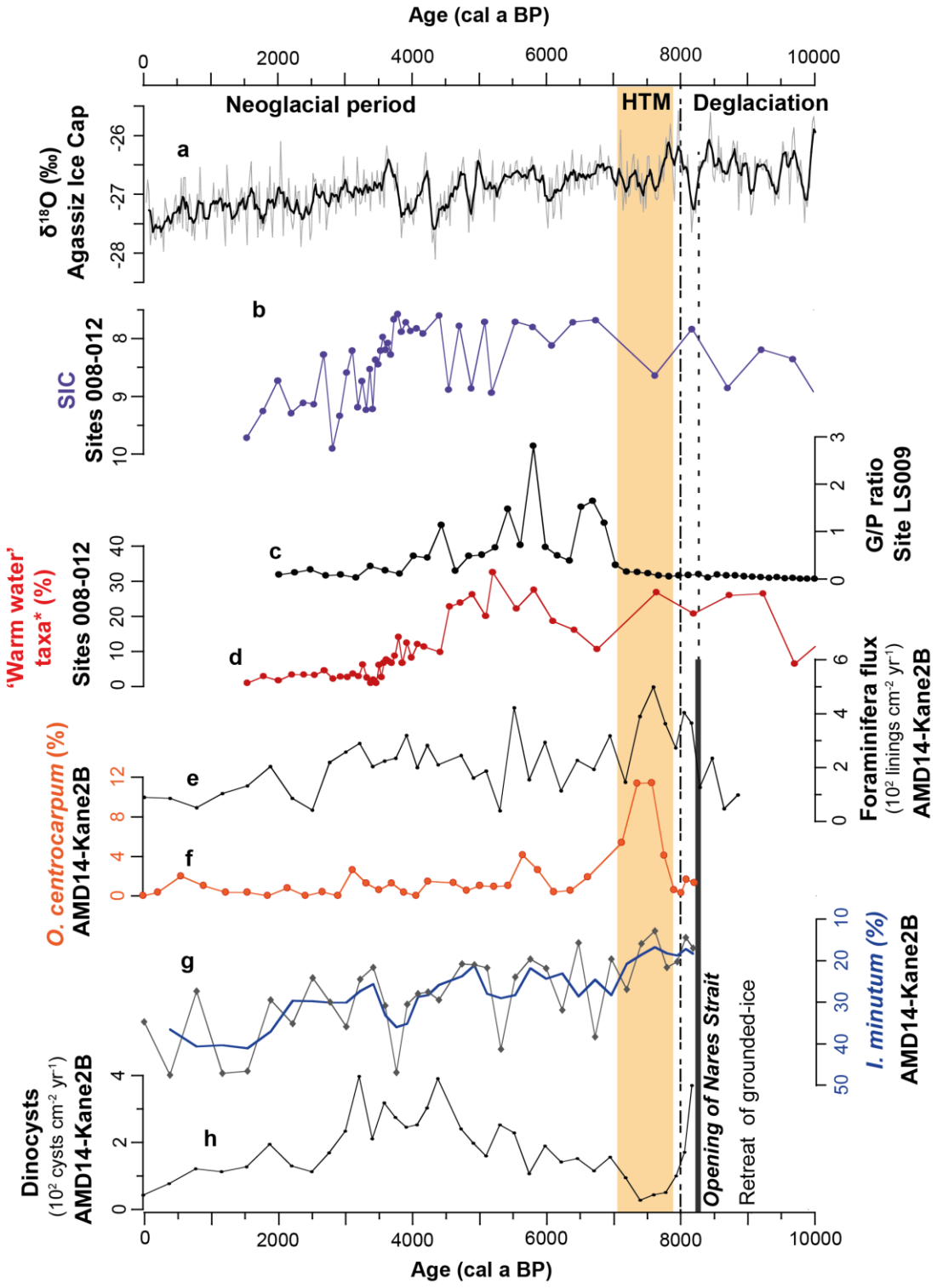
Age (cal a BP)
0
2000
4000
6000
8000
10000
Neoglacial period
HTM
Deglaciation
δ18O (‰)
Agassiz Ice Cap
-26
-27
-28
a
SIC
Sites 008-012
b
8
9
10
'Warm water'
taxa* (%)
Sites 008-012
c
d
0
10
20
30
40
G/P ratio
Site LS009
0
1
2
3
O. centrocarpum (%)
AMD14-Kane2B
e
f
0
4
8
12
Foraminifera flux
(10² linings cm⁻² yr⁻¹)
AMD14-Kane2B
0
2
4
6
I. minutum (%)
AMD14-Kane2B
10
20
30
40
50
g
Dinocysts
(10² cysts cm⁻² yr⁻¹)
AMD14-Kane2B
h
0
2
4
Opening of Nares Strait
Retreat of grounded-ice
Age (cal a BP)

**Figure 8**

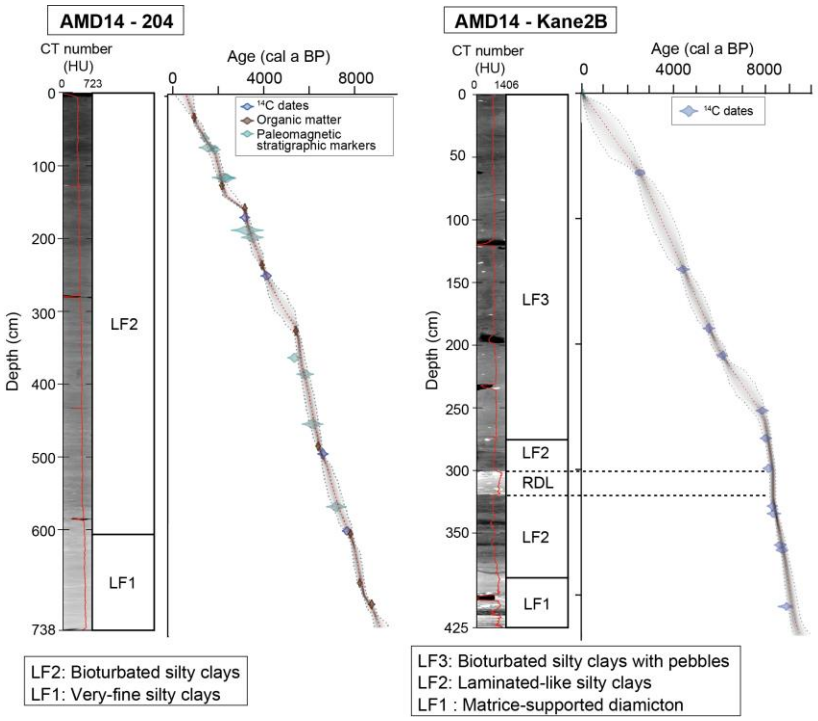


**Figure S1**

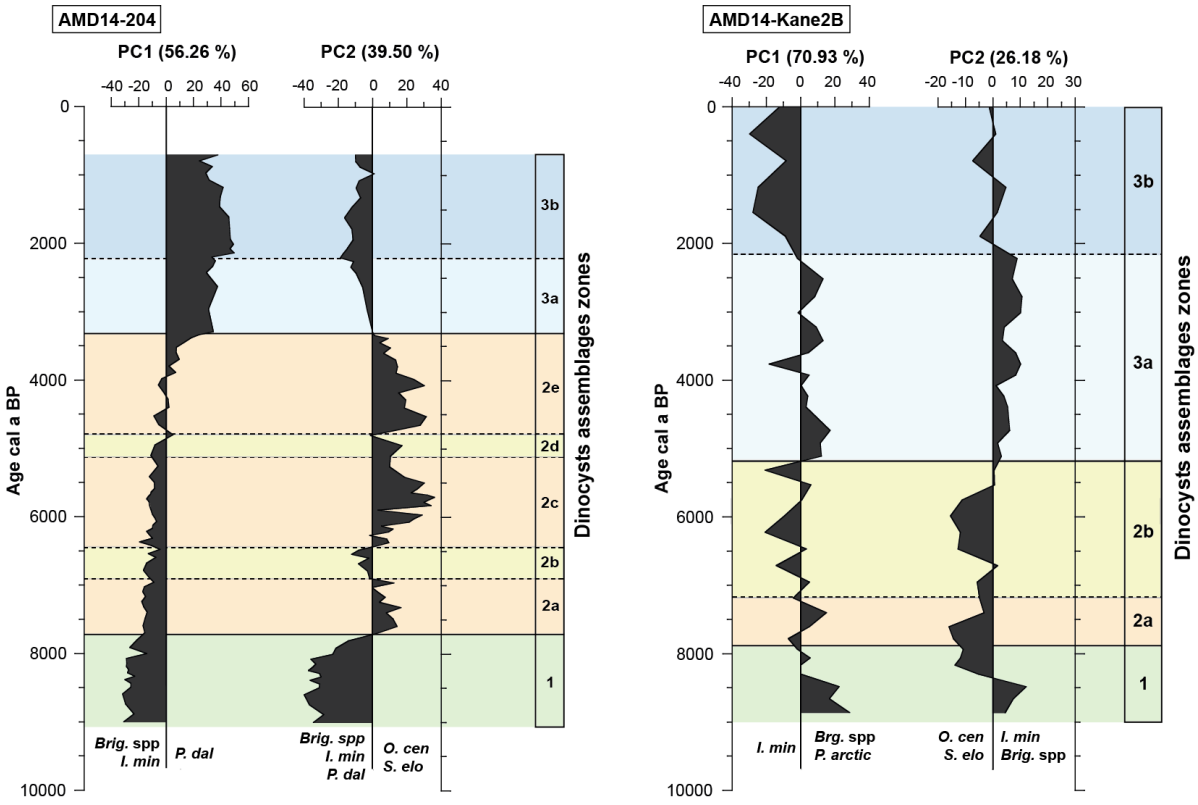


**Figure S2**